RESISTIVITY STRUCTURE ACROSS THE HUMBOLDT RIVER BASIN, NORTH-CENTRAL NEVADA

by

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ABSTRACT

Magnetotelluric data collected along five profiles show deep resistivity structures beneath the Battle Mountain-Eureka and Carlin gold trends in north-central Nevada, which appear consistent with tectonic breaks in the crust that possibly served as channels for hydrothermal fluids. It seems likely that gold deposits along these linear trends were, therefore, controlled by deep regional crustal fault systems.

Two-dimensional resistivity modeling of the magnetotelluric data generally show resistive (30 to 1,000 ohm-m) crustal blocks broken by sub-vertical, two-dimensional, conductive (1 to 10 ohm-m) zones that are indicative of large-scale crustal fault zones. These inferred fault zones are regional in scale, trend northeast-southwest, north-south, and northwest-southeast, and extend to mid-crustal (20 km) depths. The conductors are about 2- to 15-km wide, extend from about 1 to 4 km below the surface to about 20 km depth, and show two-dimensional electrical structure. By connecting the locations of similar trending conductors together, individual regional crustal fault zones within the upper crust can be inferred that range from about 4- to 10-km wide and about 30- to 150-km long. One of these crustal fault zones coincides with the Battle Mountain-Eureka mineral trend. The interpreted electrical property sections also show regional changes in the resistive crust from south to north. Most of the subsurface in the upper 20 km beneath Reese River Valley and southern Boulder Valley are underlain by rock that is generally more conductive than the subsurface beneath Kelly Creek Basin and northern Boulder Valley. This suggests that either elevated-temperature or high-salinity fluids, alteration, or carbonaceous rocks are more pervasive in the more conductive area (Battle Mountain Heat-Flow High), which implies that the crust beneath these valleys is either more fractured or has more carbonaceous rocks than in the area surveyed along the 41st parallel.

INTRODUCTION

Many sediment-hosted gold deposits occur along linear trends in northern Nevada (Struhsacker and others, 1996). The distribution and genesis of these deposits along the Battle Mountain-Eureka (BME) and Carlin gold trends is not fully understood. In general, most geologic models suggest that regional structures played an important role in the spatial distribution of these deposits (e.g. Arehart and others, 1993; Ilchik and Barton, 1997; Radtke, 1985; Shawe, 1991; Sillitoe and Bonham, 1990; Tosdal, 1998). To investigate crustal structures that may be related to the genesis of gold deposits along these trends, five profiles of magnetotelluric (MT) soundings were acquired in 1998, 1999, and 2000 (lines MT5 to MT9, Figure 1). Modeling of the MT data can be used to infer the deep resistivity structure of the crust. Such structures may reflect possible

tectonic controls on the emplacement of mineral deposits along these linear trends, and so may be used to help improve critical gold endowment estimates in the Humboldt River Basin.

MAGNETOTELLURIC METHOD

The magnetotelluric (MT) method is a passive surface geophysical technique, which uses the earth's natural electromagnetic fields to investigate the electrical resistivity structure of the subsurface. The resistivity of geologic units is largely dependent upon their fluid content, porosity, degree of fracturing, temperature, and conductive mineral content (Keller, 1989). Saline fluids within the pore spaces and fracture openings can reduce resistivities in a resistive rock matrix. Also, resistivity can be lowered by the presence of conductive clay minerals, carbon, and metallic mineralization. It is common for altered volcanic rocks to contain authigenic minerals that have resistivities ten times lower than those of the surrounding rocks (Nelson and Anderson, 1992). Increased temperatures cause higher ionic mobility and mineral activation energy, reducing rock resistivities significantly. Unaltered, unfractured igneous rocks are normally very resistive (typically 1,000 ohm-m or greater), whereas fault zones will show low resistivity (less than 100 ohm-m) when they are comprised of rocks fractured enough to have hosted fluid transport and consequent mineralogical alteration (Eberhart-Phillips and others, 1995). Carbonate rocks are moderately to highly resistive (hundreds to thousands of ohm-m) depending upon their fluid content, porosity, fracturing, and impurities. Marine shales, mudstones, and clay-rich alluvium are normally very conductive (a few ohm-m to tens of ohm-m). Unaltered, metamorphic rocks (non-graphitic) are moderately to highly resistive (hundreds to thousands of ohm-m). Tables of electrical resistivity for a variety of rocks, minerals and geological environments may be found in Keller (1987) and Palacky (1987).

The MT method can be used to probe the crust from depths of tens of meters to depths of tens of kilometers (Vozoff, 1991). Natural variations of the Earth's magnetic and electric field are measured and recorded at each MT station. The primary frequency bands used by the MT method are 10,000 Hz to 1 Hz from worldwide lightning activity and 1 Hz to 0.0001 Hz from geomagnetic micro-pulsations. The natural electric and magnetic fields propagate vertically in the earth because the very large resistivity contrast between the air and the earth causes a vertical refraction of both fields transmitted into the earth (Vozoff, 1972).

The natural electric and magnetic fields are recorded in two orthogonal, horizontal directions. The vertical magnetic field ("tipper") is also recorded. The resulting time-series signals are used to derive the tensor apparent resistivities and phases. First, the signals are converted to complex cross-spectra using

FFT (fast-Fourier-transform) techniques. Then, least-squares,
cross-spectral analysis (Bendat and Piersol, 1971) is used to
solve for a transfer function that relates the observed electric
fields to the magnetic fields under the assumption that the Earth
consists of a two-input, two-output, linear system with the
magnetic fields as input and the electric fields as output
(Rodriguez and others, 1996b). Prior to conversion to apparent
resistivity and phase, the electromagnetic tensor is normally
rotated into principal directions that correspond to the
direction of maximum and minimum apparent resistivity. For a
two-dimensional (2-D) Earth, the MT fields can be de-coupled into
transverse electric (TE) and transverse magnetic (TM) modes; 2-D
modeling is generally done to fit both modes. When the geology
satisfies the 2-D assumption, the MT data for the TE mode is for
the electric field parallel to geologic strike, and the data for
the TM mode is for the electric field across strike. The MT
method is well suited for studying complicated geological
environments because the electric and magnetic relations are
sensitive to vertical and horizontal variations in resistivity.
The method is capable of establishing whether the electromagnetic
fields result from subsurface terranes of effectively 1-, 2-, or
3-dimensions. An introduction to the MT method and references
for a more advanced understanding are contained in Dobrin and
Savit (1988) and Vozoff (1991).

MAGNETOTELLURIC SURVEYS
     Thirty-eight MT soundings were located along five profiles
(MT5 to MT9, Figure 1) of lengths varying from 20 to 110 km and
spacing that varied from 1.5 to 13.5 km.   The MT soundings
reported here span from 117.48972 degrees west and 41.14575
degrees north to 116.22838 degrees west and 40.58591 degrees
north. Results of sixty-nine other MT soundings located along
four profiles (MT1 to MT4, Figure 1) of lengths varying from 40
to 110 km and spacing that varied from 1.3 to 19.5 km are
described in Rodriguez and Williams (2001).
     The following table lists the MT station locations of the
five profiles (MT5 to MT9). These locations were found using
either on-site GPS measurements or digitized from 100,000 scale
field maps (Williams and Rodriguez, 2001b; 2001c; 2001d).
Coordinates are referenced to the 1866 Clarke spheroid and North
American 1927 Western United States datum. Longitude and
latitude format below is decimal degrees. Elevation (derived
from 100,000 scale field maps) is in meters.

Station	Longitude	Latitude	Elev(m)
		MT5	
53	-117.46039	41.03561	1450
61	-117.37602	41.05213	1430
48	-117.23083	41.04530	1340
62	-117.12730	41.03736	1350

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45          -116.99049      40.99599   1420
60          -116.95005      41.01850   1460
43          -116.88703      41.00584   1650
42          -116.80737      41.00304   1600
59          -116.64933      40.98375   1530
58          -116.54706      40.98970   1650
57          -116.45419      41.00731   1700
34          -116.41038      40.98637   1600
37          -116.39134      40.98808   1650
33          -116.36519      40.98794   1700
32          -116.33897      40.99409   1800
35          -116.26639      40.99996   1800
56          -116.22838      40.99611   1740
                        MT6
72          -116.65770      40.76352   1400
67          -116.59319      40.76856   1400
65          -116.52693      40.76426   1410
64          -116.49056      40.75211   1420
                        MT7
63          -116.46887      40.69744   1450
66          -116.49212      40.79785   1415
68          -116.50533      40.84150   1430
69          -116.49231      40.89135   1470
70          -116.47627      40.94122   1530
                        MT8
76          -116.98241      40.58591   1400
75          -116.96516      40.60841   1400
74          -116.90120      40.66956   1374
73          -116.86531      40.70358   1463
                        MT9
90          -117.48972      40.74573   1440
91          -117.41152      40.78297   1390
92          -117.38475      40.83747   1390
93          -117.31868      40.87867   1370
94          -117.27735      40.93117   1340
49          -117.16232      41.08410   1360
95          -117.04928      41.13012   1420
96          -116.95985      41.14575   1420
97          -117.17020      40.79370   1400
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MAGNETOTELLURIC DATA

Frequencies sampled for profiles MT5 to MT9 ranged from 0.002 to 200 Hz using single station recordings of both orthogonal horizontal components of the electric and magnetic fields, along with the vertical magnetic field at stations 37, 48, 49, 53, 59, 63 to 69, 73 to 76, and 90 to 97 (Williams and Rodriguez, 2001b; 2001c; 2001d). The single station recordings for profiles MT1 to MT4 are reported in Williams and Rodriguez (2000; 2001a) and Williams and others (2001a; 2001b; 2001c) along

with the vertical magnetic field recordings at stations 12A to 28A, 1, 11, 14, 71, 77, 79 to 89, 98, 108, and 109. Sampling this frequency range in other areas of widely varying geology has allowed us to probe the crust from depths of hundreds of meters to depths of tens of kilometers (Eberhart-Phillips and others, 1990; Stanley and others, 1991; 1997; Rodriguez and others, 1996a; 1996b). The recorded time-series data were transformed to the frequency domain and processed to determine a two-dimensional apparent resistivity and phase tensor at each site. Rotation of the impedance tensor to maximum and minimum directions allows for decoupling into the TE and TM modes.

A measure of the dimensionality (2- or 3-D) of MT data is provided by the impedance skew of the impedance tensor (Vozoff, 1972) and the impedance polar plots (Reddy and others, 1977). By examining the tensor impedances, the dimensionality of electrical structures and the general strikes of 2-D structures were estimated for different frequencies. MT stations whose impedances were 3-D in the lower frequencies (Williams and Rodriguez, 2001; Williams and others, 2001a; 2001b; 2001c) are indicated with triangle symbols in Figure 1. Predicted values of the electric field can be computed from the measured values of the magnetic field (Vozoff, 1991). The coherence of the predicted electric field with the measured electric field is a measure of the signal-to-noise ratio provided in the multiple coherency plot. Values are normalized between 0 and 1, where values at 0.5 signify signal levels equal to noise levels. For this data set, coherencies were generally at an acceptable level, except at times in the "dead band" (0.1 to 1 Hz) and at times in the lower frequencies (0.002 to 0.1). Overall data quality was fair to good for profiles MT6, MT7, MT8, and MT9, and fair to poor for MT5 (Williams and Rodriguez, 2001; Williams and others, 2001a; 2001b; 2001c). The worst data points were ignored and the data used in the modeling generally had predicted coherencies above 90% and 60% for the higher and lower quality data, respectively.

In the Appendix, observed data of each MT sounding for each profile (MT5 to MT9) are represented by descrete values (circle and x symbols) of the TE and TM modes from the raw data curves (Williams and Rodriguez, 2001; Williams and others, 2001a; 2001b; 2001c). Calculated data for each profile are represented by solid (TE mode) and dashed (TM mode) lines interpolated from the descrete values (0.003, 0.01, 0.03, 0.1, 0.3, 1, 3, 10, 30, 100) output from the 2-D finite element models (Figures 2 and 3).

RESISTIVITY MODELS

Wannamaker (1983) found that MT responses in the northern Basin and Range are fundamentally 3-D in nature. However, for elongated structures, 2-D modeling can be used to construct reasonable estimates of the resistivity cross-sections along each profile (Figures 2 and 3). Wannamaker and others (1984)

demonstrated that approximating 3-D structure beneath a centrally located profile with 2-D modeling is best achieved when fitting the TM curve even at the expense of a poor fit of the TE curve. However, because TM data are relatively insensitive to the depth extent of a subsurface body (Eberhart-Phillips and others, 1995), the depths to the base of the bodies in the model are not well constrained. Hence, clarifying the model limits with 3-D resistivity modeling may be necessary.

The 2-D resistivity models of Figures 2 and 3 were constructed for each profile using the forward modeling finite element algorithm of Wannamaker and others (1987). The resistivity models generally fit the TM data better than the TE data (see Appendix A, B, C, D, and E), although fits to the TE data were satisfactory for stations where 2-D structure was indicated (Figure 1). However, because of the widespread 3-D character of the results, we focus only on the gross structure determined by the models.

The variable-dimension finite-element grid cells used for our models had 127 horizontal and 38 vertical nodes (plus 9 more vertical nodes used for the air layer). The edges of the model were extended to over 800-km horizontally and 450-km vertically, so as to minimize edge effects. The resolution of the resistivity boundaries used for each model is somewhat subjective. If different resistivities are used, then boundary positions and layer depths would have to be adjusted to achieve similar fits to the observed data. The extreme case would be to use a model with a "continuous" resistivity gradient from low to high resistivities. The resolution of the resistivity boundaries is also, in part, a function of our model grid mesh design. We have attempted to keep each model simple. The MT profile model (Figures 2 and 3) depths are relative to the surface.

DISCUSSION

The deeper section of the 2-D resistivity model for profile MT4 (Figure 2) was reported previously in (Rodriguez and Williams, 2001), but did not include a description of the upper 2 km, which is important for gold endowment estimates in the Humboldt River Basin. We, therefore, present our interpretation of the upper 2 km here, following a summary of the deeper section. The sub-vertical, 2-D, north trending (Figure 4) conductor (1 to 10 ohm-m) beneath stations 22A to 24A that penetrates to mid-crustal depths (from 5 to 20 km depth) is hard to explain other than by a major crustal-dimension fault or fracture zone (Eberhart-Phillips and others, 1995). The low resistivities (conductors) are likely caused by material associated with faulting or fracture filling, such as mylonitic breccia, brine-filled fractures, argillaceous alteration from hydrothermal fluids, substantial graphitic carbon associated with shearing, or a combination of these (Eberhart-Phillips and others, 1995). Additional support for a fault zone

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interpretation is the 2-D electrical structure. In the upper 8
km, the resistive (30 to 300 ohm-m) crustal blocks near Battle
Mountain appear to be broken by a separate northeast trending
(Figure 4) conductor (10 ohm-m beneath station 16A, Figure 2)
that is characteristic of a crustal fault zone. A separate
northeast trending conductor (3 ohm-m) beneath station 12A
(Figure 2) is also characteristic of a crustal fault zone. Each
of these inferred crustal faults occur within the BME gold trend.
A north trending (Figure 4) conductor (1 to 10 ohm-m) within the
Carlin gold trend beneath station 25A (Figure 2) is also
characteristic of a crustal fault zone. The large 3-D conductive
(1 ohm-m) zone from 1 to 12 km depth between stations 27A and 28A
(between the BME and Carlin trends) is characteristic of a
geothermal system or a thick section of carbonaceous rocks. The
upper 2 km is characterized primarily by conductive (1 to 30 ohm-
m) basin fill (Reese River Valley and Boulder Valley basins).
The resistive (100 to 1,000 ohm-m) crustal blocks within the
northern Nevada rift (Zoback and others, 1994) between stations
25A and 26A are characteristic of lower Paleozoic rocks and
intruded lower Paleozoics reported by Rodriguez and Williams
(2001). The northwest trending (Figure 5) conductor (3 to 10
ohm-m) near the Good Hope fault within the Carlin trend beneath
station 27 appears to be a local fault zone (Figure 6) and not a
crustal fault zone, because its electrical character at lower
depths is 3-D.

     The deeper southwestern half of the 2-D resistivity model
for profile MT8 (Figure 2) is dominated by conductive (10 ohm-m)
rocks we interpret to be part of the Battle Mountain Heat-Flow
High (Sass and others, 1981). The entire profile lies within the
BME trend. The northeastern half is characterized by a northwest
trending (Figure 4) conductor (1 to 10 ohm-m) beneath stations 73
and 74 that is characteristic of a crustal fault zone, which also
continues in the upper 2 km (Figure 2). The southwestern half of
the shallow section appears to be characterized by resistive (30
to 300 ohm-m) rocks that are characteristic of lower Paleozoics
or other unknown volcanic and/or clastic sedimentary rocks at
depth (Grauch and others, 1998).

     The 2-D resistivity model for profile MT6 (Figure 2) shows a
northeast trending (Figure 4) conductor (3 to 10 ohm-m) between
the BME and Carlin trends beneath station 65 that is
characteristic of a crustal fault zone, which also continues in
the upper 2 km. The large 3-D conductive (1 ohm-m) zone from
about 1.5 to 15 km depth also between the BME and Carlin trends
beneath stations 64 and 28A (Figure 2) is characteristic of a
geothermal system or a thick section of carbonaceous rocks. The
western edge of the deeper section beneath station 72 shows
resistive (30 to 100 ohm-m) rocks that are characteristic of
intruded lower Paleozoics in the northern Nevada rift (NNR).
Other conductive (3 to 10 ohm-m) zones in this model that are 3-D

in the their deeper character are not interpreted as crustal fault zones.

The 2-D resistivity model for profile MT7 (Figure 2) shows a northeast trending (Figure 4) conductor (3 to 10 ohm-m) near the southwestern edge of the Carlin trend beneath stations 68 and 69 that is characteristic of a crustal fault zone, which also continues in the upper 2 km. The large 3-D conductive (1 ohm-m) zone from about 0.5 to 15 km depth between the BME and Carlin trends beneath stations 63, 64 and 28A (Figure 2) is characteristic of a geothermal system or a thick section of carbonaceous rocks. The northern edge of the deeper section beneath station 70 within the Carlin trend shows resistive (100 ohm-m) rocks that are characteristic of lower Paleozoics.

The 2-D resistivity model for profile MT5 (Figure 3) shows a northeast trending (Figure 4) conductor (1 to 10 ohm-m) within the BME trend beneath stations 48 and 62 that is characteristic of a crustal fault zone. The western end of the model, within the BME trend, beneath stations 53 and 61 shows resistive (300 ohm-m) rocks that are characteristic of Precambrian rocks that outcrop nearby (Stewart and Carlson, 1976). The central half of the deeper section between the BME and Carlin trends beneath stations 43, 42, and 59 shows resistive (300 ohm-m) rocks that are characteristic of intruded lower Paleozoics in the NNR. Other resistive (30 to 300 ohm-m) rocks beneath stations 60, 58, and 57 are also characteristic of lower Paleozoic rocks. The eastern half of the model shows a northeast trending conductor (10 to 30 ohm-m) within the Carlin trend beneath station 34 and a north-northeast (Figure 4) trending conductor (3 to 30 ohm-m) on the northeastern edge of the Carlin trend beneath station 56 (Figure 3). Both of these conductors, characteristic of crustal fault zones, continue in the upper 2 km. Resistive (30 to 1,000 ohm-m) rocks within the Carlin trend beneath stations 37, 33, 32, and 35 are characteristic of lower Paleozoic rocks. The upper 2 km of section over the entire profile length is dominated by resistive (30 to 1,000 ohm-m) rocks that are characteristic of lower Paleozoic rocks, but also includes shallow conductive (3 to 30 ohm-m) basin fill for Paradise Valley (0.5 km), Kelly Creek basin (1 km) and Boulder Valley (0.5 to 1 km).

The 2-D resistivity model for profile MT9 (Figure 3) shows a north trending (Figure 4) conductor (1 to 10 ohm-m) within the BME trend beneath station 94 that is characteristic of a crustal fault zone. The northeastern half of the model shows two northeast trending (Figure 4) conductors (1 to 30 ohm-m) on the northeast edge of the BME trend beneath stations 48, 49, and 95 (Figure 3). Since the model profile is also trending northeast, these northeast trending conductors are characteristic of and probably sub-parallel to a single crustal fault zone (Figure 7). Resistive (30 to 1,000 ohm-m) rocks within the BME trend beneath stations 90 and 91, in the upper 8 km beneath station 93, northeast of station 49, and in the upper 5 km beneath station 95

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(Figure 3), are characteristic of lower Paleozoic rocks. Other
conductive (1 to 10 ohm-m) zones in this model that are 3-D in
the their deeper character are not interpreted as crustal fault
zones. The upper 2 km of section over the entire profile length
is dominated by resistive (30 to 1,000 ohm-m) rocks that are
characteristic of lower Paleozoic rocks, but also includes
shallow conductive (3 to 30 ohm-m) basin fill for Pumpernickel
Valley (about 0.4 km) and Kelly Creek basin (about 0.3 to 1 km).

The crustal conductors, in general, are not well constrained
laterally due to wide station spacing, giving 2 to 15 km as the
overall range of widths (Figures 2, 3, and 7). The shallow (at
about 1-km depth) conductors are even less well constrained
laterally because the recording frequency is higher. Thus, the
shallow conductive inferred fault zones are shown in Figure 6
with narrow and uniform widths. The interpreted electrical
property sections also show regional changes in the resistive
crust from south to north. Most of the subsurface in the upper
20 km beneath Reese River Valley and southern Boulder Valley
(Figure 2) are underlain by rock that is generally more
conductive than the subsurface beneath Pumpernickel Valley and
northern Boulder Valley (Figure 3). This suggests that either
elevated-temperature or high-salinity fluids, alteration, or
carbonaceous rocks are more pervasive in the more conductive area
(Battle Mountain Heat-Flow High, Figure 2). This implies that
the crust beneath these valleys is either more fractured or has
more carbonaceous rocks than in the area surveyed along the 41st
parallel.

SUMMARY

Strengthening the suggestion of Shawe (1991), MT data
collected along five profiles show deep resistivity structures
beneath the Battle Mountain-Eureka (BME) and Carlin gold trends
in north-central Nevada. These appear consistent with tectonic
breaks in the crust that possibly served as channels for
hydrothermal fluids. It seems likely that gold deposits along
these linear trends were, therefore, controlled by deep regional
crustal fault systems.

Two-dimensional resistivity modeling of the magnetotelluric
data generally show resistive (30 to 1,000 ohm-m) crustal blocks
broken by sub-vertical, two-dimensional, conductive (1 to 10 ohm-
m) zones that are indicative of large-scale crustal fault zones.
These inferred fault zones are regional in scale, trend
northeast-southwest, north-south, and northwest-southeast (Figure
7), and extend to mid-crustal (20 km) depths. The individual
conductors are about 2- to 15-km wide, extend from about 1 to 4
km below the surface to about 20 km depth, and show two-
dimensional electrical structure. By connecting the locations of
similar trending (Figure 4) conductors together, individual
regional crustal fault zones within the upper crust can be
inferred that range from about 4- to 10-km wide and about 30- to

150-km long (Figure 7). One of these crustal fault zones coincides with the Battle Mountain-Eureka mineral trend.

   The interpreted structure from these profiles provide constraints for future geologic interpretations of the genesis of gold deposits along the Carlin and BME trends. The deep regional resistivity structure across the Carlin and BME trends has direct implications for large gold resources in northeastern Nevada. The resistivity structure reveals major crustal structures that are probably related to the deformation and shattering of upper crustal rocks that provided local permeable zones favorable for fluid flow and precipitation of gold ores. This implies that the potential for additional undiscovered deeper-level deposits in the Carlin trend is large.

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Figure 1.   Index map. Magnetotelluric (MT) transects (MT5 to MT9) acquired
in 1998, 1999, and 2000. Hatched zones represent two northwest-trending
mineralized belts in north-central Nevada, the well-known Carlin trend and the
less well-defined Battle Mountain-Eureka trend. The MT stations are shown as
gray circles (for data that is 2-D in character) and triangles (for data that
is three-dimensional character in the lower frequencies). KGRA is Beowawe
Known Geothermal Resource Area. The base map is adapted from Struhsacker and
others (1996).

Figure 2. Two-dimensional, resistivity models at 10 times and no vertical exaggeration of MT profiles MT4, MT6, MT7 and MT8. Depths are from ground surface. White arrows mark interpreted crustal fault zones. Black squares show stations whose MT data indicated three-dimensional character at the lower frequencies. Dashed white lines in the upper few kilometers represent estimated basin basement. Dashed black lines outline estimated crustal fault zone (CFZ) boundaries. Question marks indicate less certain definition of fault zone boundaries or basin basement. NNR is Northern Nevada Rift.

Figure 3. Two-dimensional, resistivity models at 10 times and no vertical exaggeration of MT profiles MT5 and MT9. Depths are from ground surface. White arrows mark interpreted crustal fault zones. Black squares show stations whose MT data indicated three-dimensional character at the lower frequencies. Dashed white lines in the upper few kilometers represent estimated basin basement. Dashed black lines outline estimated crustal fault zone (CFZ) boundaries. Question marks indicate less certain definition of fault zone boundaries or basin basement. NNR is Northern Nevada Rift.

Figure 4.   Deep conductor strike directions for magnetotelluric (MT)
transects MT1 to MT9 acquired from 1994 to 2000.   Hatched zones represent two
northwest-trending mineralized belts in north-central Nevada, the well-known
Carlin trend and the less well-defined Battle Mountain-Eureka trend.   The MT
stations are shown as gray circles (for data that is 2-D in character) and
triangles (for data that is three-dimensional character in the lower
frequencies).   Bow-tie symbols show estimated 2-D strike directions for 2-D
conductors at about 10-km below the ground surface.   MT strike resolution is
about ±20 degrees.   The base map is adapted from Struhsacker and others
(1996).

Figure 5.   Shallow conductor strike directions for magnetotelluric (MT) transects MT1 to MT9 acquired from 1994 to 2000.   Hatched zones represent two northwest-trending mineralized belts in north-central Nevada, the well-known Carlin trend and the less well-defined Battle Mountain-Eureka trend.   The MT stations are shown as gray circles (for data that is 2-D in character) and triangles (for data that is three-dimensional character in the lower frequencies).   Bow-tie symbols show estimated 2-D strike directions for 2-D conductors at about 1-km below the ground surface.   MT strike resolution is about ±20 degrees.   The base map is adapted from Struhsacker and others (1996).

Figure 6.   Shallow fault zones defined by magnetotelluric (MT) transects MT1 to MT9 acquired from 1994 to 2000.   Hatched zones represent two northwest-trending mineralized belts in north-central Nevada, the well-known Carlin trend and the less well-defined Battle Mountain-Eureka trend.   The MT stations are shown as gray circles (for data that is 2-D in character) and triangles (for data that is three-dimensional character in the lower frequencies).   Interpreted fault zones, shown by black lines are based on the locations of the shallow sub-vertical 2-D conductors at about 1-km below the ground surface.   The base map is adapted from Struhsacker and others (1996).

Figure 7.   Deep crustal fault zones defined by magnetotelluric (MT) transects MT1 to MT9 acquired from 1994 to 2000.   Hatched zones represent two northwest-trending mineralized belts in north-central Nevada, the well-known Carlin trend and the less well-defined Battle Mountain-Eureka trend.   The MT stations are shown as gray circles (for data that is 2-D in character) and triangles (for data that is three-dimensional character in the lower frequencies). Interpreted crustal fault zones, shown by 4- to 10-km wide gray swaths are based on the locations and lateral extents of the deep sub-vertical 2-D conductors (see Figures 2 and 3.   Dashed black line is outline of northern Nevada rift (Zoback and others, 1994).   The base map is adapted from Struhsacker and others (1996).

APPENDIX A

OBSERVED AND CALCULATED DATA – PROFILE MT5

Magnetotelluric (MT) observed (circle and x symbols) and calculated (solid and dashed lines are TE and TM modes, respectively) resistivity and phase data for profile MT5. See the "Magnetotelluric Data" section of this report for a description of the observed data and the "Resistivity Models" section for a description of the calculated data.

APPENDIX B

OBSERVED AND CALCULATED DATA – PROFILE MT6

Magnetotelluric (MT) observed (circle and x symbols) and calculated (solid and dashed lines are TE and TM modes, respectively) resistivity and phase data for profile MT6. See the "Magnetotelluric Data" section of this report for a description of the observed data and the "Resistivity Models" section for a description of the calculated data.

APPENDIX C

OBSERVED AND CALCULATED DATA – PROFILE MT7

Magnetotelluric (MT) observed (circle and x symbols) and calculated (solid and dashed lines are TE and TM modes, respectively) resistivity and phase data for profile MT7. See the "Magnetotelluric Data" section of this report for a description of the observed data and the "Resistivity Models" section for a description of the calculated data.

APPENDIX D

OBSERVED AND CALCULATED DATA – PROFILE MT8

Magnetotelluric (MT) observed (circle and x symbols) and calculated (solid and dashed lines are TE and TM modes, respectively) resistivity and phase data for profile MT8. See the "Magnetotelluric Data" section of this report for a description of the observed data and the "Resistivity Models" section for a description of the calculated data.

APPENDIX E

OBSERVED AND CALCULATED DATA – PROFILE MT9

　　Magnetotelluric (MT) observed (circle and x symbols) and calculated (solid and dashed lines are TE and TM modes, respectively) resistivity and phase data for profile MT9. See the "Magnetotelluric Data" section of this report for a description of the observed data and the "Resistivity Models" section for a description of the calculated data.