

# Pertinent Meteorological Data for Hurricane Allen of 1980 

Silver Spring, Md.

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WB 6 Climatology of Atlantic Tropical Storms and Hurricanes. M. A. Alaka, Techniques Development Laboratory, May $1968,18 \mathrm{p}$.

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WB 9 Weekly Synoptic Analyses, 5-, 2-, and $0.4-\mathrm{Mb}$ Surfaces for 1966 (based on meteorological rocketsonde and high-level rawingonde observations). Staff, Upper Air Branch, National Meteorological Center, January 1969, 169 p .

WB 10 Hemispheric Teleconnections of Mean Circulation Anomalies at 700 Millibars. James F. O'Connor, National Meteorological Center, February 1969, 103 p.

WB 11 Monthly Mean 100-, 50-, 30-, and l0-Millibar Charts and Standard Deviation Maps, 1966-1967. Staff, Upper Alr Branch, National Meteorological Center, April 1969, 124 p.

WB 12 Weekly Synoptic Analyses, 5-, 2-, and 0.4-Millibar Surfaces for 1967. Staff, Upper Air Branch, National Meteorological Center, January 1970, 169 p.

## NOAA Techaical Reports

NWS 13 The March-April 1969 Snowmelt Floods in the Red River of the North, Upper Mississippi, and Missouri Basins, Joseph L. H. Paulhus, Office of Hydrology, October 1970, 92 p. (COM-71-50269)

NWS 14 Weekly Synoptic Analyses, 5-, 2-, and 0.4-Millibar Surfaces for 1968. Staff, Upper Air Branch, National Meteorological Center, May 1971, 169 p. (Com-71-50383)

NWS 15 Some Climatological Characteristics of Hurricanes and Tropical Storms, Gulf and East Coasts of the United States. Francis P. Ho, Richard W. Schwerdt, and Hugo V. Goodyear, May 1975, 87 p. (COM-75-11088)

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Frances P. Ho and John F. Miller

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## U.S. DEPARTMENT OF COMMERCE <br> Malcolm Baidrige, Secretary

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National Weather Service
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# PERTINENT METEOROLOGICAL DATA FOR HURRICANE ALLEN OF 1980 

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#### Abstract

Al1 available meteorological data for Hurricane Allen, 1980 , have been analyzed to provide information as accurate as possible for use in dynamic storm surge models. Detailed analyses are presented of the storm track, forward speed, central pressure, and radius to maximum wind. Particular attention is given to the period surrounding landfall.


## 1. INTRODUCTION

The purpose of this report is to provide information on a single storm event useful for storm surge modeling. The amount of observed data available from past hurricanes varies greatly and almost all of it requires further analysis and interpretation before it can be of use to storm surge modelers. An effort has been made to gather all the pertinent meteorological information into one report. The amount of data available for any single storm also varies during different portions of the storm's life, from various geographic regions, and from different sections of the hurricane. These data are subject to numerous uncertainties in interpretation. We have attempted to bring this information together to make a comprehensive analysis, to develop an accurate storm track, and to present timely histories of central pressure and radius of maximum winds.

Our intention is to make this report a comprehensive, authoritative source of meteorological information for storm surge modeling. We have tried to provide the quantitative information with as little ambiguity as possible. We have provided the basic data upon which our analysis is based so that the user may judge the degree of uncertainty in our analysis.

This report is the second of a series of reports on pertinent meteorological parameters useful for storm surge modeling. The previous report on Hurricane Carla was published as NOAA Technical. Report NWS 32 in August 1982, (Ho and Miller 1982).

## 2. PREVIOUS REPORTS

The National furricane Center (NHC) of the National weather Service (NWS), National Oceanic and Atmospheric Administration (NOAA), in Miami provided a description of significant features of all Atlantic tropical storms that occurred during 1980, including Hurricane Allen. This information was published in the Monthly Weather Review (Lawrence and Pelissier 1981) and in the National Summary of Climatic Data (NHC 1980). Tmportant features mentioned in regard to Allen were the minimum central pressure of record, the rapid deepening, and the fluctuations in intensity during its life cycle. The appearance of a double eye configuration, inner and outer, was noted from a Brownsville radar picture taken when Allen was 100 nmi oft the coast.

Willoughby and Shoreibah (1982) described secondary wind maxima associated with concentric eye walls and the evolution of the hurricane vortex in Allen and a few other hurricanes. They described the sequence of events as reported near Allen's inner core by reconnaissance aircraft on August 5 and 8, 1980. Based on data collected in Allen and other hurricanes, they concluded that an outer maximum is frequently observed to constrict about a pre-existing eye and replace it. They suggest that the concentric eye phenomen is most frequently observed in intense, highly symmetric systems.

The NHC publication on annual data and verification tabulation for the 1980 Atlantic tropical cyclones (Taylor and staff 1981) also includes a list of Allen's center fix positions obtained by aerial reconnaissance penetrations, satellite images, and land-based radar. The hurricane's central pressure, maximum winds, and other data observed by reconnaissance aircraft are also included in that report.

A smoothed "best" track for Allen has been given in publications previously cited, the Monthly Weather Review and the Climatic Data, National Summary. Cry et al. (1965) combined data from all available sources into a comprehensive report showing the most accurate and consistent locations for all tropical cyclones during their life cycle for the period 1871-1963. Neumann et al. (1981) have extended the period covered and revised earlier tracks where additional data have indicated they were necessary. The objective for these studies was to provide a firm climatological base, treating the tropical cyclone solely on the synoptic scale. Positions were given along the smoothed tracks at daily intervals for the earlier years and at 12-hr intervals-subsequent to 1930.

## 3. SCOPE OF REPORT

Values of meteorological data pertinent for storm surge models are presented in tabular and graphical form in this report. The time period covered in detail starts at. 0000 CST on September 9, 1980, and ends at 1200 CST on September 10, 1980. Since we are concerned with storm surge and not with a comprehensive look at the 3 -dimensional structure of tropical storms, the data presented are limited to surface observations. Reconnaissance aircraft data and other upper air data are used to determine surface parameters such as track, central pressure, size, winds, etc. A brief history of the storm is provided from its development stage as a tropical depression some 200 nmi east of Barbados until it finally dissipated in northern Mexico on September ll. Detailed analyses were made for the period most important for storm surge generation along the Gulf coast of the United States. For this period, data were analyzed to provide a time history of central pressure, radius of maximum winds, and forward speed. This information is tabulated and presented in table 1 at 3 -hr intervals for September 8 and part of September 9 and 10 and at 1 -hr intervals for the more crucial time of September 9 and 10.

## 4. SOURCES OF DATA

The reports discussed in section 2 were used to the maximun extent possible in this investigation. To insure accuracy and completeness of this report and to enable us to provide more detailed information on track position, speed, central pressure, etc., original records were carefully examined. This permitted us to provide the most comprehensive and detailed analysis for this storm of meteorological factors important for storm surge modeling.

Table l.-Location of storm center, central pressure, and storm size at the surface, hurricane Allen, August 8-10, 1980

| Time | Lat. | Long. | Central <br> pressure | (Nautical miles) (Statute miles) <br> $(\mathrm{CST})$ |
| :---: | :---: | :---: | :---: | :---: |
| $\left({ }^{\circ}\right)\left({ }^{( }\right)$ | $\left({ }^{\circ}\right)$ | $(1)$ | (mb) | Primary Secondary Primary Secondary |

August 8, 1980

| 0000 | 22 | 48 | 89 | 12 | 946 | 15 | 65 | 17 | 75 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 0600 | 23 | 24 | 90 | 30 | 960 | 15 | 65 | 17 | 75 |
| 1200 | 23 | 57 | 91 | 48 | 946 | 15 | 65 | 17 | 75 |
| 1800 | 24 | 28 | 93 | 00 | 912 | 10 | 65 | 12 | 75 |

August 9, 1980

| 0000 | 25 | 00 | 94 | 15 | 909 | 10 | 64 | 12 | 74 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 0600 | 25 | 15 | 95 | 21 | 917 | 10 | 61 | 12 | 70 |
| 1200 | 25 | 22 | 96 | 08 | 922 | 10 | 55 | 12 | 63 |
| 1300 | 25 | 28 | 96 | 15 | 925 | 10 | 53 | 12 | 61 |
| 1400 | 25 | 34 | 96 | 17 | 927 | 10 | 52 | 12 | 60 |
| 1500 | 25 | 38 | 96 | 22 | 930 | 10 | 51 | 12 | 59 |
| 1600 | 25 | 41 | 96 | 28 | 931 | 10 | 49 | 12 | 56 |
| 1700 | 25 | 43 | 96 | 34 | 932 | 49 | 10 | 56 | 12 |
| 1800 | 25 | 46 | 96 | 39 | 934 | 48 | 10 | 55 | 12 |
| 1900 | 25 | 49 | 96 | 44 | 937 | 47 | 10 | 53 | 12 |
| 2000 | 25 | 52 | 96 | 48 | 940 | 46 | 10 | 52 | 12 |
| 2100 | 25 | 56 | 96 | 54 | 942 | 44 | 10 | 51 | 12 |
| 2200 | 26 | 01 | 96 | 59 | 944 | 43 | 10 | 49 | 12 |
| 2300 | 26 | 06 | 97 | 06 | 945 | 41 | 10 | 47 | 12 |

August 10,1980

| 0000 | 26 | 12 | 97 | 11 | 946 | 40 | 10 | 45 | 12 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 0100 | 26 | 19 | 97 | 14 | 947 | 39 | 10 | 44 | 12 |
| 0200 | 26 | 24 | 97 | 18 | 948 | 38 | - | 43 | - |
| 0300 | 26 | 29 | 97 | 22 | 948 | 37 | - | 43 | - |
| 0400 | 26 | 36 | 97 | 31 | 949 | 37 | - | 42 | - |
| 0500 | 26 | 43 | 97 | 40 | 949 | 36 | - | 41 | - |
| 0600 | 26 | 50 | 97 | 49 | 950 | 36 | - | 41 | - |
| 1200 | 27 | 13 | 99 | 00 | 967 | 36 | - | 41 | - |
| 1300 | 27 | 42 | 99 | 48 | 990 | 36 | - | 41 | - |

The basic information is obtained from the regular reporting network of weather stations operated by National Weather Service (NWS), NOAA. These reports are part of the nation's historic weather records and are maintained at the National Climatic Data Center (NCDC), National Environmental Satellite Data and Information Service (NESDIS) NOAA. Additional data routinely stored in various forms are ship weather observations, radar weather observations, radiosonde records, and weather reconnaissance flight data. Ship weather observations are available on magnetic tape and radarscope photographs are on microfilm.

In addition, the meteorological data collected by research aircraft of NOAA's Research Facilities Center (RFC) were processed as computer printouts of flight data, flight-level wind information, and other meteorological information. These listings are stored on microfilm and magnetic tapes at the Hurriane Research Division of NOAA's Atlantic Oceanographic and Meteorological Laboratory (AMOL) in Miami, Florida. This information was made available to us for this report. A detailed description of the collection of meteorological information by aircraft, including the instrumentation, its calibration, and reliabilities, was first published in the National Hurricane Research Project (NHRP) Report No. 52 (Hawkins et al. 1962). A recent evaluation of in-flight calibration of the NOAA/RFC research aircraft instruments during 1977-78 was published in a NOAA Technical Memorandum ERL RFC-6 (Merceret et al. 1980). These publications provide the most recent discussion of the calibration and instrumentation of the NOAA research aircraft.

In addition to the network of regular reporting stations, observations are taken by many private individuals and corporations for their own use. In some cases, this material is filed with NCDC as part of NOAA's Cooperative Reporting Network.

## 5. GENERAL METEOROLOGICAL SITUATION

Hurricane Allen originated near Cape Verde Islands, off the west coast of Africa, and developed into the second most severe Atlantic hurricane in modern records. It reached tropical storm strength in the early hours of August 2 1980, and attained hurricane strength in the evening. Its central pressure dropped to 951 mb by the evening of the 3 rd as the eye passed just north of Barbados and south of St. Lucia (figure 1). The hurricane continued westward into the Caribbean at about 20 kn and passed south of Puerto Rico during the evening of the 4 th. Its central pressure deepened and reached 911 mb , the lowest pressure ever recorded in the eastern Caribbean, on the early morning of the 5th.

The hurricane weakened as it passed the southwest tip of Haiti late on August 5 and moved between Jamaica and Cuba on the 6th. This was the first of three strengthening-weakening cycles that are unprecedented in hurricane records. Allen continued on a west-northwesterly course, passed almost directly over the Cayman Islands, and reintensified rapidly as the circulation moved over the warm waters of the northwestern Caribbean Sea. Arriving at the Yucatan Channel on the 7 th , its central pressure deepened to 899 mb , the lowest pressure ever observed in the western Caribbean and the second lowest ever recorded in the Atlantic. The central pressure was on 1 y 7 mb higher than the 892 mb recorded in the Labor Day, 1935 storm that struck the Florida Keys.

The hurricane weakened for the second time when it moved over shallow waters off the north coast, of Yucatan peninsula. Its central pressure rose very rapidly, reaching 961 mb on the morning of the 8 th . As the hurricane continued west-northwestward across the warm open water of the Gulf of Mexico, Allen deepened once again with a minimum pressure of 909 mb observed during the night of the 8 th. Figure 2 shows a satellite photograph of Allen during the night of August 8, when it was at its minimum central pressure over the Gulf of Mexico.

As the hurricane approached the Texas coast on the 9 th , its intensity weakened and the forward speed decreased. Allen held on to its west-northwesterly course until mid-day and turned northward and then northwestward. After crossing the


Figure 1.-Hurricane track, August 2-11, 1980, for Hurricane Allen.


Figure 2.--Sate11ite photograph for August 8, 1980.
southern end of Padre Island just northeast of Brownsville, Texas, Allen continued on a northwesterly direction. By early morning on the loth, Allen moved inland at a slightly faster speed and turned gradually towards the westnorthwest. In the early afternoon, the hurricane passed just south of Laredo, Texas and moved into Mexico. On August 11 , Allen was downgraded to a tropical storm and finally dissipated over the mountainous terrain of northern Mexico.

In addition to the damage from the hurricane winds and storm surge, Hurricane Allen also spawned at least a dozen tornadoes over Texas. Rainfall from Hurricane Allen did not approach the extremes of some other recent storms, such as Agnes in 1972. Still, rainfall amounts up to 20 in. fell across a wide swath over the southern parts of the state.

## 6. DETALLED METEOROLOGICAL ANALYSIS

A primary focus of this report is to analyze objectively, and in detail, those meteorological factors of hurricanes used in storm surge models. For this purpose, we decided to begin with the raw observational data in order to obtain an unbiased review of all available information. This section describes these analyses. The intent of these analyses is to yield specific values of the hurricane's central pressure, the radius of maximum winds, the direction and speed of its forward motion, and the location of its center at various time intervals. Particular attention was focused on the period just before and after the hurricane made landfall since this is the time interval most critial for storm surge computation. The basic observational data used in these analyses are given in the appendix.

### 6.1 Storm Track

Generally, the analyses of meteorological data are weighted toward synopticscale motion. The hurricane track, thus obtained, is the best estimate of the large-scale storm motion and not a precise location of the eye at discrete time intervals. Therefore, such an analysis of the large-scale motion does not precisely describe the track needed for storm modeling. Track differences of a few miles, insignificant in determining the large-scale motion, can be significant for replicating high water on the open coast and inside bays and estuaries. A surge model requires, among other factors, specific information on the precise landfall point, the time of landfall, and accurate positions at closely spaced intervals in time along the hurricane track for $24-\mathrm{hr}$ prior to and after the hurricane's landfall or along the track while the hurricane is bypassing the coast.

The analysis of this report emphasizes the meteorology in greater detail during August $9-10$, when Hurricane Allen was approaching and crossing the Texas coast. The final track determined for Hurricane Allen from 0000 CST August 8 through 0000 CST. August 11 is shown in figure 3, with locations of the meteorological stations used in this report. The stations are either National Weather Service offices or military installations, except for Raymondsville and Port Mansfield, and regularly report to the National Weather Service. The positions of the center of the hurricane are shown at 6-hr intervals from 0000 CST August 8 to 11 . The central pressure (mb) and the radius of maximum winds (st. mi) are plotted to the left of $12-h r$ positions.


Figure 3.--Hurricane track showing positions every 6 hr from August 8 to 11 , with central pressure (mb) and radius of maximum winds (st.mi.) plotted at $12-\mathrm{hr}$ intervals.

Since the interest in this report is a detailed determination of the path of the hurricane immediately before and after landfall, the area nearest the coast, enclosed in the box on figure 3 , is enlarged for greater clarity in figure 4. The hurricane locations at hourly intervals are indicated by open circles. The time at every other hour, together with values for the central pressure (mb) and radius of maximum winds (st. mi), are shown along the track. In addition to regularly reporting weather stations, observations by satellites, personnel of private industries, private individuals, and eyewitnesses were useful aids in determining the storm track.

Any final determination of the track and speed of forward motion of a hurricane, especially over data sparse regions, has inherent uncertainties. The selected track is finalized from a subjective analysis to account for all available information. Figure 5 is an example of the information used in our analysis. Hurricane eye positions based on radar weather observations reported from Brownsville and Corpus Christi, Texas are shown as solid dots. Aircraft reconnaissance penetration fixes are shown by triangles. Locations of the hurricane's center determined from satellite observations are given by diamonds. The selected track, fitted by eye and guided by bias in center location, is presented by the solid curve. Locations at 6-hr intervals, with the time indicated, are shown by open circles.

The data from radar fixes and aircraft penetrations are the primary resource used in determining the track and speed of forward motion of the hurricane over the open ocean. However, information obtained from satellite observation and from all ships operating in the area was considered in determining the final track and speed of motion. The information from each of these sources was cacefully evaluated before a final track was selected.

The track position selected (figure 5) deviates from the average radar position line along almost the full length of the track. We chose to follow a more northerly track that relies heavily on aircraft reconnaissance fixes. Radar locations are based upon echo returns from the wall cloud. These have shown some bias in location relative to the pressure and wind center, our primary interest. Pressure, wind, and dropsonde observations obtained by aircraft penetrations are usually more accurate than radar eye fixes in determining the pressure/wind center. Center positions determined by aircraft must, of course, be evaluated in terms of possible navigational instrument error. Holliday. (1966) compared reconnaissance aircraft center fixes with the best-fit radar track of the eyes of seven hurricanes obtained by coastal radars. Results of the comparison indicate that the maximum differences vary from 13 to 37 nmi in individual hurricanes.

Aircraft reconnaissance fixes near 0000 CST on the loth (figure 5) were based on airborne radar observations. These positions seem to be biased in a similar manner as those obtained from land-based radar observations. We decided to ignore these two reconnaissance fixes at the coast and adopted a track crossing the barrier island in a north-northwesterly direction instead of a westerly direction. This is supported by wind observations at Port Mansfield, Texas, and an eyewitness report of calm winds as the eye of the hurricane passed over the station (Hagan 1982).


Figure 4.--Hurricane track showing positions at l-hr intervals from 1200 CST on August 9 through 0900 CST on August 10, 1980, with central pressure (mb) and radius of maximum winds (st. mi.)plotted at $2-\mathrm{hr}$ intervals.


Figure 5.--Hurricane eye center obtained from radar weather observations (.), aircraft reconnaissance penetration fixes ( $\Delta$ ), and satellite observations ( $\circlearrowleft$ ) together with positions of hurricane center ( $\Theta$ ) on the selected track.

### 6.2 Forward Speed

The translation speed of the hurricane is another important ingredient in determination of the surge along the open coast and in bays and estuaries. Hourly positions were the basic building blocks for determining this forward speed of translation. First, speeds between successive hours from initial positions along the best track were determined and plotted on a time scale, and a smooth curve was drawn subjectively to minimize abrupt changes. Second, speeds from three successive hours were averaged and plotted at the mid-hour, and smooth curves drawn from these data were used to adjust the hourly locations. The new locations were examined with regard to the observed data and, if necessary, some further adjustments were made. This process was continued in an iterative fashion until the best combination between smooth forward speeds and observed eye postions was obtained. This process helped to obtain the best possible estimate of forward speed and hourly locations.

Figure 6 shows the forward speed of Allen which was unusually fast after its rapid development east of the Lesser Antilles and westward movement into the Caribbean. Its average speed stayed at about 20 kn until its center reached the north-western Caribbean Sea. The hurricane then decelerated as it moved into the Gulf of Mexico. Its forward motion slowed to a speed of about 15 kn on August 8 and to about 10 kn on the early morning hours of the 9 th . There was a further distinct slowing of forward progress in the afternoon of the 9 th when Allen was about 50 nmi from the Texas coast. Its center moved at an average speed of about 6 kn for a period of $12 \mathrm{hr}(0900-2100 \mathrm{CST})$ on the 9 th and then crossed the barrier inland (092100-100300 CST) at an average speed of 7 kn . After moving inland on the $10 t h$, Allen continued slowly northwestward at about 10 kn into the mountains of northern Mexico.

### 6.3 Central Pressure

The most important factor in storm surge models is the intensity of the hurricane which is directly related to its central pressure. Figure 7 shows the finalized track of Hurricane Allen as the storm crossed the Texas coast. Also shown are minimum pressures observed at regular reporting stations and minimum pressures obtained during reconnaissance aircraft penetrations. These observations were not all obtained at the sane time. Since the storm track did not cross any land station location, none of the values reported at land stations are equal to the minimum central pressure in the storm.

Figure 8 shows our analysis of the pressure information from land stations and aircraft reconnaissance flights that was used to obtain a time history of Allen's minimum pressure. The curve drawn is, in general, a curve fitted to the data by eye. Allen deepened for the third time in its lifespan on the evening of August 8th. A minimum pressure of 909 mb , observed by aircraft reconnaissance at 2358 CST on the 8 th, was the lowest reading ever recorded in the western Gulf of Mexico. We considered this pressuce to be the lowest that occurred in Hurricane Allen as it approached the coast. The short time interval between central pressures obtained by aircraft, combined with other information, did not indicate any lower pressure at intermediate times. As Allen continued its course westnorthwestward, approaching the Texas coast, its intensity weakened. While the hurricane's central pressure rose steadily, the characteristics of its inner core region (as indicated by its eye wall structure and maximum wind distribution) appeared to have undergone dramatic changes. We shallfirst look into pressure

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changes at individual stations; discussions of other phenomena are presented in subsequent paragraphs.

Sea-level pressure recorded at coastal stations was used to examine the pressure variation during the period when Allen approached the coast and moved inland. Hourly observations of sea-level pressure recorded at Brownsville and Kingsville, Texas were plotted against time and the distances of the hurricane's center from each station at various times as determined from Allen's track (figure 3). By further examining the rate of pressure change at each station and in Allen, we assess the extent to which the pressure variations at individual stations can be related to the movenent of the large-scale pressure distribution of the hurricane.

Figure 9 shows a plot of hourly observations of sea-level pressure against time recorded at $B$ rownsville, Texas (dots) with solid lines joining the data points. The distances of the hurricane's center from the station at the time of observation are shown in circled dots and joined by the dashed lines. These curves indicate that the pressure variation at Brownsville appears to be closely related to the distance from the hurricane's center, especially when Allen was within 85 mmi of the station. The rate of pressure drop averaged about 1.7 mb per hour when Allen was within the $85-n \mathrm{mi}$ range. The rate of pressure drop was compratively small when the hurricane was farther away from the station.

Figure 10 shows similar data from Kingsville, Texas. The curves for Kingsville reveal a rapid pressure fall on the 9 th when the center of the hurricane was about 100 nmi from the station. For the 8 -hr period starting from 2200 CST, the


Figure 7.—Minimum pressure recorded at land stations and by aircraft reconnaîssance during Hurricane Allen for period 1200-2300 CST on August 9, 1980.


## Figure 8.-Variation of central pressure with time, Hurricane Allen, August 8-10, 1980.

average drop in pressure at the station amounts to approximately 1.65 mb per hour which is almost the same as that observed at Brownsville (figure 9). It is of interest to note, as a comparison, the rate of falling central pressure in Allen during one of the rapid deepening stages in its life cycle. Allen's central pressuce dropped at an average rate of 4.0 mb per hour over a 12 -hr period from 0600 to 1800 CST on August 8 when Allen was located in the western Gulf of Mexico (Lawrence and Pelissier 1981). This change in intensity is another factor (among others) that influences pressure changes at individual stations. The central pressure in Hurricane Allen rose steadily prior to the time of landfall (figure 8). It was 909 mb at midnight on the $8 \mathrm{th}, 922 \mathrm{mb}$ at noon on the 9 th , and reached 946 mb in the next 12 hr when Allen crossed the coast. The weakening of the hurricane, as indicated by increasing central pressure with time, moderated the rate of falling pressure at land stations during the period of Allen's approach.


Figure 9.--Hourly observations of sea-leve1 pressure recorded at Brownsville, Texas (.) and distance of Allen's center from station ( 6 ) for period 1200 CST on August 8 through 2400 CST on August 10.


Figure 10.-Same as figure 9, for Kingsville, Texas.

### 6.4 Wind Analysis

We analyzed the wind field for Allen in two ways. We first examined the wind observations of the regular reporting land stations. Next, we did a streamline analysis of the windfields at the 6 -hr intervals for the period from 0600 CST August 8 through $0000 \operatorname{CST}$ August 10 . This wind analysis was used to aid in the determination of the radius of maximum wind. It also provided some guidance in determining the best track.

### 6.4.1 Analysis of Observed Winds at Weather Stations

Supplemental to the minimum pressure reported at stations during hurricane passage, surface winds were reported at several weather stations operated by the National Weather Service, the military services, and other federal agencies such as Coast Guard stations and FAA operated airport facilities. We attempted to relate the variation of surface wind speed to the distance of the storm's center from the station in the same manner described for pressure variation at individual stations. Figure 11 shows a plot of the time variation of hourly wind speed recorded at Brownsville, Texas (solid curve) and the distance of the storm's center from the station (circled dots plotted at hourly intervals). The resultant magnitudes, after the storn's speed of translation was subtracted from the observed wind speed, are also shown in the diagram (dashed line). These give


Figure 11.-Hourly observations of wind speed and distance of Allen's center from Brownsville, Texas for period 1200 CST on August 8 through 2400 CST on August 10, 1980.
the wind speeds relative to the storm's center. The observed wind directions at hourly intervals are shown by dots at the top of the diagram. The solid curve reveals that the maximum wind at Brownsville was observed at 1645 CST on the 9 th when Allen was about 49 nmi to the east. A second maximum appeared at 0500 CST when the hurricane's center, moving away from the station, was located at a distance of 50 nmi . This second maximum relative to the storm's center actually occurred 2 hr earlier, and the distance from the storm's center was 35 nmi when the speed of its forward motion was subtracted from the observed wind speed (dashed line). In general, there is an inverse corcelation of the wind speed and the distance from the storn's center except for winds inside the radius of maximum winds. The dashed line indicates that a small fluctuation of the wind speed occurred inside the region of wind maxima.

A similar plot of hourly winds for Kingsville, Texas (diagram not shown) shows that the time variation of wind speed at Kingsville also was closely correlated with distance from the storm's center. A maximum observed wind of 60 kn occurred when the storm's center was 35 nmi south of the station.

### 6.4.2 Streamline Analysis of Surface Charts

Since surface data were too limited and scattered to make an analysis of the winds when the hurricane was located some distance off the coast, all reconnaissance aircraft observations within intervals of several hours were combined and plotted on surface charts. In the course of penetrating the center, aerial reconnaissance recorded flight-level winds within a $100-n m i$ radius of the hurricane's center. No ship report was nearer than 70 nmi to the center. Surface charts at $6-\mathrm{hr}$ intervals were analyzed for the period from 0600 CST ( 1800 GMT) August 8 through 0000 CST ( 0600 GMT ) August 10 (figures 12a through 12 g ). Observations at coastal stations and ship observations taken at map time were plotted on the appropriate charts. Flight-level winds and observed minimum pressure reported by reconnaissance aircraft within 6 -hr of map time were also plotted.

As a supplemental aid in the streamline analysis, the position of each observation taken in aerial reconnaissance was measured in terms of azimuth angle and radial distance relative to the hurricane's center at the time of observation. Each wind observation was then transposed to the location relative to the hurricane's center. These transposed observations are not shown in the charts. For the purpose of illustration, examples of two transposed wind observations were plotted on figure l2a. Flight-level wind of $170^{\circ} / 80 \mathrm{kn}$ was observed at 081840 GMT when the reconnaissance aircraft was located at $23.7^{\circ} \mathrm{N}, 90.5^{\circ} \mathrm{W}$. This information was plotted on the chart for 081800 GMT. We then obtained the location of the hurricane's center at the time of observation ( 081840 GMT) by interpolation of hourly positions given by the hurricane tracking charts (figures 3 and 4). The next step was to measure the location of the plotted observation relative to the hurricane's center at the time of observation, yielding an azimuth angle of $103^{\circ}$ and a radial distance of 72 nmi . Using this relative location, the observation was transposed to a location relative to the hurricane's center and plotted on the chart with the wind direction and speed underlined. The transposed location on the chart is just slightly east of the location shown for the observation. Similarly, the observed wind of $070^{\circ} / 85 \mathrm{kn}$ was plotted at the location with azimuth angle of $315^{\circ}$ and a radial distance of 61 nmi , relative to the hurricane's center.

Figures 12a through 12 g show the stream analysis of winds within the hurricane's circulation at 6 -hr intervals from 1200 CST ( 1800 GMT) on August 8 through 0000 CST ( 0600 GMT ) on August 10 . Isotach patterns are shown in dashed lines. The maximum flight-level winds reported near the eye are not shown on the charts to make room for a clear illustration of the isotach pattern near the center. These maximum flight-level winds are shown in figure 13 and listed in table A. 3 of the appendix. From 1200 CST August 8 through 0000 CST August 9, the hurricane was in a rapidly deepening stage. The central pressure dropped from 946 mb to 909 mb (figure 8). A maximum flight-level wind of 145 kn was reported at 1719 CST on August 8 at a radial distance of 10 nmi . During the same period, an area of secondary wind maximum appeared on each of the three charts (figures 12a, b, and c) at a radial distance of $60-65 \mathrm{nmi}$. Winds of $80-85 \mathrm{kn}$ were reported in this area throughout the $12-h r$ period and remained the same for


Figure 12a.--Streamline analysis, 1200 CST (1800 GMT), August 8, 1980.


Figure.12b.--Streamline analỳsis, 1800 CST, August 8, 1980.


Figure 12c.-Streamline analysis, 0000 CST, August 9, 1980.


Figure 12d.--Streamline analysis, 6000 CST, August 9, 1980.


Figure 12e.-Streamline analysis, 1200 CST, August 9, 1980.


Figure 12f.--Streamline analysis, 1800 CST, August 9, 1980.


Figure 12g.-Streamline analysis, 0000 CST, August 10, 1980.
the next 12 hrs through 1200 CST August 9. The radial distance of the secondary maximum from the center reduced only slightly during the latter half of the $24-\mathrm{hr}$ period or from 0000 CST to 1200 CST August 9. At 1200 CST August 9, the hurricane was centered about 60 nmi of $f$ the Texas coast and wind speeds of about 100 kn were reported at radial distance of 10 nmi , as well as $55-60 \mathrm{nmi}$ from the center. By 1800 CS ' August 9, the hurricane's center was located at about 23 nmi off the coast, and Brownsville, Texas reported winds of 38 kn from the north after experiencing maximum winds of about 45 kn during the past hour (figure 11). High winds in Allen were then recorded off the Texas coast at a radial distance of about 40 nmi from the center, but a lesser wind speed maximum was still located near the eye at a radial distance of about $10-15 \mathrm{nmi}$. As the hurricane continued to approach the coast, the observed flight-level winds near the eye decreased to $60-70 \mathrm{kn}$ while the extreme winds at the outer band increased in magnitude and the area migrated inward.

### 6.5 Radius of Maximum Winds

The size of a hurricane is commonly denoted by the distance between the lowest central pressure and the band of highest winds around the center. The radius to the maximum winds was determined from all the observations available for this storm. Three different types of observations were available. The first measure includes the maximum flight-level winds and estimated surface winds as reported by reconnaissance aircraft. The second is the radar eye diameter, also reported by reconnaissance aircraft as well as by surface observing stations. Some optical reports were used when the reconnaissance aircraft was in the eye of the storm. The third measure, useful only after the hurricane was near shore, estimates the radius from surface wind records at land stations.

Flight-level winds, recorded at one-second (1-s) intervals by the reconnaissance aircraft of the NOAA Research Flight Center were processed, and 10 -s running averages of the $1-s$ intervals are available on magnetic tape. The aircraft location for each observation was translated as a relative position to the storm center. From these records, composite maps of flight-level winds at given intervals were plotted by computer and made available to us by the Hurricane Research Division of NOAA/AMOL. Analysis of these maps yielded another measure of the radius to maximum winds.

Figure 13 is an example of a composite map of flight-level winds for the period of 1330 to 1545 CST on August 9, 1980. The wind data (in $\mathrm{m} / \mathrm{s}$ ) recorded at an altitude of 2368 m were plotted at translated positions relative to the storm center. The highest wind speed of approximately 100 kn along each leg of penetration of the eye was located about 10 nmi from the center. A secondary maximum of about 90 kn can be identified at about 65 mmi from the center. Similar distributions of flight-level winds can be identified in composite maps of other time periods (diagrams not shown). The map series indicates that the secondary (outer) maximum migrated inward as the inner wind maximum weakened. The evolution of this phenomenon can be illustrated by radial wind profiles constructed from flight-level wind data recorded on August 9 and 10 . The selected periods covered the time when the hurricane was located about 75 nmi east-southeast of Brownsville, Texas until its center was some 40 nmi inland ( 6 hr after crossing the coast over the southern tip of Padre Island).

Figure 14 shows flight-leve 1 winds recorded at the $700-\mathrm{mb}$ level (minimum height of 2453 m to 2510 m ) between 1200 CST and 1500 CST on August 9. The data points


Figure 13.-An example of composite map of flight-level ( 2500 m ) winds for period 1745 to 1900 cST, August 9, 1980. Numerials indicate time of observation in GMT (hours, minutes and seconds). Wind barbs denote speed in $\mathrm{m} / \mathrm{s}$.


Figure 14.--Radial profile of flight-level ( 2500 m ) winds recorded during period $1200-1500$ CST, August 9, 1980. Wind speeds (in knots) are resultant wind speeds with speed of storms's motion subtracted from observed winds.
were plotted against the radial distance from-the center of the hurricane at the time of observation. A smooth envelopment curve drawn from the data points reveals that maximum winds of about 90 kn were observed at a distance of 10 nmi from the center with a secondary maxima at about 65 nmi . Similar radial wind profiles at three different levels for the previous day were constructed by Willoughby et al. (1981). However, the wind maxima on the 8 th (about 110 kn ) were higher than that of the $9 t h$, while the magnitude of the secondary maxima remained about the same.

Figure 15 shows flight-level winds recorded during the period 1500 to 1800 CST on the 9th. Again, the data were plotted against the radial distance from the center of the hurricane. The smooth envelopment curve drawn from these data points indicates that the wind maximum was then about 45 nmi from the center. The inner wind maximum had become the secondary maximum and was still located at


Figure 15.--Radial profile of flight-level ( 2500 m ) winds recorded during period 1500-1800 CST, August 9, 1980. Wind speed (in knots) are resultant winds with speed of storm's motion subtracted from observed winds.
about 10 nmi from the center. A small fluctuation of wind speed appeared 20 to 25 nmi from the center. A similar plot of flight-level winds recorded during the period of 0500 CST through 0800 CST on August 10 is shown on figure 16 . There were no observations taken within 45 nmi of the hurricane's center during that time period because the reconnaissance aircraft was flying over water along the coast. We assume that the wind maximum at the flight level remained at a distance of about 45 nmi from the eye.

The radial profiles of flight-level winds described in preceding paragraphs clearly indicate that the wind maximum near the eye reduced in magnitude, while the outer maximum migrated inward and became the dominant feature in the radial


Figure 16.--Radial profile of flight-level ( 2500 m ) winds recorded during period $0500-0800 \mathrm{CST}$, August 10, 1980. Wind speed (in knots) are resultant wind speeds with speed of storm's motion subtracted from observed winds.
wind profile. The magnitude of the wind maxima at a radial distance of 45 nmi appeared to remain the same at the flight level, as the hurricane moved over land on August 10 .

The second measure used in determining the radius of maximum winds is the radar eye diameter. This type of data is obtained from reconnaissance aircraft reports and from land based radar weather observing stations. Figure 17 is a reproduction of a Brownsville, Texas radarscope photograph taken at 0430 CST on August 9, 1980, showing Allen's well-defined concentric eye structure. Similar to other mature hurricanes, the eye of the storm is defined by a ringlike radar echo which is separated from the spiral bands of the storm. Inside this ring, the eye is clear of precipitation echoes. In the case of Allen, this eye structure shown in the radar photograph will be referred to as the "inner eye." There was also a relatively wider ring in the storm interior with little or no radar return. This echo-free area was surrounded by another ring of wall clouds. Thus, the structure of Allen's inner core appeared to have two concentric wall clouds.

The phenomenon of concentric eye structure was first described by Fortner (1958) and observed in Hurricane Donna of 1960 (Jordan and Schatze, 1961) and in Hurricane Carla of 1961 (Jordan 1966). Hoose and Colon (1970) documented a complete concentric eye cycle in Hurricane Beulah of 1967 and related the inner eye deterioration with the shift of maximum wind from the inner eye to the outer eyewall. Similar behavior was described by various authors in Typhoon Gloria of 1974 (Holliday 1977) and in Hurricanes Debbie of 1968 (Gentry 1970;


Figure 17.-Brownsville, Texas radarscope photograph taken at 0430 CST August 9, 1980, showing Allen's well-defined concentric eye structure.

Hawkins 1971; Black et al. 1972), Anita of 1977, David of 1979, and Allen of 1980 (Willoughby et al. 1981).

Our main concern in this study is the hurricane characteristics which closely describe surface wind distributions important for storm surge modeling. We looked in great detail at the variation of these parameters during the prescribed period prior to, and after, the hurricane crossed the coast. Land based radar weather observing stations report, among other data information, the diameter of the inner eye as a hurricane approaches the coast. In the case of Allen there was no significant variation with time in the diameter of the inner eye reported during the period of interest on the 9 th. We further examined photographs of radar echoes taken at Brownsville, Texas during the period of 0300 CST on the 9 th through 0800 CST on the 10 th. From these photographs we obtained measurements of radius from the storm's center to the outer perimeter of the inner eye wall and the radius to the rim of the outer echo free area.

Figure 18 shows smooth curves joining these measurements to indicate the variation of the eye structure with time. The data points were read off radar-scope photographs, beginning at 0330 CST and ending at 2300 CST , on the 9th. There was no reading obtained for the lower curve after 2000 CST when the inner eye was completely filled. Radar pictures indicated that the filling process began at about 1430 CST when openings of the inner eye wall appeared to the southwest. These openings, which reflected the dissipation of convective clouds in that quadrant, occurred during the period when flight-level winds recorded by reconnaissance aircraft (1500-1800 CST) were decreasing in magnitude near the inner eye wall (figure 15). The lower curve in figure 18, showing the variation of the radius to the outer perimeter of the inner eye wall, reveals a rapid increase in radius during the $2-h r$ period of 1230 to 1430 CST on the 9 th. This increase in radius indicated an outward expansion of convective precipitation in the inner eye wall prior to the filling of the inner eye. The upper curve in figure 18 shows the variation of radius from Allen's center location to the rim of the outer echo-free area with time. This curve shows a general trend of decreasing radius in the first 12 hr followed by short period oscillations and a rapid decrease in magnitude at around 2100 CST on the 9 th. The general trend of decreasing radius occurred in the same time period that the secondary wind maximum migrated inward from 65 nmi of the center. The rapid decrease in radius occurred when Hurricane Allen was about 10 nmi from the coast. Although no specified relation between the radius of the outer echo-free area and the radius to wind maxima is considered in this report, we speculate that they will tend to either increase or decrease together.

The third measure used in determining the radius of maximum winds came from surface winds recorded at land stations. These can be illustrated by radial wind profiles constructed from surface wind records. Figure 19 shows radial wind profiles for the time periods when Allen's center was approaching the coast and when it was moving over land. The wind data plotted on the diagram were resultant wind speeds after the hurricane's speed of translation was subtracted from the observed wind speed. The upper diagram shows a smooth curve fitted by eye to the data observed at Brownsville during the period 0600 through 2100 CST on the 9 th and winds recorded at other stations at 01800 CST. This curve indicates that the maximum wind at Brownsville occurred when Allen's center was at a distance of 49 nmi from the station. Since the center bypassed the station at a distance of 21 nmi , there were no observations of surface winds in the hurricane's eye region.


Figure 18.-Eye radii obitained from Brownsille, Texas radarscope for period $0300-2200 \mathrm{CST}$, August 9, 1980. Upper curve shows radial distance of outer eye from Allen's center. Lower curve shows radial distance of outside parameter of inner ege wall from the center.

The lower curve (figure 19) was fitted by eye to data points based on surface winds recorded at land stations when Allen was moving over land. Observations taken at 2200 and 2300 CST on the 9 th at Brownsville when Allen's center was 8 nmi and 4 nmi off the coast, respectively, were included to show the decrease in wind speed from the maximum towards the center. The highest wind speed plotted on this curve was recorded at Kingsville Air Force Base when the center of Hurricane Allen was 35 nimi south of the station. This indicated a decrease in the radius of maximum winds from 49 nmi to 35 nmi as the hurricane approached the coast and moved inland. This observed decrease in radius of maximum winds is supported qualitatively by the eye radius which appeared to have decreased rapidly around 2100 CST on the 9 th (figure 18).

Figure 20 provides a curve from which the radius of maximum winds can be determined. It is based on analyses of all available observations previously described. The distances of observed maximum flight-level winds from the hurricane's center reported during aerial reconnaissance are shown by triangles. Radial distances of maximum winds obtatned from analyses of flight-level winds are shown by circled dots. These radial distances were read directly from composite charts of flight-level winds (e.g., figure l3), except for those shown in the first $24-\mathrm{hr}$. The results from the earlier time period


Figure 19.-Radial profiles of surface winds constructed from observations taken at Brownsville, Texas and other stations for periods 0600-2100 CST on August 9 (upper curve) and 092200-100800 CST, August 1980, (1ower curve).
were interpolated from analyzed surface charts (figures 12a through g). Radial distances determined from surface winds recorded at land stations are given by inverted triangles. The magnitude of extreme winds recorded at a given time was classified into two categories, a primary and a secondary wind maximum. The occurrences of primary wind maxima are denoted by solid lines while occurrences of secondary wind maxima are indicated by dashed lines. A shift of wind maxima from a radial distance of 10 nmi near the eye to that of about 45 nmi from the center seems to have occurred near 1500 CST on August 9.

Analysis of flight-level wind distributions (e.g., figure 13) and radial profiles of flight-level winds constructed from reconnaissance flight data (figures 14 through 16) reveal that the primary wind maximum near the inner eyewall decreased in magnitude while the secondary maximum migrated inward and became the dominant. feature in the radial profiles. The analyzed results of flight-level ( 2500 m ) winds, yielding an estimated radius to wind maximum of 10 nmi during the period of 1800 CST on the 8 th through $1200 \operatorname{CST}$ on the 9 th, can


Figure 20.-Variation of radius of primary (solid line) and secondary (dashed line) wind maxima with time, Hurricane Allen, August 8-10, 1980.
be applied to the surface. This was supported by the consistency of observed winds at three different altitudes ( 850,600 , and $500 \mathrm{mb}-1 \mathrm{le}$ vels) recorded within a same time period on the 8th. The vertical structure of wind during that time period had little or no variation with height between flight-levels. (See figure 14 of Willoughby et al. 1982.) The secondary maximum migrated inward as the winds near the inner eyewall decreased in magnitude. By 1500 CST, the existence of a wind maximum at a radial distance of 45 nmi from the hurricane's center can be identified in both flight-level wind analyses and surface wind analyses. For Allen, then, this distance became the radius to the maximum wind, which reached $100-110 \mathrm{kn}$ at flight-levels. After 1500 CST , winds of $70-90 \mathrm{kn}$ at a radial distance of $10-20 \mathrm{nmi}$ were reported by aerial reconnaissance aircraft. The extreme winds were, in fact, secondary wind maxima observed at the flight-level.

Analyses of surface winds recorded at Brownsville; Texas yielded results which agree very we 11 with flight-level winds observed between 1600 CST and 1800 CST. These results indicated that wind maxima occurred at a radial distance of about 45 nmi from the center, implying that maximum winds would begin to strike coastal areas near 1600 CST when the hurricane's center was some 50 nmi off the coast. At Port Mansfield, maximum gusts of 120 kn were recorded at 2240 CST just before the recording instrument became inoperative. This indicated a reduction of the radial distance of maximum winds to about 36 nmi as Allen approached the coast. Surface winds observed at Kingsville, also indicated that the maximum winds remained at the same radial distance of 36 nmi from the center as Allen moved over land.

### 6.6 Summary and Discussion of Meteorological Analysis

The individual parameters from our analysis of Hurricane Allen are listed in table 1. These are listed for locations of the hurricane center at $3-\mathrm{hr}$ intervals on August 8 and part of August 9 and 10 , and at 1 -hr intervals between 1200 CST on the 9 th and 0600 CST on the 10 th . For each location, central pressure and the radius of primary and secondary wind maxima (both in nautical and statute miles) are listed. The table provides, in convenient form, the information that could be obtained from analyses of the basic data described in various sections.

It is of interest to note that there were two areas of wind maxima of approximately equal magnitudes observed during a brief time interval near 1200-1500 CST on the 9th (figure 14). This phenomenon occurred during a transition period when the secondary wind maximum migrated inward and winds near the inner eyewall weakened. During this time period, Allen's central pressure increased from 922 to 930 mb . As the intensity of Allen continued to weaken with its central pressure rising gradually, the wind maximum near the inner eyewall decreased in magnitude. The extreme winds at a greater radial distance then became the dominant feature which influenced the storm surge generation.

## 7. DISCUSSION

It would be only speculation had we attempted to explain the evolution of wind maxima based on observations of a single hurricane. However, characteristics of the hurricanes previously cited reveal similar evolutions of wind maxima assocfated with observed phenomena of concentric eye walls. Similar to Allen, Hurricanes Beulah of 1967 and Anita of 1977 are good examples of such evolutions observed in the western Gulf of Mexico. Hurricane Anita deepened on August 31 and September $l$ and its central pressure dropped to a minimum of 926 mb during the night of September 1 , before striking the Mexican coast about 130 nmi south of Port Isabel, Texas. The concentric eye walls and the associated wind maxima were observed just after the central pressure fell below 940 mb (Willoughby and Shoreibah 1982). Hoose and Colon (1970) observed that concentric eye walls appeared in Hurricane Beulah shortly before the hurricane's central pressure dropped to 940 mb when Beulah was located about 200 nmi south of San Juan, Puerto Rico. They also observed that maximum winds occurred in the precipitation echo-free area outside the inner eye wall. They deduced that the concentric eye configuration would be quite unstable and the inner eye was probably dissipated by the subsiding downdrafts generated by the development of the more stable outer eye system. This concept was contirmed by numerical computations made by Shapiro and Willoughby (1982) using a dynamic model of Eliassen (1951).

We examined reconnaissance flight data recorded during a period of five hr when Beulah was about 120 nmi southeast of Brownsville, Texas (figure 21). Our analysis shows that wind maxima appeared at radial distances of 15 and 45 nmi from the hurricane's center near 1430 CST on September 19, after Beulah reached its maximum intensity and the central pressure dropped to 923 mb .

It is gratifying to note that both Beulah and Allen weakened before crossing the southern portion of the Texas coast. In both cases, the maximum winds near the inner eye wall decreased in magnitude while the outer wind maximum contracted in radius. If the hurricanes had deepened and reached their maximum intensity just before making landfall, winds of much higher magnitudes would have occurred at close proximity to the hurricane's center. Under such a configuration of surface winds, surge generation caused by the approaching hurricane would be quite different from that of a weakening hurricane. Table 2 shows the minimum central pressure of some hurricanes and typhoons near the time when concentric eye walls were observed. An examination of the time variation of central pressure in these hurricanes and typhoons reveals that they were in a deepening stage prior to the observed events. The central pressure of all hurricanes and typhoons in this table, except for Debbie of 1969 , dropped below 940 mb around the rime when concentric eye walls were observed. We speculate that this phenomenon is associated with intense hurricanes in their deepening stage after a threshold intensity (as measured by central pressure of around 940 mb ) is reached.

It may be redundant to iterate the importance of wind maxima acting on and influencing the water levels in bays and estuaries when a hurricane crosses the coast and moves over land. However, based on the observations previously discussed, we recommend that further studies are needed to scrutinize wind configurations in intense hurricanes which deepen in the close proximity of the coast. Though this short-lived phenomenon might not change the general characteristics of a mature hurricane, it might well be an important factor to consider in the simulation of surge heights by using historical hurricane parameters as input to surge modeling.

Table 2.-Central pressure of hurricanes and typhoons near the time when concentric eye walls were observed

| Storm date | Name | Central pressure (mb) |
| :--- | :--- | :---: |
| August 26, 1954 | Typhoon Ida | 892 |
| March 24, 1956 | Typhoon Sarah | 940 |
| September 6, 1960 | Donna | 940. |
| September 9, 1967 | Beulah | 940 |
| September 19, 1967 | Beulah | 923 |
| August 20, 1969 | Debbie | 954 |
| November 5, 1974 | Typhoon Gloria | 937 |
| September 2, 1977 | Anita | 930 |
| August 28, 1979 | David | 938 |
| August 30, 1979 | David | 930 |
| August 8, 1980 | Allen | 940 |



Figure 21. Composite map of flight-level ( 990 m ) winds observed in Hurricane Beulah during period 1230-1800 CST, September 19, 1967.

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## APPENDIX - METEOROLOGICAL DATA

This appendix gives the basic meteorological data used to develop the analysis presented in this report. The tables list the observations of sea level pressure and wind data at land stations and the hourly reports from ships. They also include positions of the hurricane center as determined from reconnaissance aircraft and land-based radar data.

Table A.l lists the hourly observations of sea level pressure, wind direction, wind speed, and gustiness obtained at regular reporting stations. These hourly observations are taken from U.S. Weather Bureau Surface Weather observations forms (WBAN 10) covering the period August 8 through August 10, 1980. The sea level pressure is given in units of millibars (mb). The wind direction is given as the direction from which the wind was blowing to the nearest ten degrees, measured clockwise from north. The reported wind directions were, in most instances, in compass points, i.e., N, NNE, NE, ENE, etc., and converted to degrees from north. The observed wind speed is a 5 -min average determined from recorded observations. The gustiness is characterized by sudden, intermittent increases in speed where at least 9 kn were indicated between peaks and valleys with a time interval of less than 20 s . The wind speed is determined to the nearest knot.

The National Weather Service maintains a series of radar observing stations along the U.S. coastline from Brownsville, Texas to Eastport, Maine. These radars are used to track hurricanes for use in the hurricane warning system. Two stations in this network were in a position to track Hurricane Allen as it approached the Texas coastline. These stations are at Brownsville and Corpus Christi, Texas. The radar eye positions reported by these NWS stations, when the center was within range of the land-based radar, are listed in table A. 2 by half-hourly intervals.

The complete reconnaissance alrcraft reports were considered too voluminous to reproduce entirely in this report. Table A. 3 lists those reports that provided the locations of storm center, observed sea level pressure, estimated surface winds, and/or the dianeter of the eye. For a few of these reports, the range and maximum winds from the storm center were obtainable. This information is also presented in the remarks column. The reported position of the storm center has the same accuracy as the aircraft position determined by radar and the land navigational systems. With some exceptions, the accuracies of these positions is generally within 1 nmi . The central pressure data is given in millibars and is determined by dropsonde or extrapolation from flight-level data.

To obtain weather reports from oceanic areas, the NWS solicits the cooperation of merchant ships of U.S. and foreign registry and of non-military U.S. Government ships. There are about 200 ships that participate in this progran. Observations are visual plus barometric and occasionally cyanometric pressure and are reported by radio at synoptic time when the ship is underway. In addition to the ships in the cooperative programs, all ships are asked to send special radio reports when tropical storms or hurricanes are encountered. Data from ships that report through the regular reporting system, supplemented by data from those ships submitting weather observations after the arrival at their major destination, are listed in table A.4. We have restricted our listing to the location of the reporting ship, sea level pressure, and wind data. The data are
grouped by the time of observation from 1800 CST , August 7, through 1200 CST , August 10, 1980. This set of data was useful in the analyses of the pressure field and the wind field of the hurricane, especially when its center was located off the coast. The aneroid barometers on ships in the cooperative observing programs are calibrated by the NWS when a ship is visiting a port in the U.S. where an NWS Port Meteorological Officer is assigned. These calibrations, however, may not be as frequent as desirable.

## 

alice tx
LATITUDE $27^{\circ} 4^{\prime \prime} \mathrm{N}$ LONGITUDE $98^{\circ} 4^{\circ} \mathrm{W}$
ELEVATION


DATE AUGUST 8TH. 1980
PRESSURE(MB) 1013.9 1013.6 $1013.2 \quad 1013.9 \quad 1013.9 \quad 1012.9 \quad 1011.7 \quad 1010.3 \quad 1009.6 \quad 1010.3 \quad 1010.0 \quad 1009.6$


DATE AUGUST 9TH. 1980
 $1007.6 \quad 1005.9 \quad 1006.5 \quad 1008.5 \quad 1007.6 \quad 1005.2 \quad 1004.5 \quad 1003.5 \quad 1000.4$



DATE AJGUST $30 T H, 1900$

 GUST (KM)

## ABLE TVE - CONTINUED

EA EVEL PRESSURE ANO WINO OATA FROM REGULARLY REPORTING STATIONS

AUSTIN TX

$$
\text { EATIYUQE } 300170^{\circ} N \text { LONGITUDE } 97^{\circ} 266^{\circ} \mathrm{W}
$$



DATE AJGUST BTH, 1980
PRESSURE (MB) $10.15 .2 \quad 1014.9 \quad 1014.3 \quad 1015.6 \quad 1015.6 \quad 10: 4.9 \quad 1013.4 \quad 1011.7 \quad 1010.4 \quad 1010.7 \quad 1012.0 \quad 1012.1$



DATE AUGUST 9TH, 1930




DATE AUGUST LOTH. 1990




## TABLE AEA PRECONTINUED WIND DATA FROM REGULARLY REPORTING STATIONS

GEEVILLE, CHASE FIELD, TX LATITUDE $\angle B \circ 2 \mathrm{ag}, \mathrm{N}$ LONGITUDE $97040 \cdot \mathrm{~W}$

OATE AUGUST BTH, 1980


 GUST (KN)

LATE AUGUST 9TH. 1980
PRESSURE(HE) $1009.5 \quad 1007.0 \quad 1006.7 \quad 1007.0 \quad 1007.5 \quad 1009.4 \quad 1006.5 \quad 1006.3 \quad 1005.4 \quad 1004.3 \quad 1003.9 \quad 1002.5$

 GUST (KN)

DATE AUGUST 10TH. 1980
PRESSURE (MB) $1001.5 \quad 1000.0 \quad 999.9 \quad 1000.9 \quad 1001.7 \quad 1002.2 \quad 1002.9 \quad 1003.4 \quad 1002.9 \quad 1003.3 \quad 100.5 \quad 1008.3$


ABLE ABEL PRCONTINUED
EA LEVEL PRESSURE AMD WIND DATA FROM REGULARLY REPORTING STATIONS

BROWNSUILLE TX


DATE AUGUST 8TH. 1980
PRESSUFEE(MB 1013.2 1012.4 1011.8 1012.1 1011.9 1011.5 1010.6 1009.4 1007.9 1008.2 1006.8 1005.6 $1012.9 \quad 1012.1012 .2 \quad 1012.2 \quad 1011.6 \quad 1010.8 \quad 1010.2 \quad 1008.5 \quad 1008.2 \quad 1007.1 \quad 1006.8 \quad 1004$


DATE AUGIJST 9TH. 1980




DATE AUGUST $20 \mathrm{TH}, 1980$
 $976.93931 .2 \quad 999.8 \quad 1000.5 \quad 1002.1 \quad 1001.2 \quad 1001.1 \quad 1001.9 \quad 1002.6 \quad 1005.8 \quad 1006$


TABLE A.l - CREONTINUED NINO DATA FROM REGULARLY REPORTING STATIONS

COLLEGE STATION TX

$$
\begin{aligned}
& \text { LATITUQF } 30035 \cdot N \text { LONGITUDE } 96022 \cdot W \\
& \text { ELEVATION } 314 F T
\end{aligned}
$$





DATE AUGUST 9rH. 1980
PRESSU日E (M13) 1012.0 1011.3 1011.0 1011.6 $1012.6 \quad 1011.8 \quad 1011.1 \quad 1009.8 \quad 1009.8 \quad 1009.5 \quad 1010.4 \quad 1010.4$ $1012.0 \quad 1011.2 \quad 1011.1 \quad 1012.6 \quad 1012.2 \quad 1011.5 \quad 1010.5 \quad 1009.5 \quad 1009.51009 .51010$
 GUST (KNi)

DATE AJGIJST IOTH, 1900
 $1009.0 \quad 1009.1 \quad 1910.2 \quad 1010.8 \quad 1011.7 \quad 1011.1 \quad 1011.1 \quad 1010.5 \quad 1010.2 \quad 1010.5 \quad 1011.6 \quad 1011$


## TABLE ABA - $\overline{\text { SEAESONTINUED }}$ CINO DATA FROM REGULARLY REPORTING STATIONS

CORPIUS CHRISTI TS
LATITUDE 2704GON LO:JGITUDE $97030 . W$
ELEVATON

DATE AULUST 8TH. 1980
 $1013.9 \quad 1013.3 \quad 1013.2 \quad 1013.9 \quad 1013.2 \quad 1012.2 \quad 1011.2 \quad 1010.2 \quad 1010.2 \quad 1009.8 \quad 1009.8 \quad 1008$


DATE AUGUST 9TH. 1980




DATE AUGUST 10 TII, 1981
 $\begin{array}{lllllllllllllllllll}994.9 & 993.2 & 994.9 & 996.3 & 999.7 & 1001.4 & 1000.0 & 1002.0 & 1002.7 & 1004.7 & 1007.8 & 1008\end{array}$


## TABLE AA.1 $\quad$ PRESONTINUED <br> WIND DATA FROM REGULARLY REPORTING STATIONS

CORPUS CHRISTI, NAS, TX LATITUDE 27042IN LONGITUDE 97 IFT17.W





DATE AUGUST 9TH, 1980


 DATE AUGUST $10 T H, 1980$



## SEALEEAEI PRESONTINUED

cotulla TX
LATIVUDE $\angle 8^{\circ} 27^{\circ} \mathrm{N}$ LONGITUDE $99013 . \mathrm{W}$


DATE AUGUST BTH. 1980
PRESSURE(MB) • $\quad 1014.4 \quad 1014.8 \quad 1014.8 \quad 1014.4 \quad 1012.5 \quad 1011.0 \quad 1010.0 \quad 1010.0 \quad 1010.7$


DATE AUGUST 9TH, 1990




DATE AUGIIST 10 TH, 1980



DATA EOTERED ONLY FOR THOSE TIMES THAT WERE AVAILABLE

## 

```
OALLAS LOVE FIELi'tX LATITUDE S2051.N LONGITUDE 96051.W
OATE AUGIST 8TH, 1980
dATE AUGuSt 9Th, 1980
```



```
dAte AUGuSt 10TH, 19BO
```



```
data entered only for those times that were available
```



FORT WORTH TX FAA

```
LATITUDE 32049'N LONGITUDE 97021.W
```



DATE AUGUST BTH. 1980


DATE AUGIJST 9TH, 1900


DATE AUGIJST IOTH. 1980


DATA E JTERED only FOR those thmes that vere available

## TABLE AA - - CONTINUED WIND dATA FROM REGULARLY REPOKTING STATIONS

```
GALVESTUN TX
            LATITUDE 29018%N LONGITUDE 94048.W
avgucr 8TH, 1930
WIND OIR(DEG)
    040
DATE AUGUST 9TH. 1980
WIND OIR(OEG)
DATE AUGUST IOTH, 1980
WIND OIN(OEG)
```



```
DATA EITERED ONLY FOR THOSE TIMES THAT NERE AVAILABLE
```


## TABLE A

KINGSVILLE NAS. TX
LATITUDE $27^{\circ} 30 \cdot \mathrm{~N}$ LONGITUDE $97049 \cdot \mathrm{~W}$


QATE AUGUST BTH. 1980
WIND OIR (OEG)
WIND SPD (KN)
GUST (KN)

DATE AUGUST 9TH. 1980



OATE AUGUST $10 T H, 1980$



## TABLE AAEA -- CONTINUEO WIND DATA FROM REGULARLY REPORTING STATIONS

KINGSVILLE, NAS, TX

$$
\begin{aligned}
& \text { LATITUDE } 27030 \cdot \mathrm{~N} \text { LONGITUDE } 97049 \cdot \mathrm{~W} \\
& \text { ELEVATION }
\end{aligned}
$$

DATE AUGUST BTH. 1930
WIND OIR(DEG)
WIND SPD
GUST (KN)

DATE AUGUST 9TH, 1980




DATE AUGUST 10TH. 1980




TABLE AAB P- CONTINUED

LUFKIN TX

$$
\begin{aligned}
& \text { LATITUDE } 1 \circ 14^{\circ} \mathrm{N} \\
& \text { ELEVATION } \\
& \text { 2BAFT }
\end{aligned}
$$



DATE AUG!ST ETH, 1980







DATE AUGUST 10TH, 1980
PRESSURE(MB) $1011.7 \quad 1011.0 \quad 1011.0 \quad 1012.1 \quad 1012.6 \quad 1012.9 \quad 1012.2 \quad 1011.2 \quad 1010.9 \quad 1010.9 \quad 1012.1010 .4$



## TAELLE A. SEA LEL CONTINUED PRESSURE AUD WIND DATA FROM REGULARLY REPORTING STATIONS

MCALLEIN TX

$$
\begin{aligned}
& \text { LATITUDE } 26^{\circ} \frac{111 N}{122 F T} \text { LONGITUDE } 98^{\circ} 14^{\circ} \mathrm{W} \\
& \text { ELEVATION }
\end{aligned}
$$

DATE AUGUST 8TH. 1980
PRESSURE(MB) $1013.8 \quad 1012.8 \quad 1012.5 \quad 1012.8 \quad 1012.8 \quad 1012.1 \quad 1010.8 \quad 1009.8 \quad 1009.7 \quad 1009.4 \quad 1008.7 \quad 1008.4$



DATE AUGUST 9TH. 1980
PRESSURE (MB) $1006.4 \quad 1004.0 \quad 1004.0$
1005.41003 .41003 .0
$\begin{array}{lrrrrrr}\text { WIND OIR(DEG) } & 340 & 340 & 330 & 350 & 030 & 330 \\ \text { WND SPD (KN) } & 13 & 15 & 13 & 15 & 18 & 15\end{array}$

DATE AUGUST $10 \mathrm{TH}, 1980$


WIND DIR(DEG)
GUST (KN)

data entered only for those times that were available

TABLE AAA P PRECONTINUED WINO DATA FROM REGULARLY. REPORTING STATIONS

MIDLANO TX

## LATITUDE $31^{\circ} 571 \mathrm{~N}^{2}$ LONGITUDE $102011^{\prime \prime} \mathrm{W}$



OATE AUGUST 8TH. 1980
PRESSURE (M13) 1012.7 1012.6 $1013.1 \quad 1013.6 \quad 1014.0 \quad 1013.6 \quad 1012.1 \quad 1010.4 \quad 1009.5 \quad 1009.1 \quad 1010.4 \quad 1011.6$



DATE AUGUST 9TH. 1980
PRESSURE(MB) $1011.6 \quad 1011.0 \quad 1011.2 \quad 1011.8 \quad 1012.1 \quad 1011.7 \quad 1010.3 \quad 1008.6 \quad 1007.4 \quad 1007.5 \quad 1008.2 \quad 1009.1$ $1011.21011 .1 \quad 1011.4 \quad 1011.6 \quad 1012.1 \quad 1011.2 \quad 1009.5 \quad 1007.6 \quad 1007.5 \quad 1007.7 \quad 1008.9 \quad 1009$
 GUST (KN)

DATE AUGUST $10 \mathrm{TH}, 1980$




TABLE AAEI $\bar{O}$ CONTINUED WIND DATA FROM REGULARLY REPORTING STATIONS

PORT ARTHUR.WSO,TX


DATE AUGUST ETH, 1930
PRESSURE (MB) $1015.2 \quad 1014.4 \quad 1014.6 \quad 1015.1 \quad 1015.6 \quad 1015.3 \quad 1014.2 \quad 101$. 0 . 0 1012.1 $1012.4 \quad 1012.6 \quad 1012.2$ $\begin{array}{llllllllllllllllllllll}1014.8 & 1014.3 & 1014.8 & 1015.3 & 1015.3 & 1014.6 & 1013.5 & 1012.5 & 1012.2 & 1012.1\end{array}$


DATE AUGUST 9TH. 1930
PRESSUPE (MB) $1011.2 \quad 1010.8 \quad 1011.0 \quad 1011.2 \quad 1012.2 \quad 1012.3 \quad 1011.4 \quad 1010.6 \quad 1010.2 \quad 1010.4 \quad 1010.8 \quad 1011.1$ $1010.8 \quad 1010.41011 .0 \quad 1012.2 \quad 1012.3 \quad 1011.9 \quad 1011.3 \quad 1010.4 \quad 1010.0 \quad 1010.8 \quad 1010.8 \quad 1010$
 gUST (KN)

OATE AUGUST IOTH, 1990




TABLE AGA - PRESONTINUED WIND DATA FROM REGULARLY REPORTING STATIONS
SEA LELE

SAN ANTONIO TX

$$
\begin{aligned}
& \text { LATITUDE } 29032 \% \text { N } \\
& \text { ELEVATION } \\
& 794 \mathrm{~F} T
\end{aligned}
$$

AUGUST 8TH. 1980
PRESSURE (MB) $1015.3 \quad 1015.0 \quad 1015.1 \quad 1015.4 \quad 1015.6 \quad 1015.0 \quad 1013.2 \quad 1011.8 \quad 1010.8 \quad 1010.5 \quad 1011.7 \quad 1012.1$ $\begin{array}{lllllllllllllllllllll}1015.3 & 1015.1 & 1015.1 & 1015.6 & 1015.0 & 1013.9 & 1012.6 & 1011.1\end{array}$


OATE AUGUST 9TH. 1980




DATE AUGUST $10 \mathrm{TH}, 1980$
 $1007.0 \quad 1006.1005 .3 \quad 1006.7 \quad 1008.1 \quad 1007.0 \quad 1005.6 \quad 1005.0 \quad 1005.6 \quad 1007.0 \quad 1008.5 \quad 1009$
 GUST (KN)



OATE AUGUST 8TH. 1980



DATE AUGUST 9TH. 1980
 $\begin{array}{lllllllllllllllllllllll}1011.4 & 1010.5 & 1009.9 & 1010.9 & 1010.5 & 1010.4 & 1008.8 & 1008.4 & 1007.6 & 1007.8 & 1008\end{array}$


DATE AJGUST 10 TH . 1980




TABLEAAEG PRESONTINUED WIND OATA FROM REGULARLY REPORTING STATIONS
victoria rx

$$
\begin{aligned}
& \text { LATITUDE } \& 8051 * V \\
& \text { ELEVATION } \begin{array}{l}
\text { IO } \\
\text { FFT }
\end{array} \text { LONGITUDE } 96055 * W .
\end{aligned}
$$

DATE AJGUST 8TH. 1980
PRESSURE(MB) $1014.9 \quad 1014.2 \quad 1014.5 \quad 1014.9$ 1014.9 $1014.2 \quad 1012.8 \quad 1011.5 \quad 1011.3 \quad 1011.8 \quad 1010.8 \quad 1011.1$ $1014.5 \quad 1014.2 \quad 1014.7 \quad 1015.2 \quad 1014.3 \quad 1013.3 \quad 1012.5 \quad 1011.1 \quad 1011.3 \quad 1011.3 \quad 1011.5 \quad 1010$





DATE AUGUST 10TH, 1980



TABLEAAA PE CONTINUED WIND DATA FROM REGULARLY REPORTING STATIONS

WACO TX

$$
\begin{aligned}
& \text { LATITUDE } 31037 \text { NN LONGITUDE } 97013 . \mathrm{W} \\
& \text { ELEVATION } 500 \mathrm{~F} \text { T }
\end{aligned}
$$

DATE AUGUST 8TH. 1980




DATE AUGUST 9TH. 1980
PRESSURE(ME) 1012.4 $1012.4 \quad 1012.4 \quad 1013.0 \quad 1013.3 \quad 1013.3 \quad 1011.5 \quad 1010.2 \quad 1009.8 \quad 1009.5 \quad 1011.0 \quad 1010.6$



DATE AUGUST 10TH, 1980
PRESSURE (MB) 1010.6 1010.1 $1010.1 \quad 1011.2 \quad 1011.9 \quad 1012.5 \quad 1011.5 \quad 1010.8 \quad 1010.8 \quad 1010.5 \quad 1010.8 \quad 1011.5$ $1010.1 \begin{array}{llllllllllllllllll} & 1010.1 & 1010.5 & 1011.5 & 1012.1 & 1012.1 & 1010.8 & 1010.8 & 1010.5 & 1010.5\end{array}$



|  | datestime | $\mathrm{Sat}_{\text {Storm }}^{\text {LAT }}$ | CENTEN | SEA LEVEL PRESSURE (MB) | $\begin{aligned} & \text { FLT.LVL } \\ & \text { WIND } \\ & \text { (KN) } \end{aligned}$ | $\begin{aligned} & \text { EYEMETER } \\ & \text { CIAMEMI) } \\ & \text { (N.MI } \end{aligned}$ | Remarks |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| g | $7 \quad 1920$ | 2215 | 8815 |  | 105 | 10 | WIND $160 / 64 \mathrm{KN}$ 3UN.MI FROM CENTER |
|  | 7 2045 <br> 7 23 | 22 22 20 31 | 88 83 80 | $\begin{array}{r}937 \\ 945 \\ \hline 95\end{array}$ | 67 | 10 | WALL CLOUD 15N.MI THICK |
|  | 8 8 8 | 2248 <br> 23 <br> 18 | 8988 | 946 | 182 | 10 |  |
|  | $\begin{array}{ll}8 & 431 \\ 8 & 544\end{array}$ | 23 12 |  | 950 | ${ }_{80}^{82}$ |  |  |
|  | 8 ${ }^{8} 844$ | ${ }^{23} 18$ | 9124 | 959 | 71 | 20 |  |
|  | $\begin{array}{ll}88 & 1125 \\ 8 & 1436\end{array}$ | 23 <br> 24 <br> 24 <br> 24 | $912{ }^{42}$ | 949 | 110 120 |  |  |
|  | 81719 | 24 24 24 24 24 3 | 9250 | 913 <br> 914 <br> 18 | 145 107 | 10 12 |  |
|  | $\begin{array}{lll}8 \\ 8 & 2000 \\ 8 & 2100\end{array}$ | 24 24 24 42 | 9330 | ${ }^{914}$ |  | 12 |  |
|  | 82300 | 2454 <br> 2459 <br> 25 | 94 <br> 90 <br> 14 <br> 14 |  | 103 | 7 |  |
|  | 8 <br> 9 <br> 9358 | - 245 | ${ }^{94} 14$ | 916 | 193 | $10 \times 8$ |  |
|  | $\begin{array}{ll}9 & 351 \\ 9 & 427\end{array}$ | ${ }^{25} 515$ | 94 99 |  | 104 103 |  |  |
|  | 9510 | 2514 | 9511 | 916 | $\begin{array}{r}92 \\ 92 \\ \hline\end{array}$ | $10 \times 7$ |  |
|  | 96816 | 2513 <br> 25 <br> 14 | 955 |  | 92 |  |  |
|  | 9811 | ${ }^{255}$ | 95 <br> 958 <br> 98 | 920 | ${ }_{9} 9$ |  | clouds fili. the eye |
|  | 91210 |  | 9610 | 920 | 100 | ¢ ${ }_{10}$ |  |
|  |   <br> 9 1436 <br> 9  <br> 174  | 25 25 25 | 968 | 929 | 93 |  | MAX WIN WINO S'SKN |
|  | 9. 2039 | 25 23 | $\begin{array}{r}96 \\ 96 \\ 96 \\ \hline 9\end{array}$ |  | 88 | $20 \times 12$ | HUKRICANE iWINIS EXTEND LIENMI DUE NORTH |
|  | 9 ${ }^{2} 236$ | 2605 2604 | 9745 | 945 | 60 | 10 | EyE over land |
|  | 10 | 2604 2607 2609 | $9717 * *$ |  | 60 80 |  | EYE OVER LAND |

[^0]


SEA LEVFL PRFASUREA ANO - WIND DATA FROM SHIP REPORTS
LAT LONG WIND DPD PRESSURE

## $1500 C S T$ BTH AUGUST

| 25 | 48 | 86 | 6 | 110 | 18 | 1013.0 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 26 | 0 | 93 | 30 | 070 | 31 | 1007.0 |
| 21 | 30 | 92 | 54 | 270 | 36 | 0999.9 |
| 22 | 0 | 93 | 12 | 270 | 30 | 0999.7 |


$2100 C S T$ 8TH AUGUST

| 26 | 0 | 93 | 30 | 090 | 33 | 1005.9 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 25 | 54 | 89 | 42 | 150 | 31 | 100778 |
| 21 | 48 | 92 | 54 | 240 | 23 | $10022: 2$ |
| 25 | 54 | 96 | 24 | 350 | 45 | 1003 |
| 27 | 30 | 96 | 6 | 040 | 40 | 100508 |
| 27 | 36 | 91 | 47 | 080 | 40 | $1008: 5$ |
| 27 | 48 | 90 | 30 | 100 | 35 | 1012.0 |
| 28 | 42 | 95 | 18 | 040 | 02 | 1010.0 |

0000CST GTH AUGUST

| 25 | 30 | 85 | 4 A | 110 | 16 | 1016.3 |
| :--- | ---: | :--- | :--- | :--- | :--- | :--- |
| 26 | 0 | 93 | 30 | 100 | 37 | 10055.0 |
| 21 | 54 | 9248 | 210 | 30 | 1002.0 |  |
| 27 | 0 | 89 | 30 | 130 | 37 | 1010.7 |
| 25 | 54 | 89 | 42 | 150 | 29 | 1008.8 |
| 28 | 42 | 95 | 18 | 040 | 02 | 10080.4 |
| 28 | 0 | 90 | 6 | 100 | 30 | 1011.4 |



LAT LONG WIND | WIR SPD PRESSURE |
| :---: |
|  |
|  |

1200CST 9TH AUGUST

| 23 | 12 | 91 | $4 A$ | 150 | 34 | 1006.4 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 23 | 42 | 87 | 24 | 140 | 04 | 1010.1 |
| 25 | 6 | 87 | 54 | 130 | 14 | 1012.2 |
| 28 | 42 | 95 | 18 | 060 | 02 | 1007.0 |
| 27 | 30 | 90 | 12 | 120 | 20 | 1014.0 |
| 27 | 18 | 91 | 12 | 130 | 32 | 1010.6 |
| 27 | 54 | 94 | 30 | 090 | 40 | 1003.0 |
| 27 | 54 | 92 | 12 | 130 | 33 | 1002.0 |

1500CST 9TH AUGUST

| 25 | 54 | 89 | 42 | 150 | 19 | 1010.4 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 26 | 0 | 93 | 30 | 110 | 31 | 1007.3 |
| 24 | 24 | 94 | 0 | 160 | 40 | 0998.5 |
| 23 | 30 | 91 | 30 | 150 | 34 | 1006.1 |
| 27 | 12 | 89 | 54 | 120 | 24 | 1011.5 |
| 27 | 54 | 94 | 36 | 110 | 40 | 1002.4 |
| 28 | 42 | 95 | 14 | 050 | 02 | 1006.3 |

1800 CST 9TH AUGUST

| 26 | 0 | 93 | 30 | 120 | 29 | 1007.1 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 27 | 18 | 90 | 30 | 130 | 32 | 1012.4 |
| 26 | 54 | 90 | 36 | 130 | 19 | 1011.9 |
| 25 | 54 | 89 | 42 | 140 | 21 | 1009.3 |
| 27 | ${ }^{6}$ | 8 A | $1 ?$ | 090 | 05 | 1011.0 |
| 27 | 54 | 86 | 1 A | 110 | 13 | 1014.0 |
| 23 | 42 | 91 | 18 | 120 | 27 | 1005.8 |
| 24 | 36 | 88 | 30 | 090 | 08 | 1007.3 |
| 23 | 54 | 86 | 1 A | 120 | 15 | 1009.8 |
| 24 | 24 | 86 | 36 | 100 | 15 | 1012.5 |
| 24 | 18 | 86 | 6 |  |  |  |
| 28 | 42 | 95 | 1 1A | 060 | 02 | 1006.1 |
| 27 | 24 | 91 | 18 | 120 | 32 | 1008.2 |
| 27 | 42 | 94 | 30 | 090 | 40 | 1002.0 |
| 210 | OCST |  | H |  |  |  |
| 25 | 54 | 89 | 42 | 140 | 19 | 1010.7 |
| 26 | 0 | 93 | 30 | 120 | 31 | 1007.7 |
| 27 | 12 | 90 | 4 A | 140 | 29 | 1010.0 |
| 27 | 12 | 91 | 12 | 120 | 18 | 1010.8 |
| 24 | 12 | 91 | 16 | 110 | 31 | 1006.9 |
| 28 | 42 | 95 | 1 A | 070 | 02 | 1006.8 |

SEA LEVEL PRESSURE AND WIND DATA FROM SHIP REPORTS
LAT LONG WIND SPO PRESSURE

## OOOOCST 10 TH AUGUST

| 27 | 0 | 90 | 6 | 120 | 21 | 1010.1 |
| :--- | ---: | :--- | :--- | :--- | :--- | :--- |
| 26 | 0 | 93 | 30 | 110 | 27 | $1008: 5$ |
| 25 | 54 | 89 | 42 | 140 | 17 | 1010.6 |
| 22 | 54 | 86 | 6 | 080 | 17 | 1011.5 |
| 25 | 30 | 88 | 42 | 130 | 18 | 10100.6 |
| 28 | 42 | 95 | 18 | 070 | 02 | 1006.5 |
| 27 | 42 | 91 | 36 | 130 | 30 | 1012.9 |

O300CST 10TH AUGUST

| 26 | 0 | 93 | 30 | 120 | 29 | 1007.4 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 25 | 54 | 89 | 42 | 150 | 17 | $1009: 8$ |
| 28 | 42 | 95 | $1 \cap$ | 080 | 02 | 1006.8 |

0600CST 10 TH AUGUST

| 2554 | 89 | 43 | 150 | 17 | 1010.4 |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 2642 | 85 | 12 | 110 | 39 | 1011.9 |
| 2636 | 88 | 42 |  |  |  |
| 26 | 93 | 30 | 120 | 25 | 1008.0 |
| 2636 | 94 | 30 | 150 | 25 | 1004.4 |
| 2436 | 85 | 30 | 140 | 13 | 1011.3 |
| 256 | 87 | 12 | 130 | 18 | 1012.0 |
| 256 | 86 | 3 | 140 | 12 | 1012.0 |
| 2418 | 87 | 24 | 130 | 10 | 1011.1 |
| 2254 | 86 | 1 B | 110 | 20 | 1010.2 |
| 2842 | 95 | 18 | 080 | 02 | 1006.8 |
| 2724 | 93 | 24 | 120 | 25 | 1008.1 |
| 0 | 92 | $4\rangle$ | 130 | 37 | 1011.5 |
| 220 | 85 | 36 | 140 | 14 | 1009.4 |
| 2154 | 85 | 36 | 120 | 09 | 1009.5 |

0900 CST 10TH AUGUST

| 2554 | 89 | 42 | 150 | 17 | 1011.3 |
| :--- | :--- | :--- | :--- | :--- | :--- |
| 26 | 18 | 87 | $4 \Omega$ | 130 | 17 |
| 26 | 0 | 93 | $3 n$ | 130 | 25 |
| 26 | 30 | 94 | 36 | 120 | 25 |
| 28 | 42 | 95 | 18 | 090 | 02 |
|  | 100501 |  |  |  |  |
|  |  |  | 1008.5 |  |  |

1200CST 10 TH AUGUST

| 26 | 0 | 93 | 30 | 130 | 19 | 1010.5 |
| ---: | ---: | ---: | ---: | ---: | ---: | ---: |
| 26 | 48 | 93 | 0 | 150 | 27 | 1011.0 |
| 26 | 30 | 95 | 0 | 120 | 20 | 1005.8 |
| 25 | 54 | 89 | 42 | 140 | 14 | 1011.9 |

SEA LEVFL PRFSBLEREAANT WIND DATA FROM SHIP REPORTS

| LAT | LONG | WIND DIR | $S P D$ | $\begin{gathered} \text { PRFSSURE } \\ \text { (MB) } \end{gathered}$ |
| :---: | :---: | :---: | :---: | :---: |

1200CST 10 TH AUGUST

| 24 | 24 | 85 | 30 | 100 | 10 | 1012.0 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 23 | 18 | 86 | 18 | 110 | 10 | 1012.1 |
| 28 | 42 | 95 | 18 | 090 | 02 | 1009.1 |
| 28 | 42 | 93 | 42 | 150 | 30 | 1013.3 |
| 27 | 30 | 92 | 6 | 120 | 16 | 1008.2 |
| 27 | 36 | 90 | 36 | 140 | 16 | 1013.2 |
| 21 | 42 | 95 | 30 | 110 | 15 | 1010.0 |

$1500 C S T 10 T H$ AUGUST

| 25 | 54 | 89 | 42 | 120 | 12 | 1010.6 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 26 | 0 | 93 | 30 | 130 | 19 | 1010.1 |
| 26 | 54 | 94 | 42 | 100 | 24 | 1007.1 |
| 28 | 42 | 95 | 18 | 100 | 02 | 1009.2 |

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[^0]:    ** STORM CENTER FIXED BY RAOAR

