

ATMOSPHERIC MEASUREMENTS BY THE MLS EXPERIMENTS: RESULTS FROM UARS AND PLANS FOR EOS

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ABSTRACT

The Microwave Limb Sounder (MLS) on the Upper Atmosphere Research Satellite (UARS) has provided measurements of O_3 , H_2O , ClO , SO_2 , HNO_3 , temperature and pressure in Earth's atmosphere. These measurements, which have been used for a variety of scientific studies, are made near-globally both day and night and are not degraded by the presence of aerosols, cirrus or polar stratospheric clouds. The MLS now being developed for the NASA Earth Observing System (EOS) uses new technology for measurements of additional trace species and improved global coverage.

INTRODUCTION

Microwave limb sounding obtains remote measurements of atmospheric parameters by observations of millimeter and submillimeter-wavelength thermal emission as the instrument field-of-view (FOV) is scanned through the atmospheric limb from above. The technique is described by Waters (1989; 1992a,b; 1993). Its features include: (1) the ability to measure many atmospheric gases, with emission from molecular oxygen providing temperature and pressure; (2) measurements which are not degraded by aerosols, cirrus or polar stratospheric clouds; (3) the ability to make measurements at all times of day and night; (4) the ability to spectrally-resolve emission lines at all altitudes which allows measurements of very weak lines in the presence of nearby strong ones; (5) composition measurements which are very insensitive to uncertainties in atmospheric temperature; (6) a very accurate spectroscopic data base; (7) instrumentation which has very accurate and stable calibration, excellent sensitivity without necessarily requiring cooling, can be modularly designed for ease in accommodating changing measurement priorities, can provide good vertical resolution which is set by size of the antenna, and with new array technology can provide good horizontal resolution including complete coverage between orbits.

Development of the MLS experiments, which began at the Jet Propulsion Laboratory in the mid-1970's, included instruments deployed on aircraft (Waters *et al.*, 1979, 1980) and balloon (Waters *et al.*, 1981, 1984, 1988; Stachnik *et al.*, 1992). The MLS launched 12 September 1991 on the Upper Atmosphere Research Satellite (e.g., Reber, 1993; Reber *et al.*, 1993; Waters 1996; Jackman *et al.* this volume) is the first application of the microwave limb sounding technique from space. The Millimeter-Wave Atmospheric Sounder, MAS (Croskey *et al.*, 1992), has also used the technique from the Space Shuttle. UARS MLS, at the time of writing this review, continues to operate after 5 years in orbit with no degradation in its 63 and 205 GHz measurements, except for (1) time-sharing of these measurements with those of other UARS instruments due to power constraints from the spacecraft and (2) using a lower stratospheric "limb-tracking" mode about every third day of measurements to extend lifetime of the antenna scan mechanism. The MLS 183 GHz measurements ceased in April 1993 after 18 months of excellent data had been obtained. Development is underway for a next-generation MLS instrument to be deployed on the NASA Earth Observing System (EOS), with launch planned in 2002." This paper summarizes results to date obtained from UARS MLS, and describes the planned capability of the EOS MLS.

THE UARS MLS AND RESULTS TO DATE

Development of the UARS MIA experiment was led by the California Institute of Technology Jet Propulsion Laboratory, with collaboration from Rutherford Appleton Laboratory, Heriot-Watt University, and Edinburgh University in the United Kingdom. The instrument (Barath *et al.*, 1993) contains ambient-temperature heterodyne radiometers which operate near 63 GHz for measurements of temperature and pressure; 205 GHz for measurements of stratospheric O₃, ClO, SO₂, HNO₃ and upper tropospheric H₂O; and 183 GHz for stratospheric and mesospheric H₂O and O₃. Calibration of the instrument, described by Jarnot *et al.* (1996), is accurate to ~3% overall. Validation of the MIS primary data products, and their accuracies and precision, are described in the *Journal of Geophysical Research* special issue on UARS data evaluation: temperature/pressure by Fishbein *et al.* (1996); O₃ by Froidevaux *et al.* (1996), Cunnold *et al.* (1996 a, b), and Ricaud *et al.* (1996); H₂O by Lahaot *et al.* (1996a); ClO by Waters *et al.* (1996). Additional results relevant to validation of these MIS measurements are included in Aellig *et al.* (1996), Crewell *et al.* (1995), Redaelli *et al.* (1995), Singh *et al.* (1996) and Wild *et al.* (1995). Other data products which have been obtained from UARS MLS, beyond that for which the instrument was primarily designed, include SO₂ injected into the stratosphere by the Pinatubo volcano (Read *et al.*, 1993), upper tropospheric H₂O (Read *et al.*, 1995), lower stratospheric HNO₃ (Santee *et al.*, 1995), temperature variances associated with atmospheric gravity waves in the stratosphere and mesosphere (WU and Waters, 1996a), and geopotential height (Fishbein *et al.*, in preparation). Fourier-transform techniques applied to mapping MIS data are described by Elson and Froidevaux (1993). Information on the spectroscopic data used in obtaining the MIS atmospheric measurements can be found, for example, in Pickett *et al.* (1981, 1992), Poynter and Pickett (1985), Oh and Cohen (1992, 1994).

Starting within 10 days of launch, and continuing for approximately 2 months, MIS observed the 3-dimensional distribution and decay of residual SO₂ injected into the tropical stratosphere by the Mt. Pinatubo eruption which occurred about 3 months before launch of UARS. These observations (Read *et al.*, 1993) showed the Pinatubo SO₂ mixing ratio maximum to occur around 26 km altitude with abundances of ~15 ppbv on 21 September 1991. The observed SO₂ decay had e-folding times of 29 days at 26 km and 41 days at 21 km, consistent with expectations that the primary destruction of SO₂ is due to reaction with OH leading to formation of stratospheric sulfate aerosols. Projected backward to time of eruption, the total amount of SO₂ injected by Pinatubo is estimated from MIS data to be 17 Mtons, consistent with estimates inferred from other measurements.

Early results from UARS MIS also included the first global maps of stratospheric ClO, the predominant form of chemically-reactive chlorine involved in the destruction of stratospheric O₃. The initial MIS results (Waters *et al.*, 1993a,b; see also Chipperfield, 1993) showed the lower stratospheric Antarctic vortex to be filled with ClO in the region where O₃ was depleted, confirming earlier conclusions from ground-based and "aircraft instruments that chlorine chemistry is the cause of the Antarctic ozone hole. They showed, for the first time, that ClO in the Antarctic vortex can become enhanced by June, and that O₃ destruction by ClO is masked in the early Antarctic winter by influx of O₃ expected from diabatic descent. These results also showed that the Arctic winter lower stratospheric vortex can become filled with enhanced ClO, leading to calculated vortex-averaged O₃ destruction rates of ~0.7%/day. Results from 3D models (Douglass *et al.*, 1993; Geller *et al.*, 1993; Lefevre *et al.*, 1994), produced shortly after the MIS results were obtained, showed the observed distribution of enhanced Arctic ClO was consistent with chemical-transport model predictions. A clear relationship was found between predicted polar stratospheric cloud formation along back trajectories and enhanced Arctic ClO observed by MIS, and sporadic large values of ClO seen by MIS outside the vortex were shown to be consistent with that expected to be caused by instrument noise (Schoeberl *et al.*, 1993). Definitive loss of Arctic ozone due to chemistry associated with the enhanced ClO was determined from analyses of combined MIS and UARS CLAES data by Manney *et al.* (1994). Bell *et al.* (1994) found the expected anticorrelation between enhanced Arctic ClO measured by MIS and HCl measured from the ground. Additional confirmation of the paradigm of chemical processing by polar stratospheric clouds leading to activation of stratospheric chlorine is shown in the analyses of northern hemisphere CLAES, MIS and HALOE data by Geller *et al.* (1995), and in southern hemisphere MIS and CLAES data by Ricaud *et al.* (1995). Differences between the Arctic and Antarctic winter vortex conditions as deduced from MIS observations are described by Santee *et al.* (1995), and deduced from combined MIS,

CLAES and HALOE data by Douglass *et al.* (1995). Mackenzie *et al.* (1996) compare lower stratospheric vortex ozone destruction calculated from the MLS ClO with the MLS-observed change in O₃ for the northern winter of 1992-93 and southern winter of 1993. Additional comparisons between MLS observations and model results for polar chemistry are given by Ekman *et al.* (1995), Chipperfield *et al.* (1996) and Santec *et al.* (1996a). Schoeberl *et al.* (1996a) use MLS, HALOE and CLAES data in an analysis of the development of the Antarctic ozone hole, MLS measurements of Arctic ClO and O₃ for the five northern winters observed to date are described in the collective papers of Waters *et al.* (1993a,b; 1995), Manney *et al.* (1994; 1995a,b; 1996a,b), and Santec *et al.* (1995, 1996b). Low ozone "pockets" in the middle stratospheric winter anticyclone have been observed in MLS data and analyzed by Manney *et al.* (1995c), who conclude these cannot be explained solely by transport.

MLS observations have been used in several studies to provide information on vortex and high-latitude dynamics. Early results (Harwood *et al.*, 1993) showed large parcels of air from the Antarctic vortex migrating to midlatitudes. Other studies of high-latitude dynamics which use MLS stratospheric data include those of Fishbein *et al.* (1993), Lahoz *et al.* (1993, 1994, 1996b), Manney *et al.* (1993; 1995d,c), and Morris *et al.* (1995). Orsolini *et al.* (1996) use MLS O₃ data to initialize a high-resolution transport model and examine ozone laminae along the Arctic polar vortex edge.

An overview of zonal mean O₃ results from the first two and one-half years of MLS operation is given by Froidevaux *et al.* (1994); in addition to features observed in stratospheric O₃, this work includes initial results of examining residual differences between the stratospheric O₃ column from MLS and the total O₃ column from TOMS -- with information on tropospheric ozone as the ultimate goal. Analyses by Ziemke *et al.* (1996) using these data sets have shown zonal asymmetries in southern hemisphere column ozone that have implications for biomass burning. Elson *et al.* (1994) describe large-scale variations observed in MLS O₃, and Elson *et al.* (1996) show zonal and large-scale variations in MLS H₂O. Randel *et al.* (1995) include MLS and HALOE data in analyzing changes in stratospheric ozone following the Pinatubo eruption. Dessler *et al.* (1995; 1996a,b) used MLS ClO and O₃ data, along with that of other UARS instruments, to provide information on various aspects of stratospheric chlorine chemistry. The latitudinal distribution of ClO in the upper stratosphere (Waters *et al.*, 1996) shows a minimum in the tropics as expected from quenching by increased amounts of upper stratospheric CH₄ in the tropics. Two-day waves in the stratosphere have been analyzed by Limpasuvan and Leovy (1995) using MLS H₂O data, and by Wu *et al.* (1996) using MLS temperatures. Four-day waves observed in MLS ozone, temperature and geopotential height have been analyzed by Allen *et al.* (1996). MLS data have been used in calculations of stratospheric residual circulation by Rosenlof (1995) and Eluszkiewicz *et al.* (1996). Stone *et al.* (1996) used MLS upper tropospheric H₂O measurements to investigate the structure and evolution of eastward-traveling medium-scale wave features in the southern hemisphere summertime, and found results consistent with paradigms for the structure and evolution of baroclinic disturbances.

Kelvin waves observed in MLS tropical data have been analyzed by Canziani *et al.* (1994, 1995) and Stone *et al.* (1995), and MLS observations of the semiannual oscillation by Ray *et al.* (1994). Randel *et al.* (1993) describe CLAES and MLS observations of stratospheric transport from the tropics to middle latitudes by planetary wave mixing. Carr *et al.* (1995) performed initial analyses of MLS tropical stratospheric H₂O data, and Mote *et al.* (1995) found variations in these data which could be related to the annual cycle in tropical tropopause temperatures. More extensive analyses by Mote *et al.* (1995), greatly aided by the use of HALOE H₂O and CH₄, confirmed that tropical air entering the stratosphere from below is marked by its water vapor mixing ratio and retains a distinct memory of tropical tropopause conditions for 18 months or more; this analysis implies that vertical mixing is weak and that subtropical stratospheric "transport barriers" are effective at inhibiting transport into the tropics. Schoeberl *et al.* (1996b) also use MLS and other UARS data to estimate the dynamical isolation of the tropical lower stratosphere. Newell *et al.* (1995a,b) have shown that the preliminary upper tropospheric H₂O from MLS are reasonably consistent with NASA ER-2 aircraft measurements, and consistent with the expected tropical Walker circulation. Newell *et al.* (1996c) found variations in MLS tropical upper tropospheric H₂O over the 1991-1994 period to be closely related to sea surface temperature variations in the eastern tropical Pacific, including both seasonal and interannual components.

Analyses of the 63 GHz radiances from MLS have produced the first global maps of atmospheric temperature variances associated with gravity wave activity in the stratosphere and mesosphere (WU and Waters, 1996a,b).

These data provide information on gravity waves with spatial scales of 30- 100 km in the horizontal and 10 km in the vertical. The mapped variances show high correlation with regions of strong background winds which are expected to play a major role in determining gravity wave amplitudes in the stratosphere and mesosphere. The observed variance grows exponentially with height in the stratosphere, and saturates in the mesosphere as expected from wave breaking and dissipation at the higher altitudes. The data also show some correlation with surface topography features and regions of tropospheric convective activity which are expected sources of gravity waves. The analysis of Alexander (1996) indicates that the MLS maps are consistent with model predictions of atmospheric gravity wave behavior but that the dominant patterns in the maps can be explained, without requiring any geographical variation in the sources, by Doppler-shifting effects of background winds on gravity waves having the vertical scales observed by MLS. The extent to which MLS can provide information on atmospheric gravity wave sources is a current topic of investigation. Analyses of data taken when the MLS FOV was tracking the limb have extended the range of horizontal scales to 15-5000 km (WU and Waters, 1996c).

PLANS FOR EOS MLS

The MLS planned for the NASA Earth Observing System will be improved over UARS MLS in providing, (1) additional stratospheric species, (2) more tropospheric/tropopause measurements, (3) better global coverage, and (4) better precision. These improvements are possible because of advances in microwave technology since the UARS MLS design was frozen. The additional species include OH, HO₂, HCl, HOCI, BrO, N₂O, CO and perhaps others. The broad spectral coverage of the EOS MLS radiometers operating near 100 and 200 GHz will provide measurements of H₂O, temperature and pressure to much lower altitudes in the troposphere than is possible with UARS MLS; initial analyses have indicated that, at least in cloud free regions, H₂O and temperature measurements can be made in the lower troposphere. Measurements of O₃, temperature, N₂O (and possibly HNO₃ and CO), as well as H₂O, will be made in the upper troposphere and tropopause regions. Figure 1 shows the planned EOS MLS measurement capability in the stratosphere and troposphere, and indicates the general range of frequencies of the radiometer spectral bands which will be used for the measurements.

A focal plane array of Millimeter-wavelength Monolithic Integrated Circuit (MMIC) radiometers (Weinreb, 1996) is now planned for the spectral bands at 120 and 190 GHz. The number of array elements is currently under study, but is expected to be between 5 and 20 in a horizontal row which projects conically on the atmospheric limb in the cross-track direction and provides orbit-to-orbit coverage. The horizontal resolution provided by such an array is 500 km for 5 elements and 130 km for 20 elements. Figure 2 shows horizontal coverage for a 6-element array, and is much improved from that of UARS MLS which only measures along a single track each orbit. The radiometers at 240 and 640 GHz will use planar mixer technology (e. g., Siegel *et al.*, 1993). High-temperature-superconductor hot-electron bolometer-mixer radiometers (e.g., McGrath, 1995), whose local oscillator can be generated by mixing of solid-state near-IR diode lasers (e.g., Brown *et al.*, 1995; Pickett *et al.*, 1996) are being considered for the 2.5 THz radiometer; such radiometers have substantially reduced power and mass requirements than the alternate Schottky-diode radiometer with gas-laser local oscillator being considered for this radiometer, and can provide measurements of additional atmospheric species beyond those shown in Figure 1. EOS MLS will have better precision for trace gas measurements than UARS MLS, both due to improvements in instrument sensitivity since UARS and by measurements of stronger spectral lines at higher frequencies. For example, the ClO precision will be at least 5x better from EOS MLS than from UARS.

The EOS MLS FOV will look in the general direction of the orbital path, in contrast to UARS MLS which looks perpendicular to the orbital path. This provides pole-to-pole coverage on each orbit every day, whereas UARS MLS high-latitude coverage switches approximately monthly between the northern and southern hemispheres.

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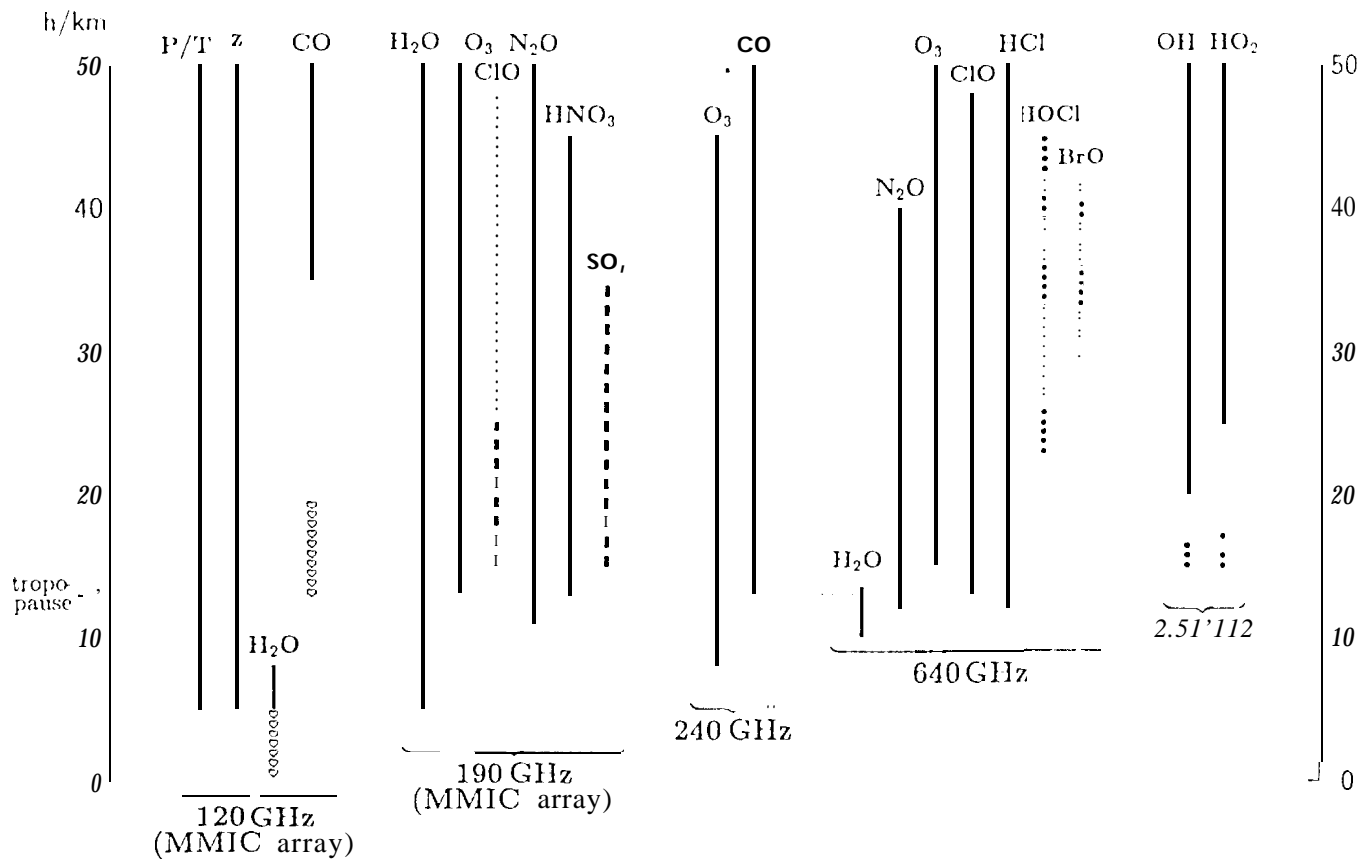


Figure 1. Measurement capability of the MIRS experiment now planned for the Earth Observing System. Solid are individual profile measurements, dotted arc zonal means, dashed arc special situations, and hearts indicate research goals. Many measurements extend higher than shown here.

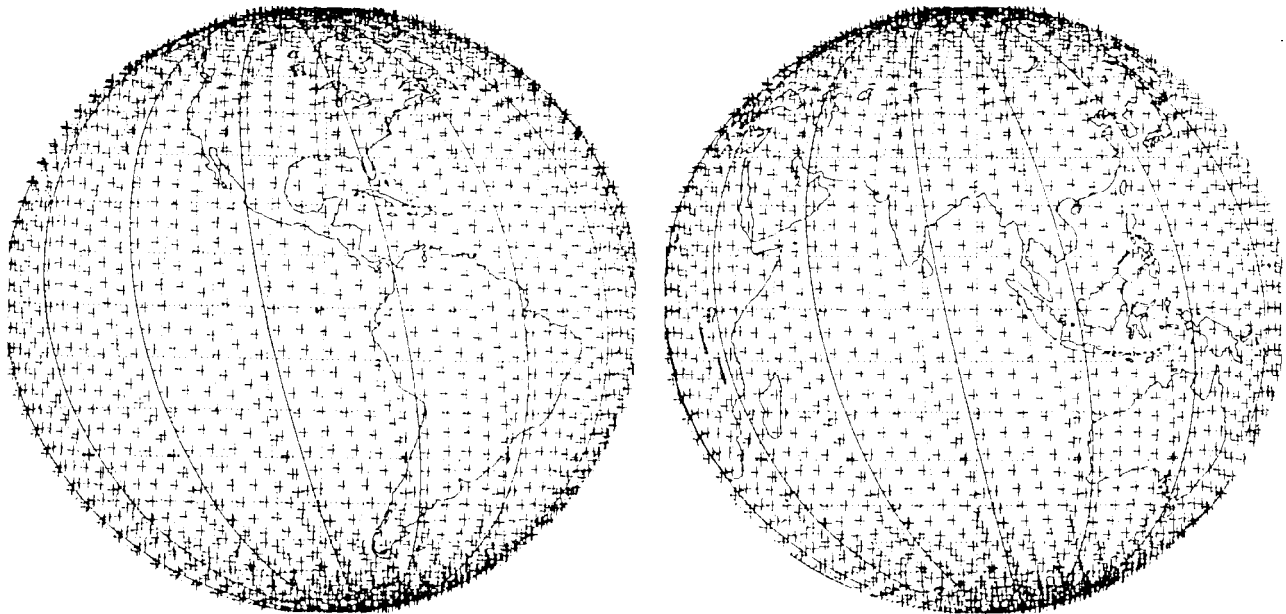


Figure 2. EOS MIRS 120 and 190 GHz measurement coverage obtained during each 12-hour period with an array having 6 "horizontal" elements. Crosses indicate locations of independent vertical profile measurements, and lines show suborbital paths. The number of array elements is currently under study, but the maximum number in a horizontal row is expected to be between 5 and 20.

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