



Gravity Study of the Guernsey Landfill Site, Guernsey, Wyoming

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Open-File Report 2004-1383

U.S. DEPARTMENT OF THE INTERIOR
U.S. GEOLOGICAL SURVEY

Introduction

A gravity survey was conducted near the Camp Guernsey cantonment area and the town of Guernsey, Wyoming, to characterize the bedrock topography and surficial deposits. The survey was performed by personnel from the United States Geological Survey (USGS) Geologic Discipline in cooperation with the USGS Water Resources Discipline, the Wyoming Department of Military Affairs, and the town of Guernsey. From July 9 to July 19, 2003, fifty-two stations were established in the inactive landfill north of town, and eleven stations were established across the North Platte River valley through the town of Guernsey and west of the cantonment area.

The landfill is situated on the southern flank of the Hartville uplift (Harris, 1997) on a river terrace approximately 16 meters above the North Platte River valley (fig. 1 and 2). Dense metadolomite is quarried immediately east of the landfill. The metadolomite and overlying metabasalt crop out in the area of the quarry. In the area of the landfill at the bottom of the deepest ravine is an outcrop of limestone interpreted to be a member of the Hartville formation that overlies the dolomite and basalt. The remainder of the area is covered with terrace gravels.

An overall accuracy in the range of tens of μGal ($1 \mu\text{Gal} = 10^{-8} \text{ meter/second}^2$) and station spacing on the order of tens of meters were considered appropriate to characterize the basement and enable stripping of the alluvial signature. A digital gravimeter and real-time kinematic (RTK) Global Positioning System (GPS) were used to gather data, and some comments are noted for their use in this setting. In addition, vertical datum conversion and gravity anomaly computation are described in relation to emerging standards.

In addition to the gravity survey, a direct-current (DC) resistivity survey (McDougal and others, 2004) was conducted to characterize the water table and the stratigraphic section.

Gravity data

A LaCoste-Romberg (L&R) gravimeter with Aliod100 electrostatic beam nulling and digital readout was used to measure the gravity field. Data were recorded on a Palm Pilot connected by a serial cable to the Aliod electronics box attached to the outside of the gravimeter. The gravimeter sensor is in the standard L&R thermostatically controlled enclosure, and when the Aliod is in operation, pendulum levels inside the enclosure drive meters the operator uses to level the instrument instead of the standard spirit levels outside the enclosure. Although all readings in this survey were obtained with the electronic nulling, the Aliod electronics may be turned off and the meter leveled via the spirit bubbles and read visually through a microscopic eyepiece.

Each station requires about 15 to 20 minutes to obtain a reading, including setup and tear-down. The digital output of the meter stabilizes within 2 minutes of unclamping the beam and then slowly drifts about $10 \mu\text{Gal}$ to the final reading. To avoid the limitations

of reading for a short period at a constant time after unclamping, the meter was read at 15-second intervals starting 6 minutes after unclamping and continuing for 3 to 4 minutes. If the readings were consistent and stable, then the last 1 or 2 minutes of readings were averaged to produce a single reading. The digital output has a resolution of 1 μGal , and 0.1 μGal was used for data reduction. Laboratory tests with this gravimeter indicate repeat measurements on the order of $\pm 2 \mu\text{Gal}$ are common with the electrostatic nulling system, which is about 3 times better than measurements obtained using optical and manual beam nulling.

Base stations were read at 1- to 1.5-hour intervals which allowed, at most, three to four station readings in each base loop. Up to six base station loops were collected each day. Most of the stations in the landfill area were read at least twice, and when the difference exceeded about 10 μGal , the stations were reread. Two base stations (g01 and g21, fig. 3) were used in the landfill area, and the seven repeat readings of the second base with respect to the first yield a standard deviation of 5 μGal . A single tie to the International Gravity Standardization Net 1971 (IGSN71) base station (Morelli, 1974) in Wheatland was done. The principal facts for the survey appear in Appendix A.

The long-term drift rate of the meter operated in the laboratory is in the range of 1 to 2 μGal per hour. In the field, the drift rate is substantially higher with 10 to 20 μGal per hour not uncommon, and several loops had up to 30 μGal per hour. Some of the higher drift rate is from small tares caused by normal handling, but some is caused by abrupt changes in temperature. To compensate, the meter was equilibrated to the ambient temperature for at least 1 hour before starting measurements and was exposed to the ambient temperature, either in the shade of the operator or outside the truck, whenever possible. On several occasions, the digital readout would not stabilize, and the meter would require re-equilibration in a location with reduced radiant heating.

Appendix B contains a summary of statistics from repeated stations. Overall, the range of the relative gravity readings at a station did not exceed 37.3 μGal with an average range of 10.5 μGal . As mentioned previously, tares are suspected for some of these larger variations, and examination of the entire data set shows almost all the data collected on July 12th and 13th contribute an extreme value to the range of station values. When these days are removed from the data along with four other extremes, the maximum range of the gravity readings drops to 14.4 μGal with an average range of 5.9 μGal . The poor repeat values for the 12th and 13th were noticed in the field, and the procedure was changed to ensure the meter and base plate were always shaded. A persistent light wind helped keep the meter at ambient temperature even in the presence of heat reflected from the ground as the temperature rose to 106 degrees Fahrenheit later in the survey.

Position surveying

Stations were located with a Javad Legacy-E dual frequency, phase comparing, real-time differential Global Positioning System (GPS) unit. A benchmark southeast of town, National Geodetic Survey (NGS) id-number NQ0176 and stamped "Z 63" (Appendix C),

is the base for the coordinates of the gravity survey. This benchmark is a Federal Base Network Control Station (<http://www.ngs.noaa.gov/>) and has up-to-date GPS derived horizontal and vertical coordinates in the North American Datum of 1983 (NAD83). The specific horizontal datum is denoted as NAD83(93) in Appendix C. A local Javad base near gravity station g21 was established and was used for all subsequent surveying. Two other benchmarks in the vicinity were occupied as a check on the overall position surveying. These benchmarks are listed in Appendixes A and B as gravity stations bh2 and bh4, and the surveyed elevation, rounded to the nearest foot, agrees with the elevation listed on the published USGS topographic map.

The Javad was operated in static mode with the antenna mounted on a range rod while surveying station points. It provides horizontal coordinates with an error of about 17 millimeters and a slightly greater vertical error with an occupation time of about 1 minute. Topographic mapping was performed with the unit mounted in a backpack and the user walking at a normal pace. The sampling rate was one per second.

Most of the landfill was surveyed in kinematic mode, and the coverage may be seen in figure 3 as the color contour background. About 9 hours, plus in-field upload and processing time, were required to collect the topographic data. Coverage is more dense near gravity stations to allow good definition for terrain correction and less dense away from the gravity stations, where the relief is low. The surveyed area combined with the USGS 10-meter digital elevation model (DEM), shifted to NAD83 horizontal and vertical coordinates, is shown in perspective view in figure 2 gridded at a 1 meter interval.

Gravity stations were located on small cleared areas with vegetation and rocks removed. Surveyed locations are at ground level in the center of these cleared areas. The ground clearance of the gravity meter base plate was noted for each reading, but the plate location may be summarized as either pressed into the ground for stability or sitting above ground level on the points of the plate tripod.

Vertical datum conversion

With the advent of high accuracy GPS surveying and geoid models, new gravity surveys may use either orthometric elevations based on the geoid or height above the reference ellipsoid. The geoid is the presumed location of sea level in a location, however, that location or surface varies according to large-scale gravity variations, and referencing gravity calculations on it causes what is known as the indirect effect (Chapman and Bodine, 1979 or Li and Götze, 2001).

Emerging gravity standards likely will favor the International Terrestrial Reference Frame (ITRF) coordinate system as the single global reference, but currently those standards can be met only if there is a benchmark with updated coordinates nearby or if the survey party can occupy a point with a dual frequency GPS for several hours over multiple days. Documentation of the details of station altitude will allow this and other new gravity surveys to be merged with older surveys, which used the NGVD29 in the

conterminous United States. The relations among the various vertical coordinates can be summarized in this area with a set of algebraic shifts.

$$\text{NAVD88} = \text{VERTCON} + \text{NGVD29}$$

$$\text{NAD83} = \text{GEOID99} + \text{NAVD88}$$

explanation:

NGVD29 National Geodetic Vertical Datum of 1929, approximately orthometric elevations (height above geoid). The datum commonly found on all but the latest USGS topographic maps.

VERTCON A grid of elevation shifts to convert the NGVD29 orthometric elevations to a more consistent orthometric elevation (available at <http://www.ngs.noaa.gov>).

NAVD88 North American Vertical Datum of 1988, orthometric elevations (height above the geoid, the presumed sea level that would occur in the area).

GEOID99 A grid of geoid heights above the GRS80 ellipse in the NAD83 vertical datum (Smith and Roman, 2001).

NAD83 North American Datum of 1983 height above the GRS80 ellipsoid, but with an approximately 1 to 2 meter origin shift that does not make the ellipsoid geocentric. This is a specific datum in reference to the ITRF. The ITRF varies slightly in time due to improved technique and continental drift and is the basis for WGS84 coordinates output by GPS units.

WGS84 World Geodetic System of 1984 is the geocentric system of coordinates used by the Global Positioning System. WGS84 epoch G1150 is equivalent to ITRF00, and both are current as of mid 2003.

In this area $\text{VERTCON} = 0.78$ meter and $\text{GEOID99} = -14.95$ meter. It should be noted that the NAD83 and WGS84 height above the ellipsoid is not the same, and in this vicinity, the difference is about 0.8 meter.

Regardless of the above, the USGS 10-meter DEM generated from the NGVD29 contours of the original 1955 topographic map required a locally determined shift of -16.95 meters to more closely match the Javad surveying. This shift was determined by the average difference between the two data sets in undisturbed low-relief areas outside the landfill. Since the VERTCON and GEOID99 shifts sum to -14.17 and -16.95 was applied to the 10-meter DEM, there is a difference of 2.78 meters between the surveyed data and the DEM. The GPS surveyed altitudes agree with the mapped benchmark altitudes rounded to the nearest foot and the 20 foot (6.1 meter) contoured DEM visually agrees with the contours of the topographic map. The combined elevation model is used for terrain correction to a radius of 2.615 kilometers from each gravity station.

Gravity data reduction

In the field, meter readings were converted to relative gravity in mGal ($1 \text{ mGal} = 10^{-5} \text{ m/s}^2$) and tide corrected (Longman, 1959). Linear drift correction was applied by using station g01 in the northwest corner of the landfill as the primary base or station g21

as a secondary base on the south side of the area. As the survey progressed, the tie between g01 and g21 was reread until the distribution of the error could be quantized. As previously mentioned, the error has a standard deviation of 5 μGal .

A single loop to connect g01 to the IGSN71 base station in Wheatland was performed. The Wheatland gravity benchmark was not recovered, although a stamped description was found on a supporting member of the beacon tower under which the benchmark should have been located. The length of the loop was 3 hours, with a modest -15 μGal per hour drift, so the tie should be well within the stated accuracy of 100 μGal , including lack of altitude recovery. The absolute gravity at the Wheatland base is given as 979,913.79 mGal (Morelli, 1974).

Terrain correction was computed from the station to a radius of 166.7 kilometers in two separate computations, with the division at 2.615 km. The inner portion ends at Hammer zone H (Hammer, 1939), and the outer portion ends at Hayford-Bowie zone O (Hayford and Bowie, 1912). The outer correction was computed with computer program OUTERTC (Plouff, 1977). The inner zone was computed with program DCTC (Webring, unpublished program). The Plouff algorithm uses elevation models of .25, 1, and 3 arc-minute resolution and a spherical Earth. The DCTC (Digitized Contour Terrain Correction) program creates the elevation model on the fly by using random surveyed points and, in this instance, the USGS 10-meter DEM, converted to XYZ, for the area immediately around the landfill. DCTC starts with a grid of 192-meter interval for zone H, interpolated with minimum curvature (cf. Webring, 1981) enhanced with a tension factor (Smith and Wessel, 1990), and then subdivides and reconverges the grid for successively closer zones ending with one meter at zone A.

The outer zone correction varies from -47 to -63 μGal at 2.67 gm/cc, while the inner correction varies from 98 to 337 μGal . Negative outer zone terrain corrections can occur in areas of low relief. The largest inner correction occurs in a section of profile that descends 16 meters from the alluvial terrace to the level of the Platte River Valley. Numerical tests with this section of profile indicate horizontal movement of the station by 0.1 meter varies the inner zone correction less than 1 μGal while vertical movement of the station by 0.01 meter varies the correction by about 5 μGal .

The ellipsoidal Free Air and Bouguer anomalies were computed following recommendations of the North American Gravity Database Standards Committee (oral commun.). As defined by the geodetic community, the gravity anomaly is referenced to the geoid, and the gravity disturbance is referenced to the ellipsoid (Heiskanen and Moritz, 1969); however, the emerging terminology is to prefix the adjective “ellipsoidal” to anomaly calculations based on ellipsoidal heights. The ellipsoidal Free Air anomaly was computed using a theoretical gravity and atmospheric correction (Moritz, 1980) and second-order height correction (Heiskanen and Moritz, 1969). Each is referenced to the GRS 1980 ellipsoid. The ellipsoidal Bouguer anomaly was computed using a spherical terrain cap (LaFehr, 1991) and a terrain correction on an Earth with a radius of 6,371 kilometers. The terrain model (spherical cap minus terrain correction) for the Bouguer anomaly was computed using a density of 2.67 grams per cubic centimeter (gm/cc).

The station altitudes and topographic data were surveyed in the NAD83 vertical datum rather than the recommended ITRF vertical datum (the one used by GPS) because the software for that conversion was not available at the time the data were reduced. However, the NAD83/ITRF altitude shift varies smoothly from an amplitude of about 0.3 meter to about 1.6 meter over the conterminous United States, so the shift may be approximated as a constant in a local area. The computed anomalies for this survey therefore have a small constant offset that may be properly ignored.

Bouguer contours and station locations for the landfill area and the profile south across the North Platte River valley are shown on figure 4. Curvature of the contours at the edge of the data, especially at the south end of the North Platte River valley profile, are artifacts of the gridding process. No prior smoothing was done on any figure, however, the display program seems to have a spline-based interpolation operation in place.

Regional gravity field

The existing gravity stations for about a 100 kilometer radius around the study area were reduced to the Bouguer anomaly and the resulting regional field lowpassed with a 100 kilometer filter. This regional gravity field can be interpreted as crustal thickness variations and the effects of distant sedimentary basins. In the area of this survey, the regional gravity field is a plane dipping southeast with a 1 mGal difference in amplitude from north to south. The regional gravity field was subtracted from the Bouguer anomaly to produce the residual Bouguer anomaly shown on figure 5. A strong north-south gradient remains that indicates the core of the Hartville uplift to the north. In retrospect, this survey did not go far enough north to define the gravity response of the shallow versus the deeper, larger scale structure of the Hartville uplift.

Density of hand samples

Several hand samples were collected and measured for density to aid the interpretation. Two samples of each rock type were used:

- Metadolomite from the quarry: 2.82 gm/cc.
- Metabasalt from hill north of the quarry: 2.64 gm/cc.
- Recrystallized limestone from ravine outcrop in landfill area: 2.66 gm/cc.
- Gray friable sandstone from outcrop south of town, tentatively identified as Arikaree formation (Harris, 1997): dry 1.43 gm/cc, saturated 1.83 gm/cc.

Density of the surficial terrace gravel

The terrain corrections listed in Appendix A are computed at a standard 2.67 gm/cc (Hinze, 2003) that, when subtracted from a constant thickness spherical cap, result in a gravity model from the topographic surface down to the reference ellipsoid and radially out from the station to 166.7 kilometers. Adding this terrain model effect to the gravity predicted in the Free-Air anomaly calculation results in a terrain-free gravity known as

the Bouguer anomaly. If the terrain has a consistent density of 2.67 gm/cc, as many mountainous areas of granitic composition have, then all topographic effects are removed, and only anomalous subsurface variations remain.

The recrystallized limestone found in the ravine as well as the metabasalt have densities very close to the standard 2.67 gm/cc and are accounted for with the standard anomaly calculation. The metadolomite is more dense and can cause gravity anomalies if it varies from a uniformly horizontal slab. The overlying terrace gravels found at the surface in the landfill will have a lower than standard density, because there is a high percentage of pore space between the uncemented grains, and the irregular topographic surface will, therefore, cause an anomaly. Other rock formations with nonstandard densities are far enough distant to affect all stations equally.

The gravity of the topography close to the station can be computed in a fashion similar to the standard Bouguer topographic model. A terrace model was constructed that extends 170 meters from the stations (Hammer zone D) and down to an altitude of 1,323 meters, just below the elevation of station g42, the lowest landfill area station. The model is a flat-bottomed cylinder with the topography as the top surface of the cylinder, and the density used in the calculation is relative to 2.67 gm/cc. Separate model computation was centered on each station. The results for profiles 1, 2, and 3 are shown on figures 6, 7, and 8 as distance versus station altitude (red) in the bottom panel and Bouguer anomaly (green) in the top panel. The several curves located above the Bouguer anomaly curve are the result of terrace gravel models at different densities added to the Bouguer anomaly.

Nettleton (1939) first proposed the examination of density with relation to topography, and the three profile figures show the results for terrace densities of 1.4, 1.6, 1.8, 2.0, and 2.2 gm/cc. Examination of the south end of profile 2 on figure 7 reveals a strong topographic effect caused by using a Bouguer reduction density of 2.67 gm/cc. The Bouguer anomaly curve in green shows an abrupt change in curvature and smoothness from stations g24 to g30, which are located on the greatest slope where the terrain correction is highest (station g26 has the highest inner zone correction of 337 μ Gal). The south end of profile 2 indicates the density of the terrace gravel is somewhat less than 2.67 gm/cc, since all the model curves are smoother than the Bouguer curve. However, the 2.5-D model shown later indicates a dense basement mass with top-surface relief (for instance, not flat topped) inside the topographic mass outlined by the altitude curve on figure 7. Therefore, the overall curvature and derived density from stations g24 to g29 is driven by a combination of terrace gravel and basement rock density .

The specific curvature referred to in the following is an estimate of the second horizontal derivative centered on the i^{th} point. The stations are assumed to be equally spaced in distance, and an estimate of the curvature may be computed with the relation:
$$c_i = z_{i-1} - 2 z_i + z_{i+1}$$
, where z is any measured quantity with relation to distance. When the three z values are collinear, the curvature is zero, and the quantity in question has no correlation with another quantity that has a nonzero curvature.

Another location to apply the Nettleton method is in the center of profile 1, enlarged on figure 9. Stations g8, g4, and g11 are located at slope breaks as seen in the bottom panel of figure 9, and the green Bouguer anomaly curve is inversely correlated with the station altitude. The terrace model curves above the Bouguer curve become straight (for instance, no correlation with topography) between densities 2.2 and 1.4 gm/cc. Specifically, station g8 at density 1.7 gm/cc, station g4 at density 2.1 gm/cc, and station g11 at density 1.8 gm/cc, which average to 1.87 gm/cc. Tests on the profile 1 stations with more localized models (Hammer zone B at 16.6 meter radius and shorter cylinders) indicate densities of 1.6, 2.0, and 1.8 gm/cc respectively, which average to 1.8 gm/cc. There are three candidates for density determination in profile 2 using the 16.6 meter model; stations g23 with 1.6 gm/cc, g25 with 2.0 gm/cc, and g27 with 1.6 gm/cc. Station g27 is located in a section of the profile with consistently negative curvature, and the density derived at this location likely is driven by the basement rock previously mentioned.

In summary, the terrace gravel bulk density derived by the Nettleton technique is about 1.8 gm/cc. There is a large variation of the densities included in this determination, and too few samples for meaningful statistics. To give an idea of the data accuracies required, the terrain corrections to a radius of 16.6 meters for the stations used in profile 1 (stations g4 and g7 through g12) vary from a low of 0.8 μ Gal to a high of 10.8 μ Gal with an average of 5.9 μ Gal, but given the gentle topography at the station locations the computed effect probably is accurate to better than 1 μ Gal. The observed gravity for these stations have an average range of 6.7 μ Gal (Appendix B) and a standard deviation of ± 1.9 μ Gal. The terrain correction, which drives the density determination via the gravity effect of the cylindrical model, is not very large in relation to the probable error in the observed gravity.

The final selection of the appropriate density for the terrace gravel model was via inspection of the landfill area terrace residual gravity in contour form for densities 1.50 gm/cc to 2.10 gm/cc at intervals of 0.1 gm/cc. In general, the contours of figure 5 straightened until a reversal of overall curvature was noted at the lower densities. The density at which the contours contained the least information (the smoothest and straightest) is 1.8 gm/cc, and the resulting gravity contours can be seen on figure 10.

Local trend removal from the terrace residual gravity

The terrace residual gravity on figure 10 is the result of removing the overcorrection by the standard terrain model, but while the obvious dependence on local topography has been removed, the north-south trend of the regional residual gravity remains. Examination of the contour spacing on figure 4 reveals no obvious slope breaks near the landfill upon which to base an estimate of the deeper structures of the Hartville uplift. The simplest structural model that produces a planar gravity field is a plane but one with any of a wide range of density contrast and slope. Therefore, a plane with a slope of 6.0 mGal/km north was removed from the terrace residual gravity with the understanding there is no particular deep structural information implied.

The de-trending plane was fit in the north-south direction by observation of the average gravity values near stations g06 and g42, and when it is subtracted from the terrace residual gravity, the near-surface gravity anomalies remain (fig. 11). The specific value of slope and direction is somewhat arbitrary, and small variations produce similar results. Secondary east-west slopes on the order of 0.5 mGal/km were examined and discarded for lack of constraints.

De-trended terrace residual gravity interpretation

The de-trended gravity was produced by removal of known gravity effects, including deep-seated structures and the effect of the most shallow source (the topographic surface). The most striking of the resulting anomalies is an elongate low that is centered to the northwest of the mapped landfill cell along profile 1 and an elongate high that starts in the southwest corner at station g43, is most prominent up on the terrace near station g29, extends down into the ravine at station g20, and continues up onto the terrace at station g40. The relative difference in gravity between the low over profile 1 and the high on the terrace to the south remains a fairly constant 1 mGal for a variety of terrace gravity densities and de-trending planes. In addition, the low along profile 1 remains north of station g52 (the center of the landfill cell) for all tested de-trending planes.

Station g52 is located in the center of the largest landfill cell (outlined on figure 11) and, in that position, is sensitive to a change of gravity caused by the addition of waste material and the disturbance of the soil. As can be seen on figure 11, the low gravity value in the area is near the intersection of profiles 1 and 2, and there is no distortion of the low that tends to follow the outline of the cell. A cylindrical model of the landfill cell 50 meters in radius, 8 meters thick and a relatively high density contrast of -0.2 gm/cc with respect to the terrace gravel generates a gravity effect of only -62 μ Gal, far less than that necessary to explain the 1 mGal anomaly. In addition, profile 3 crosses several minor ravines as determined from examination of the pre-landfill topographic map, but the gravity shows no indication of variation caused by density changes. From this, it may be supposed that compacted native fill dirt with a minor amount of landfill debris has about the same density as the undisturbed soil.

The data now have two density datums: the local 1.80 gm/cc layer above altitude 1,323 meters (NAD83 vertical) caused by the addition of the 3D terrace model, continuing with the standard 2.67 gm/cc model down to the ellipsoid. Further structural modeling across the 1323 meter boundary is most easily performed using a single density datum so a model body has a single density contrast. Provided the bottom of the terrace model is level and uninterrupted to a distance sufficient to minimize edge effects, approximately five times the model depth, the density reference below 1323 meters can be changed to 1.80 gm/cc and the resulting constant change in the gravity added to the regional field.

A 2.5D model was constructed with SAKI (Webring, 1985) along profile 2 and can be seen on figure 12. The top panel has the de-trended gravity values in red and the computed response of the model as a black line. The structural model in the bottom panel plots the gravity station locations and two layers or bodies without vertical exaggeration.

The density reference for the model is 1.80 gm/cc so the terrace gravels have a density contrast of 0.0 gm/cc and contribute no effect and are, therefore, not enclosed in a body.

The thin area between the two undulating lines is body 1 and is interpreted to be a combination of Hartville limestone and metabasalt with a density contrast of 0.87 gm/cc (2.67 gm/cc real density). Below this layer is body 2 that is interpreted as metadolomite with a density contrast of 1.02 gm/cc (2.82 gm/cc real density). The bottom of body 2 extends below the area of the panel to an arbitrary constant depth, and because it is level, does not contribute to the lateral variation of the gravity response. Both bodies extend to the south and north about 1 kilometer to minimize edge effects. The bodies extend 100 meters at 90 degrees from the strike of the profile. The calculated gravity effect floats to minimize the difference with respect to the observed gravity to compensate for deep-seated mass that is not being modeled.

The Hartville formation is included in the model, because it outcrops near station g20 at distance coordinate 80,440 meters and is modeled as continuing west of the outcrop under profile 2 at approximately the same altitude as the outcrop. The Hartville layer, body 1, is modeled as being continuous with a thickness of about 10 meters. This is an assumption that is not constrained by the gravity data, and the layer is shown as continuous based on the observation that there are no major discontinuities in the gravity field. The metadolomite layer, body 2, can be used alone to fit the gravity field, in which case it would occupy some fraction of the thickness now shown as Hartville.

Once the basic model on figure 12 was fit to the de-trended terrace residual gravity, another possible constraint became apparent. The first model had the Hartville at a depth of approximately 30 meters below the surface at station g06 on the north end near coordinate 81,000 meters. The model was brought to about 15 meters below the surface to match the depth at station g20. The trend of the gravity then was recomputed to fit the model, and the end result is the 6.0 mGal/km de-trending surface could be reduced to 4.9 mGal/km. The model was reconverged, and the final result is shown on figure 12. In other words, the original assumption that the entire trend was deep seated could be modified by assigning some of the trend to the top surface of the basement formations.

The relief in the hard-rock layers between distance coordinates 80,440 and 80,700 meters is about 30 meters and, as has been mentioned is not too dependent on the specific composition of the layer or the details of de-trending. For example, the initial model fit to the 6.0 mGal/km de-trend data had the same overall basement relief as the reconverged model using 4.9 mGal/km de-trend data.

The model described in the preceding sections is relatively simple, and the interpretation must likewise be hedged with regard to the possible features that are simplified. In particular, the east-west extent of these features is not well constrained. The north-south direction has more stations but, in that direction, is the major gradient which does not have any major slope breaks to help define the deep-seated causative bodies. The outcrop near station g20 is assumed to be in place, which then fixes the altitude of the

basement model at one point. The assumption that the basement remains shallow to the north helps define the range of probable trend surfaces.

Conclusions

The Bouguer anomaly on figure 4 shows a south to north increase of about 9 mGal, increasing in slope to the north toward the presumed core of the Hartville Uplift. Stations to the east and west of the cross-valley profile indicate the slope of the gravity field is almost due north-south. The interpreted dip of the basement and(or) thickening of the North Platte River sediments then would be approximately due south.

The residual Bouguer anomaly around the landfill (fig. 5) shows a high-low pair defined by the north-south profile 2, the southeast portion of profile 3 and the outlying perimeter stations. While station g20 (southeast of the junction of profiles 1 and 3) appears to be causing a one-station anomaly, the overall trend of profile 3 tends to confirm that there is a local high-density structure to the southeast of the landfill cells outlined in red.

The gravity anomaly values of the southwest to northeast profile 3 vary smoothly. This profile crosses several minor ravines, as determined from examination of the old topographic surface, but is now a smooth dirt road. From this, it may be supposed that compacted landfill debris combined with fill dirt has about the same density as the undisturbed terrace gravels. In addition, a simple cylindrical model of the estimated dimensions of the cell beneath station g52 would only generate a fraction of the effect seen between profile 1 and the south end of profile 2.

Modeling of the terrace gravels using a density of 1.80 gm/cc removes some obvious terrain effects and leaves a predominantly northward increasing gravity field. A planar fit to this trend reveals anomalies with strong east-west elongation that are stable with respect to small variations of the plane. Further 2.5-D modeling of the terrace residual gravity indicates 30 meters of basement relief with the low area to the north of the landfill cells and the high area on the south boundary of the landfill.

Data availability

This report, the gravity and surveying data sets are available as separate files located on greenwood.cr.usgs.gov.

Figure credits

- Digital Line Graphics (DLG) file of roads on figure 1 were acquired from the USGS/EROS Data Center, Sioux Falls, S. Dakota.
- Digital Exchange Format (DXF) file of landfill features on figures 1, 2, 3, 4, 5, 10, 11 were acquired from the USGS/Water Resources Division, Cheyenne, Wyoming.
- Digital Elevation Model (DEM) files of topographic data on figures 1 and 2 were acquired from the USGS/EROS Data Center.

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Appendix A

Principal facts for the Guernsey landfill gravity survey:

station - gravity station ID
 longitude - NAD83 datum in decimal degrees
 latitude - NAD83 datum in decimal degrees
 elevation - NAD83 datum, ellipsoidal height in meters
 obsgrv - observed gravity relative to IGSN71
 itc - inner terrain correction (zones A-H) at 2.67 gm/cc
 otc - outer terrain correction (zones I-O) at 2.67 gm/cc
 faa - free air anomaly, referred to ellipsoidal height
 ba - Bouguer anomaly referred to ellipsoidal height, reduction density 2.67 gm/cc
 (proposed name: ellipsoidal Bouguer anomaly, NAGDB committee, 2003)

station	longitude dec degree	latitude dec degree	elevation meter	obsgrv mGal	itc mGal	otc mGal	faa mGal	ba mGal
bh1	-104.7466471	42.2694335	1310.089	979975.0075	0.1515	-0.024	6.6847	-141.1598
bh2	-104.7294881	42.2727569	1329.375	979972.9342	0.2021	-0.061	10.2580	-139.7423
bh3	-104.7179542	42.2600324	1316.303	979970.3679	0.0372	-0.069	4.8077	-143.8952
bh4	-104.7414126	42.2652887	1308.624	979973.4484	0.1159	-0.033	5.0473	-142.6771
bh5	-104.7414843	42.2607775	1305.713	979972.4056	0.3472	-0.032	3.5134	-143.6513
bh6	-104.7427084	42.2579626	1301.496	979972.6068	0.1309	-0.026	2.6680	-144.2326
bh7	-104.7417840	42.2538060	1305.101	979971.2414	0.0801	-0.039	2.7884	-144.5815
bh8	-104.7402831	42.2678780	1309.686	979974.3269	0.0584	-0.031	6.0200	-141.8793
bh9	-104.7382304	42.2699783	1312.306	979974.6819	0.0629	-0.034	6.9936	-141.1990
bh10	-104.7362703	42.2716287	1318.933	979974.2724	0.1520	-0.044	8.4785	-140.3803
bh11	-104.7340465	42.2745282	1338.761	979972.1292	0.1160	-0.065	12.1871	-138.9589
g01	-104.7422469	42.2787332	1350.343	979972.3490	0.1098	-0.063	15.5988	-136.8539
g02	-104.7410960	42.2780578	1347.285	979972.4588	0.1087	-0.063	14.8267	-137.2832
g03	-104.7379229	42.2793510	1349.371	979973.3697	0.1201	-0.062	16.2642	-136.0679
g04	-104.7401129	42.2789204	1347.042	979973.2149	0.0983	-0.061	15.4302	-136.6608
g05	-104.7441790	42.2787577	1344.866	979973.5931	0.1196	-0.059	15.1522	-136.6708
g06	-104.7424152	42.2807802	1351.700	979973.9555	0.0879	-0.061	17.4393	-135.1859
g07	-104.7404479	42.2788425	1346.794	979973.2301	0.1023	-0.061	15.3759	-136.6831
g08	-104.7407943	42.2787618	1346.825	979973.1410	0.1011	-0.061	15.3037	-136.7601
g09	-104.7411476	42.2786875	1348.367	979972.6999	0.0947	-0.062	15.3446	-136.8999
g10	-104.7397649	42.2789245	1346.005	979973.4625	0.1103	-0.060	15.3577	-136.6036
g11	-104.7393566	42.2790429	1345.238	979973.7754	0.1116	-0.060	15.4235	-136.4503
g12	-104.7389987	42.2791162	1346.954	979973.5224	0.1054	-0.061	15.6929	-136.3810
g13	-104.7386465	42.2791760	1347.686	979973.4792	0.1108	-0.062	15.8700	-136.2819
g14	-104.7415104	42.2786706	1349.305	979972.5030	0.0997	-0.063	15.4384	-136.9076
g15	-104.7418683	42.2786253	1349.866	979972.3600	0.1067	-0.063	15.4725	-136.9297
g16	-104.7422308	42.2786135	1349.827	979972.3906	0.1091	-0.063	15.4921	-136.9032

g17	-104.7425978	42.2786111	1348.737	979972.6193	0.1087	-0.062	15.3850	-136.8872
g18	-104.7382857	42.2792282	1348.777	979973.3144	0.1112	-0.062	16.0368	-136.2373
g19	-104.7375622	42.2793894	1349.365	979973.4874	0.1238	-0.062	16.3766	-135.9511
g20	-104.7394102	42.2769132	1332.423	979975.6152	0.1907	-0.052	13.5043	-136.8412
g21	-104.7405707	42.2762724	1342.433	979972.9954	0.1823	-0.062	14.0283	-137.4614
g22	-104.7400425	42.2754573	1326.359	979975.6177	0.1469	-0.046	11.7684	-137.9328
g23	-104.7401053	42.2755914	1327.792	979975.5352	0.1751	-0.047	12.1156	-137.7196
g24	-104.7401566	42.2757141	1331.199	979974.9596	0.2240	-0.052	12.5793	-137.5951
g25	-104.7402164	42.2758278	1336.771	979973.8140	0.2803	-0.058	13.1413	-137.6095
g26	-104.7402640	42.2759346	1341.418	979972.8742	0.3366	-0.062	13.6246	-137.5966
g27	-104.7402873	42.2760254	1342.472	979972.7595	0.2907	-0.063	13.8266	-137.5600
g28	-104.7403886	42.2761841	1342.291	979972.9731	0.2124	-0.062	13.9701	-137.4734
g29	-104.7404685	42.2763061	1342.417	979973.0438	0.1843	-0.062	14.0687	-137.4171
g30	-104.7405576	42.2765695	1342.948	979973.0817	0.1565	-0.062	14.2466	-137.3268
g31	-104.7406484	42.2768297	1343.531	979973.0300	0.1358	-0.062	14.3512	-137.3084
g32	-104.7407187	42.2770832	1342.830	979973.2564	0.1137	-0.061	14.3386	-137.2633
g33	-104.7408461	42.2773425	1344.402	979972.9765	0.1120	-0.062	14.5200	-137.2614
g34	-104.7409388	42.2776004	1345.737	979972.7446	0.1118	-0.062	14.6764	-137.2552
g35	-104.7410317	42.2778610	1346.285	979972.6499	0.1101	-0.063	14.7272	-137.2688
g36	-104.7411106	42.2781223	1347.080	979972.5271	0.1053	-0.062	14.8260	-137.2632
g37	-104.7411936	42.2783811	1347.736	979972.5633	0.0979	-0.062	15.0411	-137.1293
g38	-104.7395122	42.2778285	1339.404	979974.3629	0.1101	-0.057	14.3218	-136.8944
g39	-104.7391435	42.2807764	1355.565	979973.4818	0.1021	-0.062	18.1575	-134.8891
g40	-104.7377617	42.2774699	1345.549	979973.1501	0.1136	-0.063	15.0357	-136.8740
g41	-104.7391682	42.2744566	1323.989	979975.1210	0.1266	-0.045	10.6312	-138.8228
g42	-104.7407200	42.2747922	1323.452	979975.8407	0.1031	-0.042	11.1551	-138.2590
g43	-104.7432873	42.2755470	1325.917	979975.6922	0.1316	-0.043	11.6985	-137.9653
g44	-104.7398008	42.2776660	1339.737	979974.1958	0.1099	-0.058	14.2720	-136.9829
g45	-104.7400875	42.2775000	1340.979	979973.8450	0.1105	-0.059	14.3190	-137.0759
g46	-104.7404157	42.2773916	1343.469	979973.2428	0.1116	-0.061	14.4942	-137.1816
g47	-104.7393539	42.2780719	1339.781	979974.3736	0.1159	-0.058	14.4268	-136.8270
g48	-104.7390873	42.2783175	1339.062	979974.6325	0.1357	-0.056	14.4419	-136.7092
g49	-104.7385917	42.2788490	1346.291	979973.4657	0.1177	-0.061	15.4559	-136.5312
g50	-104.7387593	42.2797355	1352.441	979973.0073	0.1126	-0.063	16.8136	-135.8722
g51	-104.7389856	42.2802896	1352.834	979973.5698	0.1065	-0.062	17.4474	-135.2877
g52	-104.7404677	42.2783780	1349.780	979972.1648	0.1207	-0.064	15.2730	-137.1064

Appendix B

The following is a list of stations, the number of readings for the station, and the range of gravity readings for the station. Some stations have ranges larger than the target of about 10 μGal and, where feasible, the extreme reading is deleted and a new range is given in the comment field.

Two days, July 12 and 13, consistently yielded either the maximum or minimum value, and so these days were deleted entirely. This procedure leaves two stations, g34 and g35, without a repeat but given the range was less than 19 μGal , the remaining value likely is accurate to within the desired limits. Examination of the data in profile form confirms this expectation.

The statistics for the revised data are then:

number of samples = 46
maximum range = 14.4 μGal
average range = 5.93 μGal
standard deviation = 4.11 μGal .

The overall accuracy of the relative gravity values can be stated as: 95.5% of the time (2 standard deviations) the range of repeated reading will fall between 0 and 14.15 μGal . A 95.5% confidence level is then about ± 7 μGal . Averaging multiple readings further decreases the error.

The tie between g01 and g21 (primary and secondary bases), without deletion of two July 12 readings, has a range of 11.9 μGal and a standard deviation of 4.7 μGal . Deleting the July 12 readings yields a range of 5.1 μGal and a standard deviation of 2.2 μGal .

station id	number of readings	range μGal	comments
bh1	3	12.3	Tied to g1 and Wheatland base.
bh2	1	na	Benchmark(BM) 4408 (southeast of landfill)
bh3	1	na	BM Z-63, reference mark for all position surveying.
bh4	2	3.8	BM 4340 (in Guernsey)
bh5	2	1.5	
bh6	2	0.2	
bh7	2	2.2	
bh8	1	na	
bh9	1	na	
bh10	1	na	
bh11	1	na	
g01	3	12.3	Primary base for landfill survey

g02	3	9.1	
g03	3	16.5	Remove minimum, new range 0.1
g04	2	3.2	
g05	3	3.7	
g06	4	22.5	Remove both high and low, new range 11.7
g07	3	11.1	
g08	2	2.2	
g09	2	6.4	
g10	3	10.3	
g11	3	9.8	
g12	2	3.8	
g13	3	15.3	Remove(Rm) max value, new range 3.1
g14	2	1.7	
g15	2	11.0	
g16	2	7.2	
g17	2	2.1	
g18	3	14.4	
g19	2	10.8	
g20	2	1.4	
g21	7	11.9	Secondary landfill base, removing two July 12 readings lowers range to 5.1
g22	3	16.0	Rm 7/12 max, new range 1.6
g23	2	2.2	Rm 7/12 min, no range
g24	3	10.4	Rm 7/12 max, new range 4.2
g25	2	3.0	Rm 7/12 max, no range
g26	3	2.7	Rm 7/12 min, no change in range
g27	4	37.3	Rm 7/12 min, and the max, new range 6.1
g28	3	16.7	Rm 7/12 min, new range 4.8
g29	3	13.2	Rm 7/12 min, new range 2.4
g30	3	2.8	Rm 7/13 max, new range 0.5
g31	3	16.0	Rm 7/13 max, new range 6.3
g32	3	17.3	Rm 7/13 max, new range 4.1
g33	3	18.8	Rm 7/13 max, new range 4.1
g34	2	18.6	Rm 7/13 min, no range
g35	2	15.5	Rm 7/13 min, no range
g36	2	8.6	
g37	2	9.0	
g38	2	0.4	
g39	1	na	
g40	1	na	
g41	1	na	
g42	1	na	
g43	1	na	
g44	2	11.0	
g45	2	2.2	
g46	2	8.0	

g47	2	12.9	
g48	3	26.5	Remove minimum, new range 10.6
g49	2	11.3	
g50	2	10.1	
g51	1	na	
g52	2	3.5	

Appendix C

Benchmark data sheets are available at <http://www.ngs.noaa.gov>.

NQ0176 *****

NQ0176 FBN - This is a Federal Base Network Control Station.

NQ0176 DESIGNATION - Z 63

NQ0176 PID - NQ0176

NQ0176 STATE/COUNTY- WY/PLATTE

NQ0176 USGS QUAD - GUERNSEY (1990)

NQ0176

NQ0176 *CURRENT SURVEY CONTROL

NQ0176

NQ0176* NAD 83(1993)- 42 15 36.11667(N) 104 43 04.63639(W) ADJUSTED

NQ0176* NAVD 88 - 1331.270 (meters) 4367.68 (feet) ADJUSTED

NQ0176

NQ0176 X - -1,201,352.483 (meters) COMP

NQ0176 Y - -4,573,443.943 (meters) COMP

NQ0176 Z - 4,267,909.595 (meters) COMP

NQ0176 LAPLACE CORR- -7.92 (seconds) DEFLEC99

NQ0176 ELLIP HEIGHT- 1316.30 (meters) (02/28/01) GPS OBS

NQ0176 GEOID HEIGHT- -14.95 (meters) GEOID99

NQ0176 DYNAMIC HT - 1330.450 (meters) 4364.98 (feet) COMP

NQ0176 MODELED GRAV- 979,959.1 (mgal) NAVD 88

NQ0176

NQ0176 HORZ ORDER - B

NQ0176 VERT ORDER - SECOND CLASS 0

NQ0176 ELLP ORDER - SECOND CLASS I

NQ0176

NQ0176.The horizontal coordinates were established by GPS observations

NQ0176.and adjusted by the National Geodetic Survey in October 1994.

NQ0176

NQ0176.The orthometric height was determined by differential leveling

NQ0176.and adjusted by the National Geodetic Survey in June 1991.

NQ0176

NQ0176.Photographs are available for this station.

NQ0176

NQ0176.The X, Y, and Z were computed from the position and the ellipsoidal ht.

NQ0176

NQ0176.The Laplace correction was computed from DEFLEC99 derived deflections.

NQ0176

NQ0176.The ellipsoidal height was determined by GPS observations

NQ0176.and is referenced to NAD 83.

NQ0176

NQ0176.The geoid height was determined by GEOID99.

NQ0176

NQ0176.The dynamic height is computed by dividing the NAVD 88

NQ0176.geopotential number by the normal gravity value computed on the

NQ0176.Geodetic Reference System of 1980 (GRS 80) ellipsoid at 45

NQ0176.degrees latitude (g = 980.6199 gals.).

NQ0176

NQ0176.The modeled gravity was interpolated from observed gravity values.

NQ0176

NQ0176; North East Units Scale Converg.
NQ0176;SPC WY E - 195,556.714 237,022.125 MT 0.99995436 +0 18 06.3
NQ0176;UTM 13 - 4,678,686.487 523,263.003 MT 0.99960666 +0 11 22.8

NQ0176

NQ0176 SUPERSEDED SURVEY CONTROL

NQ0176

NQ0176 ELLIP HT - 1316.32 (m) (10/19/94) GP() 4 1

NQ0176 NGVD 29 - 1330.494 (m) 4365.13 (f) ADJ UNCH 2 0

NQ0176

NQ0176.Superseded values are not recommended for survey control.

NQ0176.NGS no longer adjusts projects to the NAD 27 or NGVD 29 datums.

NQ0176.See file dsdata.txt to determine how the superseded data were derived.

NQ0176

NQ0176_U.S. NATIONAL GRID SPATIAL ADDRESS: 13TEG2326378686(NAD 83)

NQ0176_MARKER: DB = BENCH MARK DISK

NQ0176_SETTING: 7 = SET IN TOP OF CONCRETE MONUMENT

NQ0176_STAMPING: Z 63 1934

NQ0176_MARK LOGO: CGS

NQ0176_MAGNETIC: N = NO MAGNETIC MATERIAL

NQ0176_STABILITY: C = MAY HOLD, BUT OF TYPE COMMONLY SUBJECT TO

NQ0176+STABILITY: SURFACE MOTION

NQ0176_SATELLITE: THE SITE LOCATION WAS REPORTED AS SUITABLE FOR

NQ0176+SATELLITE: SATELLITE OBSERVATIONS - November 04, 1999

NQ0176

NQ0176 HISTORY - Date Condition Report By

NQ0176 HISTORY - 1934 MONUMENTED CGS

NQ0176 HISTORY - 1941 GOOD NGS

NQ0176 HISTORY - 19930725 GOOD WYDT

NQ0176 HISTORY - 19991104 GOOD NGS

NQ0176

NQ0176 STATION DESCRIPTION

NQ0176

NQ0176'DESCRIBED BY NATIONAL GEODETIC SURVEY 1941

NQ0176'1.6 MI SE FROM GUERNSEY.

NQ0176'1.6 MILES SOUTHEAST ALONG THE CHICAGO, BURLINGTON AND QUINCY

NQ0176'RAILROAD FROM THE STATION AT GUERNSEY, 0.4 MILE NORTHWEST OF

NQ0176'MILEPOST 93, AT THE WEST END OF A LONG CUT ON A CURVE, 55 FEET

NQ0176'SOUTHWEST OF THE CENTER OF A DIM-ROAD CROSSING, 25 FEET SOUTH

NQ0176'OF THE SOUTH RAIL, AND 14.5 FEET WEST OF A POLE. A STANDARD DISK,

NQ0176'STAMPED Z 63 1934 AND SET IN THE TOP OF A CONCRETE POST.

NQ0176'NOTE-- THE MARK IS 4 POLES EAST OF A YARD LIMIT SIGN AND 10 FEET

NQ0176'WEST OF A TELEPHONE POLE.

NQ0176

NQ0176 STATION RECOVERY (1993)

NQ0176

NQ0176'RECOVERY NOTE BY WYOMING DEPARTMENT OF TRANSPORTATION 1993 (RR)

NQ0176'STATION IS LOCATED ABOUT 1.5 KM (0.95 MI) SOUTHEAST OF GUERNSEY, 0.9

NQ0176'KM (0.55 MI) SOUTHWEST OF US HIGHWAY 26, ALONG THE OLD HIGHWAY, NEAR

NQ0176'THE SOUTHWEST CORNER OF SECTION 36, T 27 N, R 66 W. OWNERSHIP--STATE

NQ0176'HIGHWAY DEPARTMENT. TO REACH FROM THE JUNCTION OF US HIGHWAY 26

AND

NQ0176'WHALEN STREET IN THE EAST SECTION OF GUERNSEY, GO SOUTHEAST ON

HIGHWAY

NQ0176'26 FOR 1.17 KM (0.70 MI) TO A PAVED ROAD RIGHT. TURN RIGHT,

NQ0176'SOUTHERLY, ON ROAD FOR 0.2 KM (0.10 MI) TO THE GUERNSEY AIRPORT.

NQ0176'CONTINUE AHEAD FOR 0.93 KM (0.55 MI) TO THE STATION ON THE LEFT.

NQ0176'STATION MARK IS A DISK SET IN THE TOP OF A 15-CM SQUARE CONCRETE POST
NQ0176'PROJECTING 15 CM ABOVE GROUND. IT IS 37.2 M (122.0 FT) NORTHEAST OF,
NQ0176'AND 3 M (9.8 FT) LOWER THAN THE ROAD CENTER, 7.5 M (24.6 FT) SOUTH OF
NQ0176'THE SOUTH RAIL OF RAILROAD TRACK, 4.9 M (16.1 FT) WEST OF A TELEPHONE
NQ0176'POLE, 28.7 M (94.2 FT) SOUTHEAST OF A TELEPHONE POLE, AND 0.3 M (1.0
NQ0176'FT) NORTHEAST OF A METAL WITNESS POST. DESCRIBED BY WTD, TYPED BY
GRH.

NQ0176

NQ0176

STATION RECOVERY (1999)

NQ0176

NQ0176'RECOVERY NOTE BY NATIONAL GEODETIC SURVEY 1999 (CSM)

NQ0176'RECOVERED BY NATIONAL GEODETIC SURVEY 1999. RECOVERED IN GOOD
NQ0176'CONDITION, NEW DESCRIPTION FOLLOWS. THE STATION IS LOCATED ABOUT 3.14
NQ0176'KM (1.95 MI) EAST-SOUTHEAST OF GUERNSEY, 0.97 KM (0.60 MI) SOUTHEAST
NQ0176'OF THE ENTRANCE ROAD OF THE GUERNSEY AIRPORT, ON THE SOUTH SIDE OF
THE

NQ0176'BURLINGTON NORTHERN AND SANTA FE RAILROAD TRACKS.

NQ0176'OWNERSHIP--BURLINGTON NORTHERN AND SANTA FE RAILWAY COMPANY, 100
NQ0176'CEMETERY ROAD, GUERNEY, WY 82214. CONTACT STEVE KETCHUM PHONE
NQ0176'307-836-5200. NOTE--MUST HAVE REQUIRED OSHA APPROVED SAFETY GLASSES
NQ0176'WITH SIDE SHIELDS, HARD HAT AND ABOVE THE ANKLE SAFETY TOE BOOTS TO
NQ0176'ACCESS THE SITE. TO REACH THE STATION FROM THE EAST END OF THE US
NQ0176'HIGHWAY 26 BRIDGE OVER THE NORTH PLATTE RIVER ON THE WEST SIDE OF
NQ0176'GUERNSEY, GO EAST ON HIGHWAY 26 (WHALEN STREET) FOR 2.01 KM (1.25 MI)
NQ0176'TO A PAVED ROAD RIGHT (AT SIGN--AIRPORT AND JUST PAST HIGHWAY
MILEPOST

NQ0176'16) , TURN RIGHT AND GO SOUTH THEN EAST ON THE ROAD FOR 0.16 KM (0.10
NQ0176'MI) TO THE ENTRANCE ROAD TO THE AIRPORT ON THE RIGHT, CONTINUE AHEAD
NQ0176'EAST ON THE ROAD FOR 0.97 KM (0.60 MI) TO THE STATION ON THE RIGHT.

NQ0176'FROM THIS POINT PACK ABOUT 30.48 M (100.00 FT) DOWNHILL TO THE STATION
NQ0176'ON THE SOUTH SIDE OF THE RAILROAD TRACKS. THE STATION IS AN CGS BENCH
NQ0176'MARK DISK SET IN THE TOP OF 15 CM SQUARE CONCRETE POST ABOUT FLUSH
NQ0176'WITH THE GROUND, 3.06 M (10.04 FT) LOWER THAN THE ROAD AND 0.3 M (1.0
NQ0176'FT) LOWER THAN THE TRACKS. IT IS, 37.2 M (122.0 FT) NORTH OF THE
NQ0176'CENTER OF THE ROAD, 7.5 M (24.6 FT) SOUTH OF THE SOUTH RAIL OF THE
NQ0176'TRACKS, 0.3 M (1.0 FT) WEST OF AN NGS FIBERGLASS WITNESS POST AND 0.77
NQ0176'M (2.53 FT) WEST OF A WYDOT FIBERGLASS WITNESS POST AND SURROUNDED BY
NQ0176'3 4X4 WOOD POSTS PROJECTING ABOUT 0.77 M (2.53 FT) ABOVE THE GROUND.
NQ0176'DESCRIBED BY M.L. MCCREADY WITH NOTES BY R.G. BAILEY.

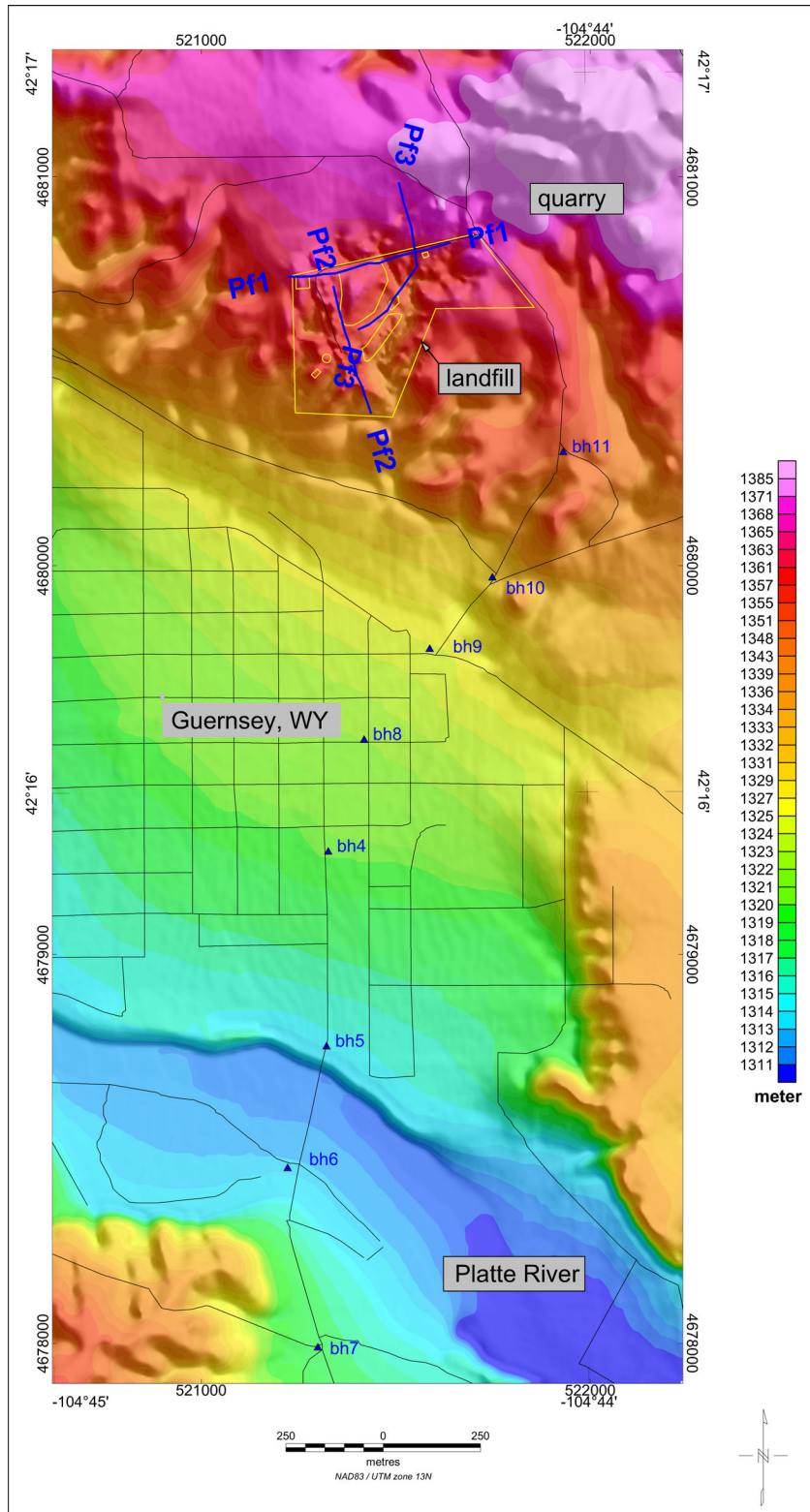


figure 1. Topography of Guernsey and landfill at 5-meter resolution with gravity profiles and landfill features superimposed.

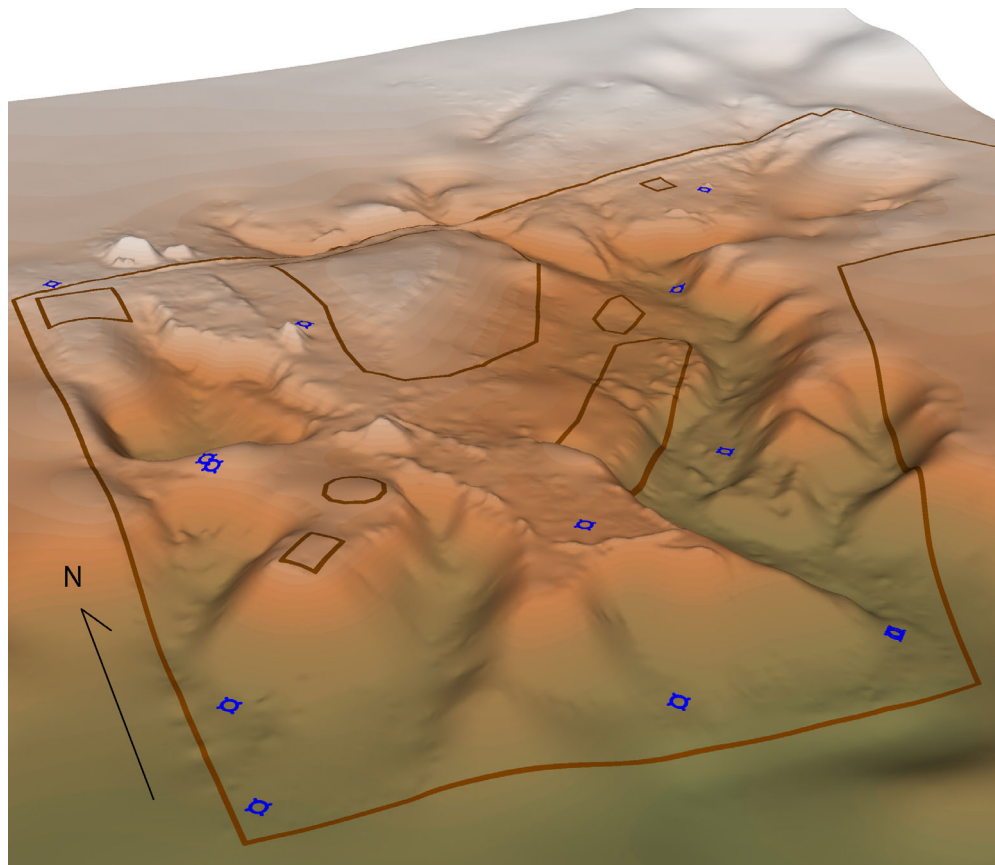
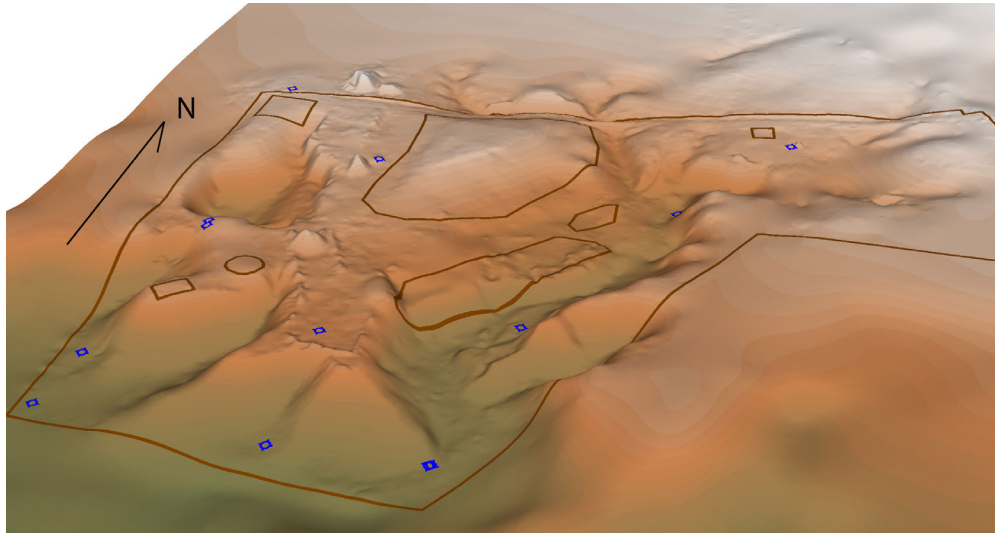


figure 2. Shaded-relief maps of topography from real-time kinematic GPS survey combined with DEM 10-meter data. Vertical exaggeration is 2X. Brown lines define landfill boundary and key features. Blue circles indicate well locations.

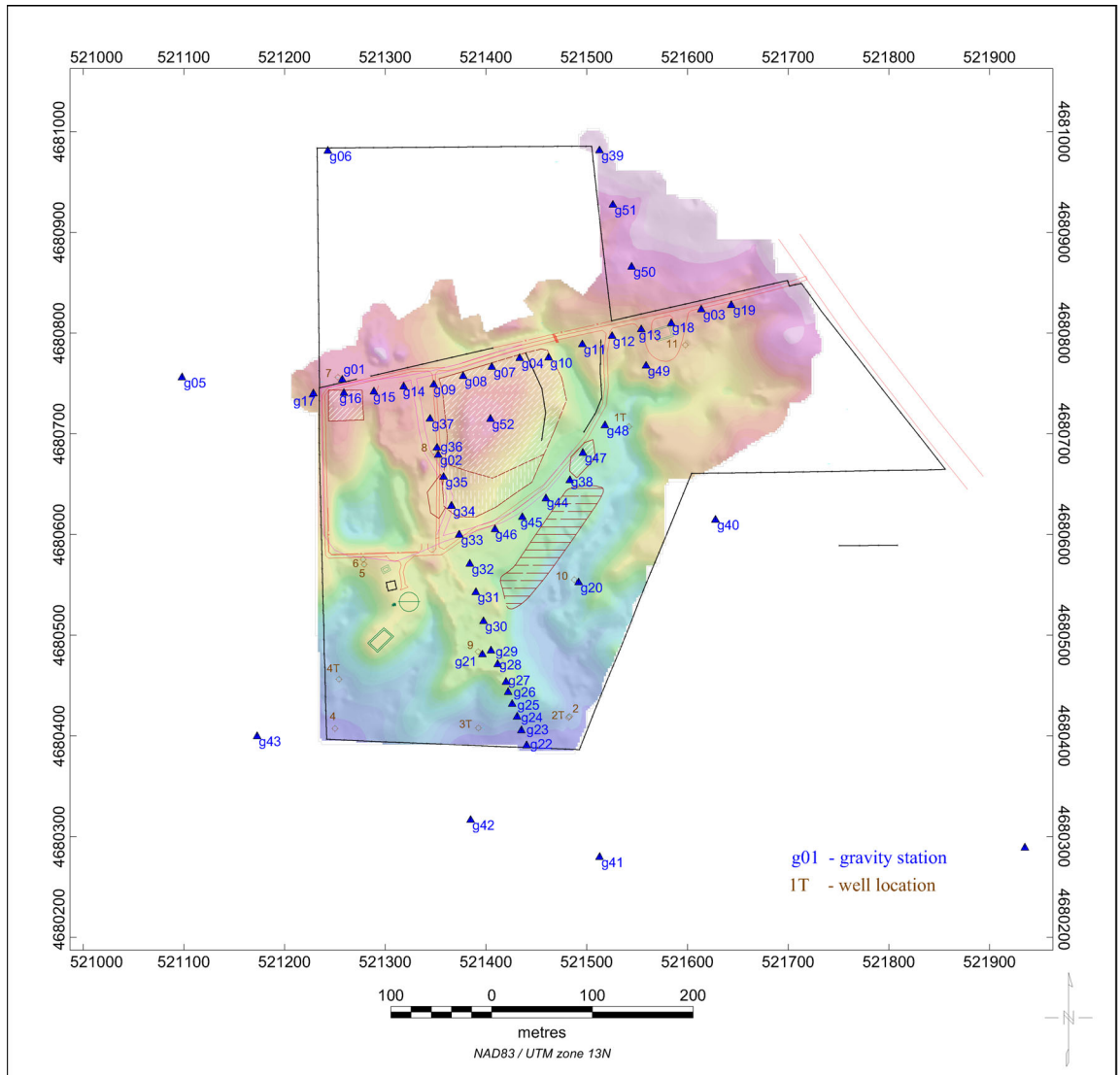


figure 3. Site map of Guernsey landfill showing placement of gravity stations superimposed over 2-meter resolution shaded-relief topography.

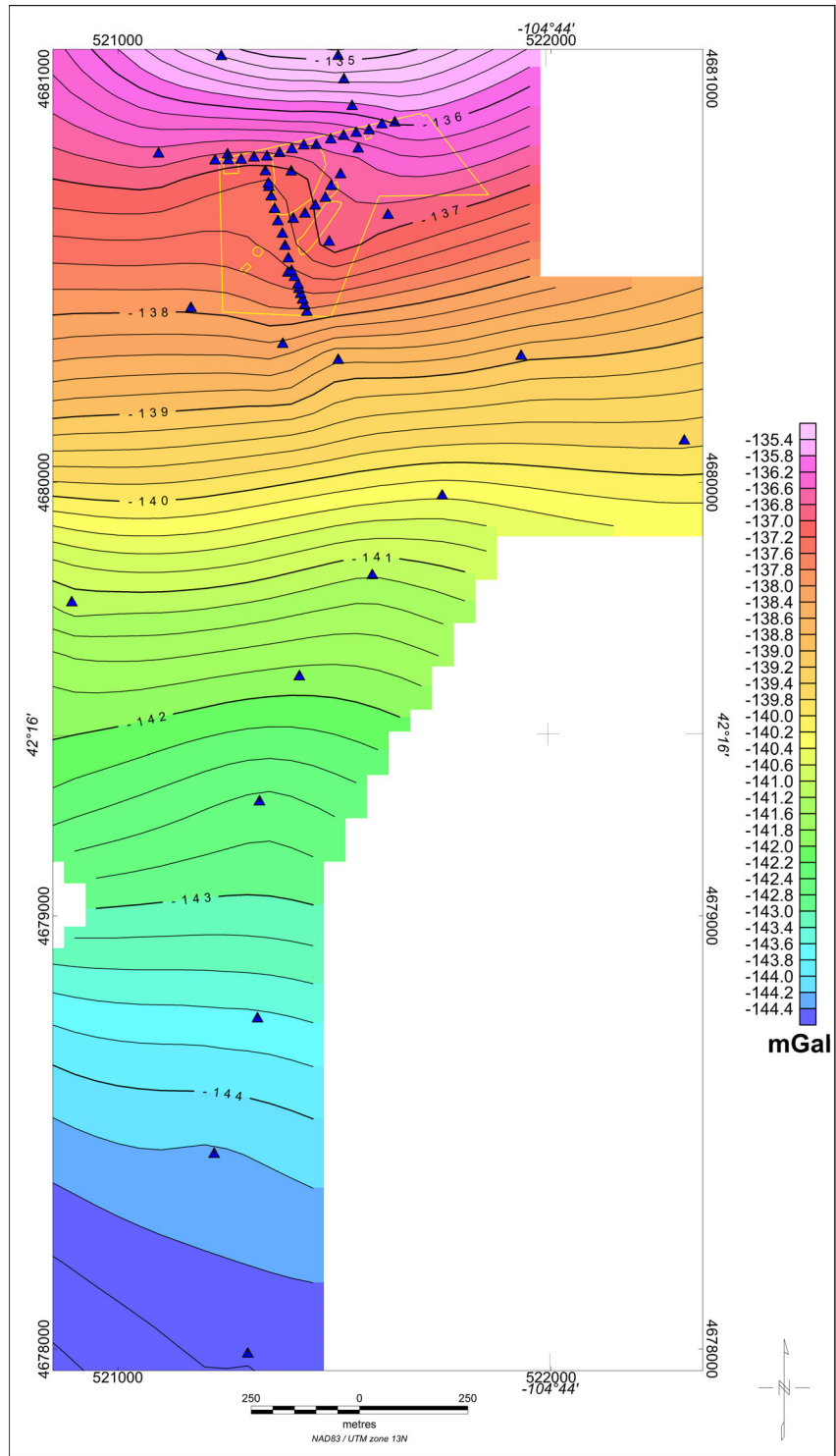


figure 4. Bouguer gravity anomaly with station locations superimposed.

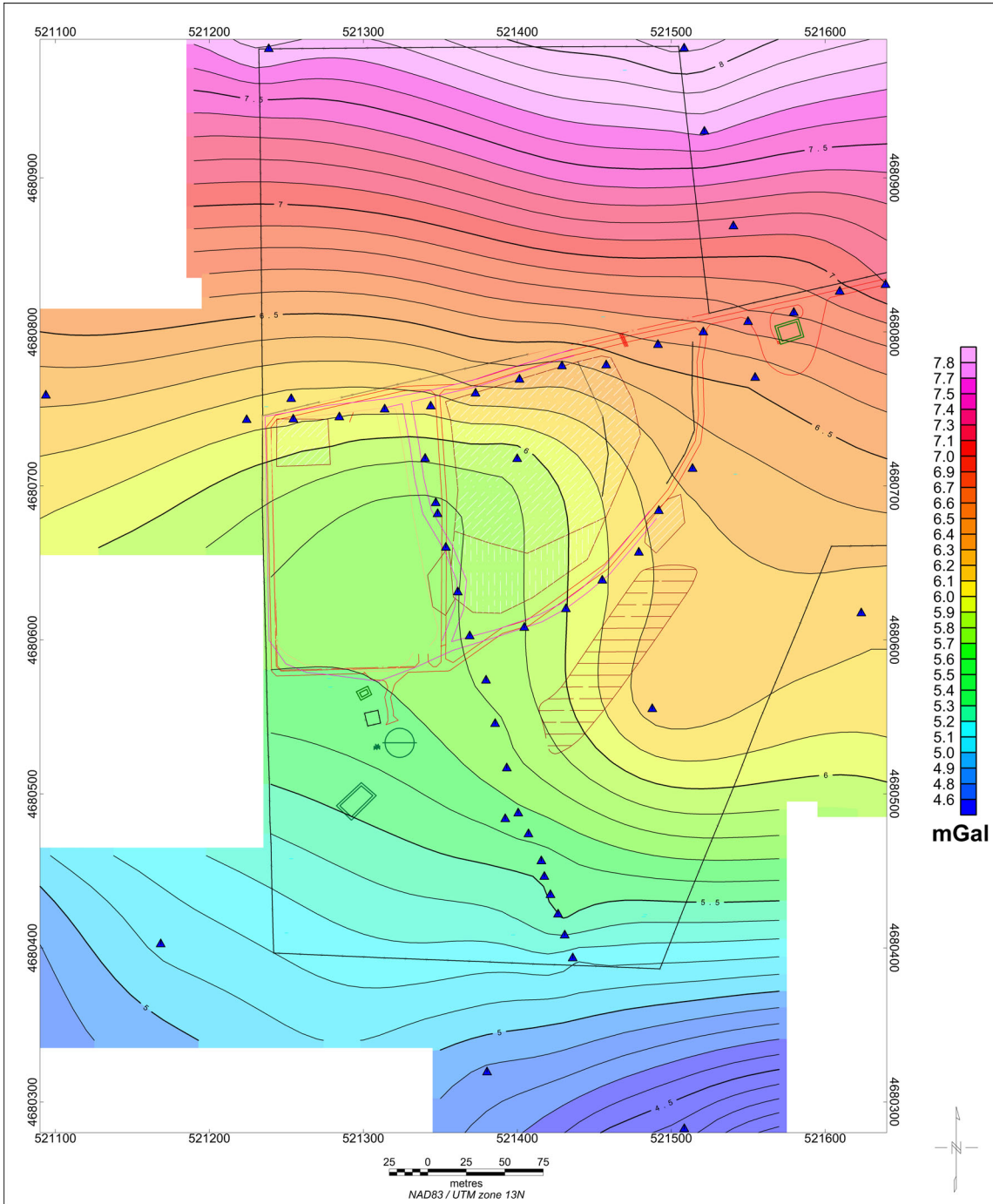


figure 5. Regional residual Bouguer gravity anomaly of the Guernsey landfill.

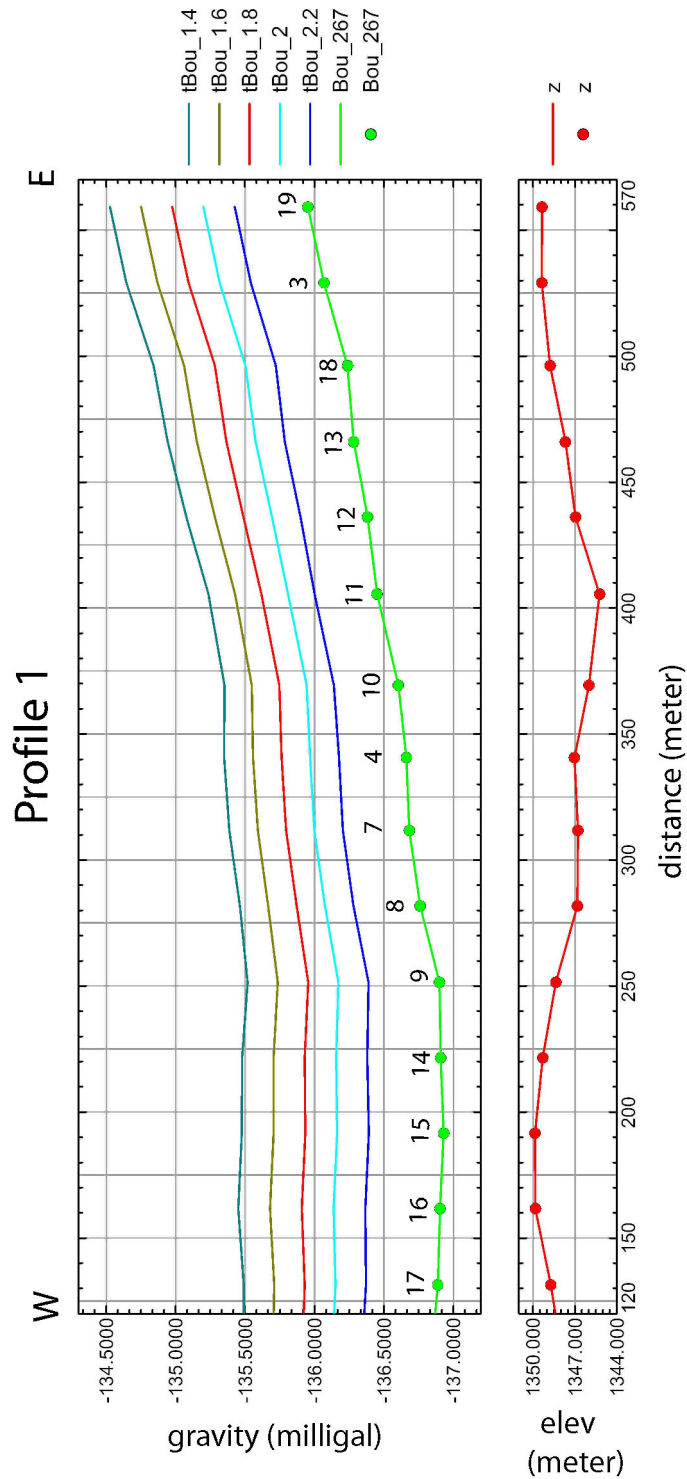


figure 6. Profile 1 showing Bouguer anomaly in red and the effect of terrace models at several densities. Gravity station locations are shown on the Bouguer and elevation curves.

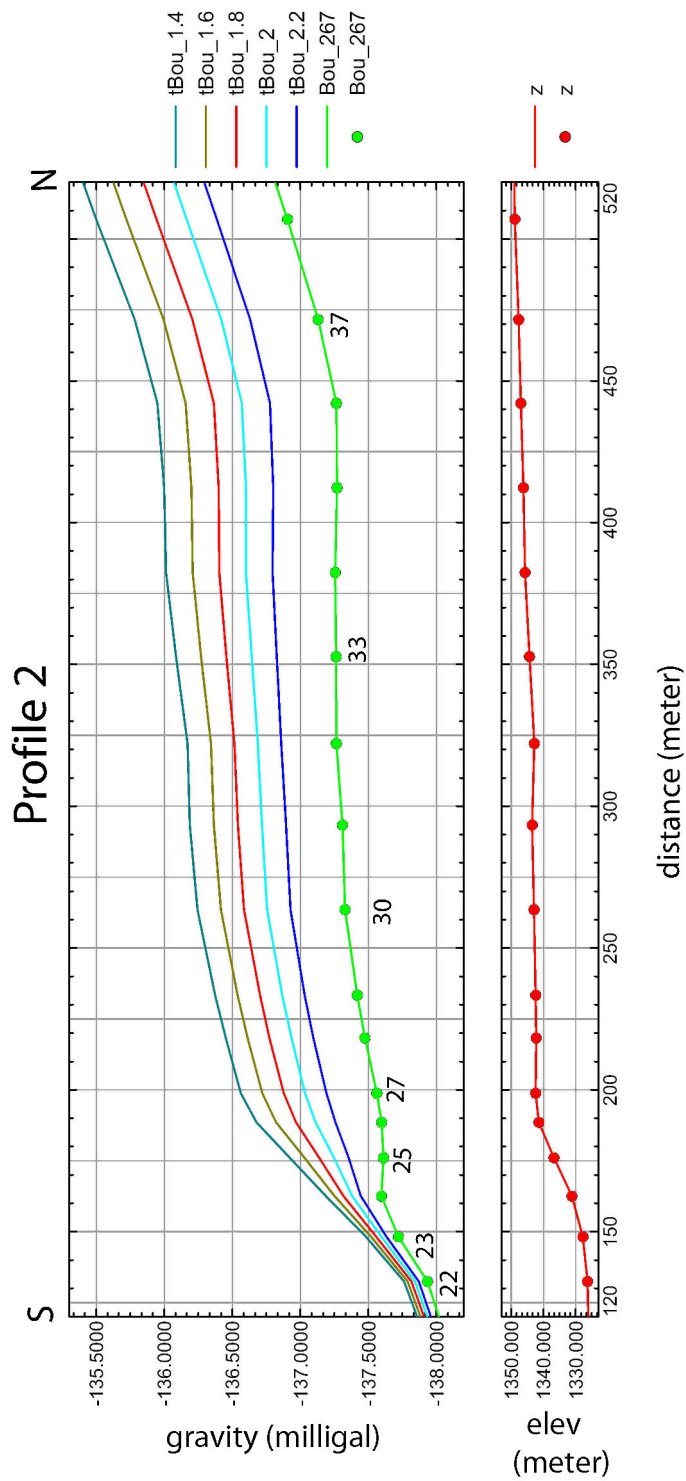


figure 7. South end of profile 2 showing Bouguer anomaly in red and effect of terrace models at several densities.

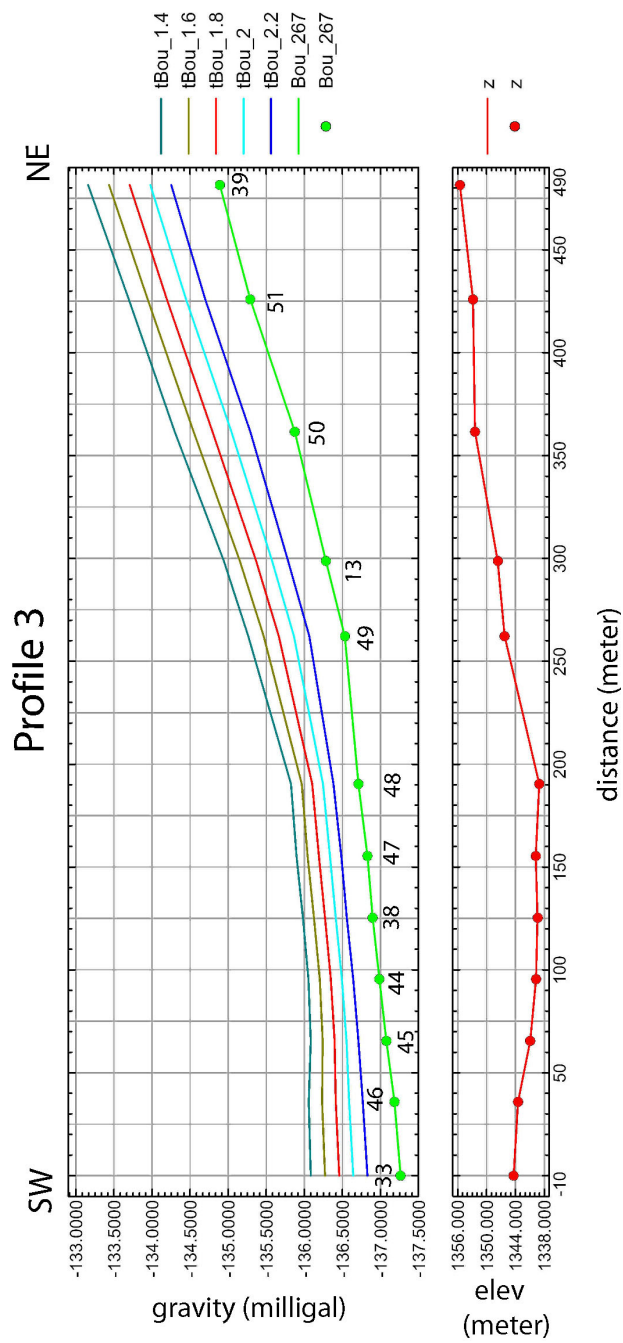


figure 8. Profile 3 showing Bouguer anomaly in red and the effect of terrace models at several densities.

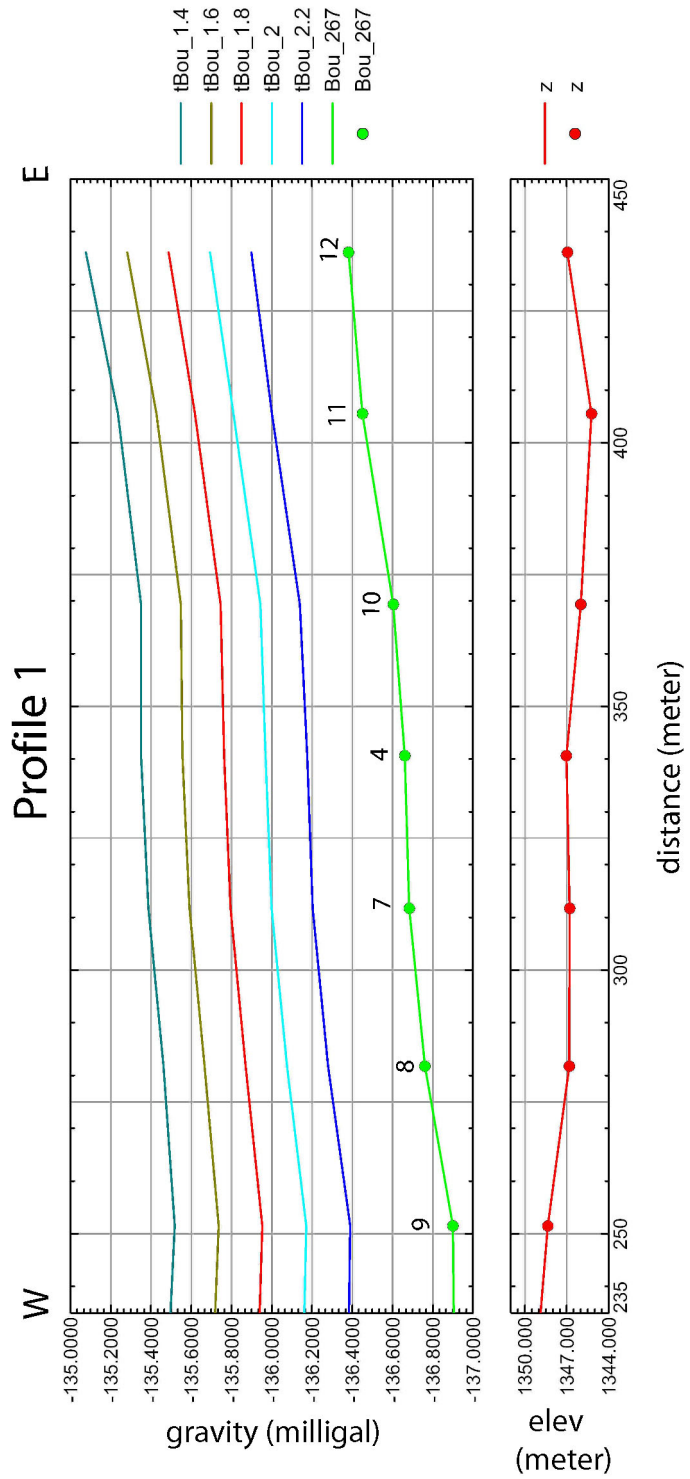


figure 9. Detail of profile 1 showing anomaly curvature reversals at stations g8, g4, and g11.

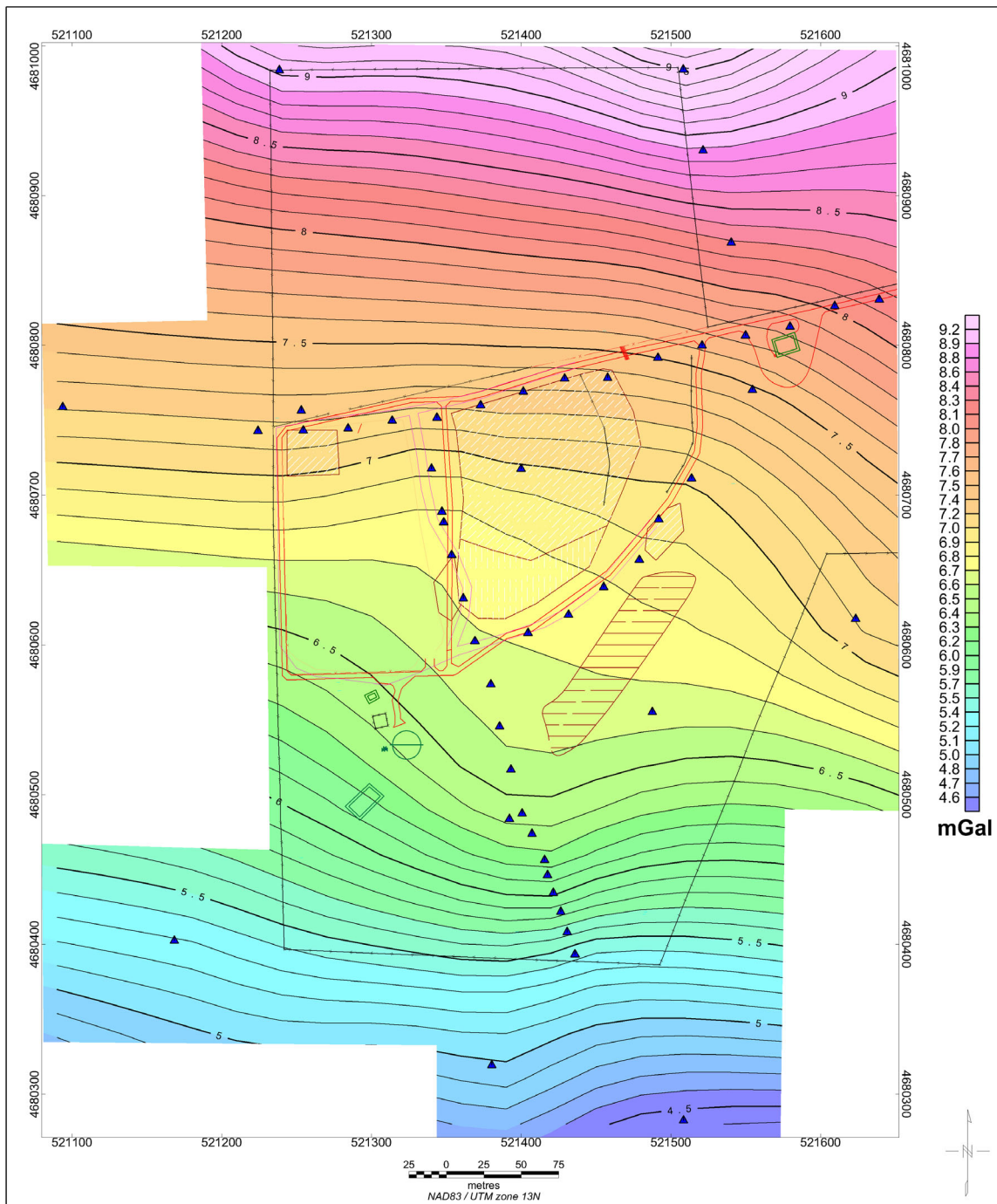


figure 10. Terrace residual gravity assuming a density of 1.80 gm/cc.

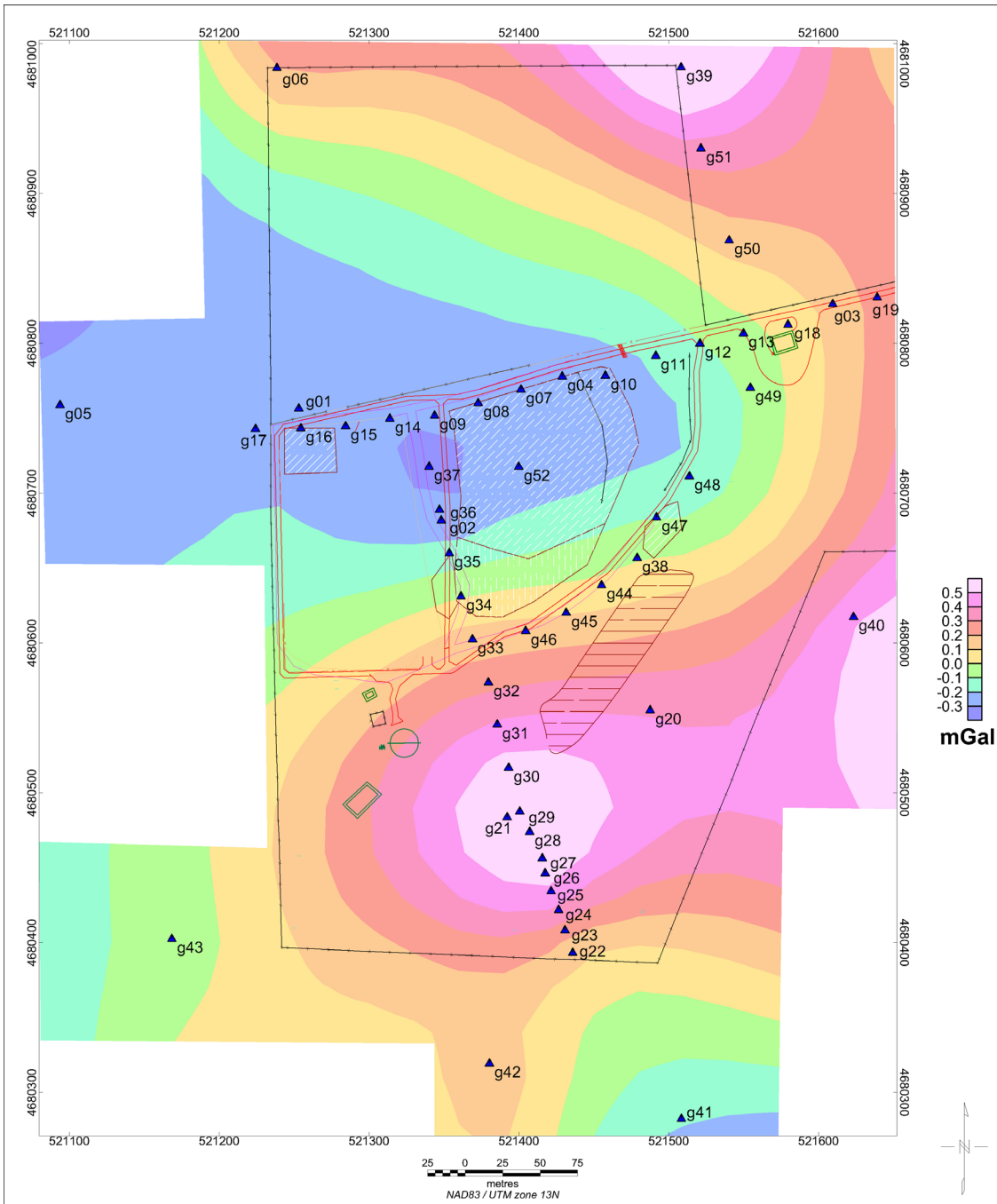


figure 11. De-trended terrace residual gravity. Landfill cells and roads in red, buildings in green.

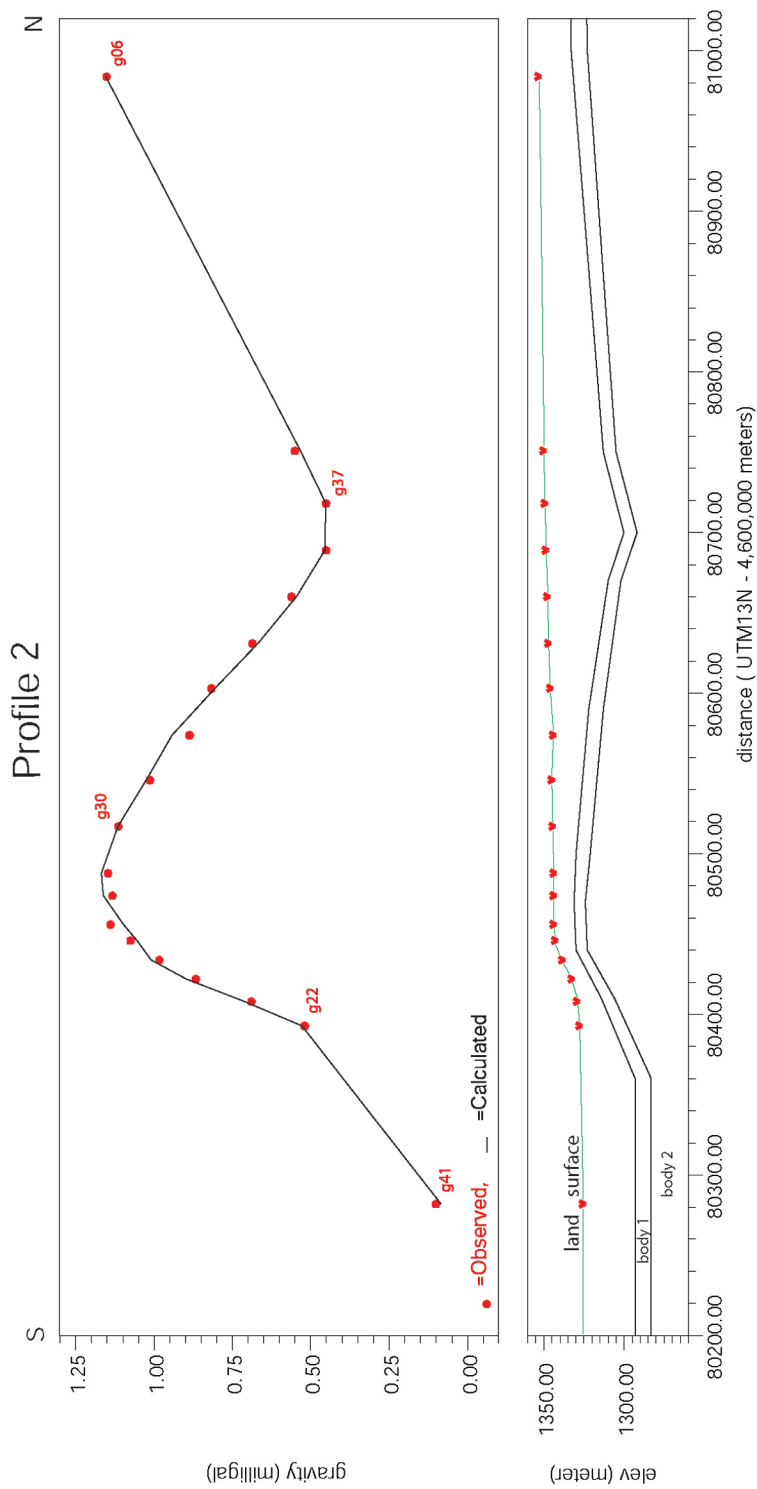


figure 12. 2.5D model from de-trended terrace residual gravity. Stations and gravity values shown in red, density reference is 1.80 gm/cc, body 1 has a density contrast of 0.87 gm/cc and body 2 a contrast of 1.02 gm/cc.