DRAFT

15555 Olivine Basalt



Figure 1: Photo of S1 surface of 15555, illustrating large zap pit and vuggy nature of rock. NASA photo no. S71-43954. Scale is in cm.

Introduction

Lunar sample 15555 (called "Great Scott", after the collector Dave Scott) is one of the largest samples returned from the moon and is representative of the basaltic samples found on the mare surface at Apollo 15. It contains olivine and pyroxene phenocrysts and is olivine normative in composition (Rhodes and Hubbard 1973). The bulk composition of 15555 is thought to represent that of a primitive volcanic liquid and has been used for various experimental studies

related to the origin of lunar basalts (e.g. Walker et al. 1977).

15555 has a large zap pit (\sim 1 cm) on the S1 face, various penetrating fractures and a few percent vugs (figure 1). It has a subophitic, basaltic texture (figure 3) and there is little evidence for shock in the minerals.

Ryder (1985) carefully reviewed all aspects of 15555. It is one of the most allocated and most studied samples

Mineralogical Mode of 15555								
0	Longhi et al. 1972	McGee et al. 1977	Heuer et al. 1972	Nord et al. 1973				
Olivine	12.1	5-12	20	20				
Pyroxene	52.4	52-65	40	40				
Plagioclase	30.4	25-30	35	35				
Opaques	2.7	5						
Mesostasis	2.3	0.2-0.4	5	5				
Silica		0.3-2						



Figure 2: Closeup photo of sawn surface illustrating texture and vuggy nature of 15555,838. *Sample is* 3 x 5 *inches. Photo* # S93-045961.

from the moon and it has often been used in public displays.

Petrography

Lunar sample 15555 is a coarse-grained, porphryritic rock with rounded olivine phenocrysts (1 mm) and subhedral zoned pyroxene phenocrysts (0.5-2 mm) set in a matrix of poikilitic plagioclase (up to 3 mm). Interstices between plagioclase megacrysts are filled with minor opaque minerals, silica, glass and pore space. Inclusions of small euhedral chromite crystals occur in olivine and pyroxene. Inclusions of olivine and pyroxene are found in plagioclase. Ni-Fe metal is rare. Small vugs are about 2-4 % (figure 2).

Dalton and Hollister (1974) determined the crystallization sequence of 15555 by carefully studying the mineral zoning. At 1 atmosphere, Kesson (1975) determined experimentally that olivine crystallized at 1283 deg.C, followed by spinel at 1227 deg.C, pyroxene at 1154 deg.C and plagioclase at 1138 deg.C.

Walker et al. (1977) and Taylor et al. (1977) determined the cooling rate of 15555 (5 degC/day) by modeling





Figure 3: Photomicrograph of thin section of 15555. Scale is 2.5 mm across.

the diffusion of Fe in olivine phenocrysts, while Bianco and Taylor (1977) determined a cooling rate (at time of olivine nucleation) of 12-24 degC/day from the number density of olivine crystals (grains/mm²).

Kesson (1975) and Walker et al. (1977) performed highpressure experiments on 15555 composition to obtain the pressure-temperature relation for multiply saturated phases. In this way, they obtained estimates of the



compiled by C Mever

Figure 4: Pyroxene and olivine compositions for 15555.



Figure 5: Chemical zoning of pyroxene in 15555 from Mason et al. (1972).

depth of origin for this composition of 240 km and 100-150 km respectively. However, it would be remarkably good fortune if a lunar basalt sample was representative of a true primary magma, because limited accumulation of olivine and/or opaques would greatly alter the liquid composition, and hence the phase diagram. Lunar basalts have very low viscosity. Never the less, the composition of 15555 (and or 15016) was chosen for experiments, because it was highest in Mg, and thus, most likely to be the primitive end member.

Mineralogy

Olivine: Bell and Mao (1972), Brown et al. (1972), Longhi et al. (1972), Walker et al. (1977) and Taylor et al (1977) studied the zoning in olivine phenocrysts. Dalton and Hollister (1974) reported two kinds of olivine; large (1mm), normal-zoned olivine phenocrysts with Fo₆₇₋₂₉ and small (0.1mm) euhedral inclusions in plagioclase with Fo₄₉₋₁₆.



POIKILITIC CRYSTALS Figure 6: Fe/Mg zoning in poililitic plagioclase in 15555 and 15065 (from Longhi et al. 1976).



Figure 7: Variation of minor elements in large plagioclase crystal in 15555 as determined by ion microprobe (Meyer et al. 1974).

Pyroxene: Pyroxene compositions are given in plots by Brown et al. (1972), Bence and Papike (1972) and Walker et al. (1977). Coexisting, intergrown augite and pigeonite zone to a common focus and then the outer portions of pyroxene crystals zone to be extremely Fe-rich (figure 4). Mason et al. (1972) presented a traverse of the zoning in a complex pyroxene in 15555 (figure 5). Boyd (1972) found pyroxene cores had sector-zoned mantles of more Ca-rich pyroxene. Heuer et al. (1972), Nord et al. (1973) and Papike et al. (1972) studied microscopic exsolution.

Plagioclase: The cores of large plagioclase crystals are relatively unzoned $(An_{94.91})$ but the rims approach An_{78} . Longhi et al. (1976) studied the change in FeO/MgO from center to rim (figure 6) and Meyer et al. (1974) studied zoning of trace elements in 15555 (figure 7). Schnetzler et al. (1973) and Brunfelt et al. (1972) determined the trace element content of plagioclase separates.



Figure 8: Composition of mare basalts, including 15555.

Spinel: Dalton and Hollister (1974) found that chromite inclusions in olivine had ulvöspinel overgrowths. The spinels in 15555 have also been studied by Haggerty et al. (1972) and El Goresy et al. (1976). Some chromite has been reduced to form exsolution of ilmenite plus Fe metal.

Silica: Heuer et al. (1972) identify silica found in the mesostasis as cristobalite.

Phase Y: Brown et al. (1972) and Peckett et al. (1972) reported a Zr-Ti-Fe phase ("phase Y") in the mesostasis which Andersen and Hinthorne (1973) were able to date by ion microprobe.

Chemistry

The chemical composition of 15555 is given in tables 1 and 2. The lack of agreement is due to the relatively large crystal size and the small amounts distributed for analysis (see complaints lodged by Mason et al. 1972 and Rhodes and Hubbard 1973). Even Ryder and Schuraytz (2001) found significant variation in 4 gram duplicate splits.

15555 is found to be typical of Apollo 15 basalts (figure 8). It is thought to be the primitive end member of the olivine-normative Apollo 15 suite (Chappell and Green 1973), and therefore suitable for high-pressure melting experiments. The rare-earth-element pattern is flat (figure 9).

Radiogenic age dating

15555 proved to be a good sample to resolve analytical techniques and inter-laboratory comparisons and was apparently allocated (by CAPTEM) to each laboratory



Figure 9: Normalized rare-earth-element diagram for 15555 (data from table 1a,b,c). Note the variation.

for that purpose. Argon 39/40 plateau ages were obtained by Alexander et al. (1971), Husain et al. (1971), Podosek et al. (1971) and York (1972). The most dependable ages were those obtained on plagioclase separates (figure 10).

Chappell et al. (1971), Wasserburg and Papanastassiou (1971), Murthy et al. (1971), Cliff et al. (1972) and Birck et al. (1975) determined internal mineral isochrons by the Rb/Sr method (figures 11-14). These results are discussed in Papanastassiou and Wasserburg (1973).

Tatsumoto et al. (1972) found that U/Pb data for 15555 lie on a discordia line from 4.65 to 3.3 b.y., while Tera and Wasserburg found that the whole rock data lie on a discordia line from 3.3 and 4.42 (*magic point*, see figure 15). Andersen and Hinthorne (1973) dated Urich, Y-Zr phases by ion microprobe.

Lugmair (1975) and Unruh et al. (1984) present Sm-Nd and Lu-Hf whole rock data and Nyquist et al. (1991) have dated 15555 by Sm-Nd internal isochron (figure 16). Nyquist et al. also heated the sample to 790 deg C and 990 deg C for 170 hours to see what disturbance there was to age dating (see table).

Lee et al. (1977) reported Hf/W and ¹⁸²W/¹⁸⁴W.

It seems clear that this sample should be dated every few years, by whatever new technique, or new laboratory, that comes along, and that the data needs to be compared by CAPTEM, to understand precision

Table 1a. Chemical composition of 15555.

reference weight	Chappe	72	Schnetzler	72	Brunfelt	72	Ganapathy	73	Rhodes	73	Cuttitta 73	6	Fruchter	73	Janghorb	ani 73 1 8 a	
SiO2 % TiO2 Al2O3 FeO MnO MgO CaO	43.82 2.63 7.45 24.58 0.32 10.96 9.22	 (a) (a) (a) (a) (a) (a) (a) 	45 1.6 9.37 21.18 0.26 12.22 9.25	(c) (c) (c) (c) (c) (c) (c)	2.03 22.13 0.3 10.35	(d) (d) (d)			44.24 2.26 8.48 22.47 0.29 11.19 9.45	 (a) (a) (a) (a) (a) (a) (a) 	45.21 1.73 10.32 20.16 0.25 11.2 9.96	(f) (f) (f) (f) (f) (f) (f)	2.25 8.5 23.16	(d) (d) (d)	45.14 2.33 9.45 20.97 0.26 11.77	44.1 2.86 8.5 22.4 0.27 10.11	(d) (d) (d) (d) (d) (d)
Na2O K2O P2O5 S % sum	0.24 0.04 0.07 0.06	(a) (a) (a) (a)	0.26 0.03 0.07	(c) (c) (c)	0.28	(d)			0.24 0.03 0.06 0.05	(a) (a) (a) (a)	0.35 0.05 0.05	(f) (f) (f)	0.26	(d)	0.39	0.39	(d)
Sc ppm V					38.4 244	(d)					40 240	(f) (f)	40	(d)		38	(d)
Cr Co Ni Cu Zn	4036	(a)	3216	(c)	4820 61.8 90 6.6 1.3	(d) (d) (d) (d) (d) (d)	0.78	(e)	42	(a)	4516 87 96 0.13	(f) (f) (f) (f) (f)	4100 50	(d) (d)		3580 55	(d) (d)
Ga Ge ppb					2.9	(d)	8.5	(e)			4.6						
As Se Rb Sr Y	0.68 89.7	(b) (b)	0.445 84.4	(b) (b)	<0.05 0.085 0.75 84	(d) (d) (d) (d)	0.156 0.65	(e) (e)	0.6 92 23	(b) (b) (a)	1.1 93 23						
Zr Nb Mo Ru Rh Bd pph			57.3	(b)					76 4.3	(a) (a)	58 17						
Ag ppb Cd ppb In ppb					<7 2	(d) (d)	1 2.1 0.55	(e) (e) (e)									
Sh ppb Sb ppb Te ppb Cs ppm					0.026	(d)	0.067 3.4 0.03	(e) (e) (e)									
Ba La Ce Pr			32.2 8.06	(b) (b)	47 3.5 10	(d) (d) (d)					30		5.4	(d)			
Nd Sm Eu Gd			6.26 2.09 0.688 2.9	(b) (b) (b) (b)	3.2 0.75	(d) (d)							3.5 1.18	(d) (d)		1	(d)
Tb Dy			3.27	(b)	0.51 3.2	(d) (d)							0.7	(d)		0.92	(d)
Ho Er			1.7	(b)	0.78 2.7	(d) (d)											
Tm Yb Lu Hf			1.45	(b)	1.64 0.43 2.1	(d) (d) (d)					4.2		2.1 0.37 2.2	(d) (d) (d)			
la W ppb Re ppb Os ppb					0.29 1200	(d) (d)	0.0013	(e)								1.45	(d)
lr ppb Pt ppb					<0.1	(d)	0.006	(e)									
Au ppb Th ppm U ppm					0.48 0.3 0.14	(d) (d) (d)	0.139	(e)									

technique (a) XRF, (b) IDMS, (c) AA, colormetric, (d) INAA, (e) RNAA, (f) various, see paper

44.75 SiO2 % 44.22 (f) 45.86 (f) TiO2 2.36 2.07 (f) 2.4 (f) 7.54 AI2O3 8.29 8.67 (f) (f) FeO 24.24 23.4 (f) 23.45 (f) 0.3 MnO 0.29 (f) 11.48 11.55 MgO 11.11 (f) (f) CaO 9.18 9.14 (f) 9.24 (f) 0.24 Na2O 0.29 (f) 0.34 (f) K2O 0.04 0.05 (f) 0.039 (b) 0.09 (f) 0.042 (b) P2O5 0.05 (f) 0.06 S % 0.07 (f) 0.065 0.0855 (f) sum 49 Sc ppm (f) V 240 (f) Cr 4174 (f) Со 59 (f) 4700 (f) (f) (f) Ni 86 21 Cu Zn Ga Ge ppb As Se Rb 0.62 (b) 0.7 Sr 84 91 (b) 85.3 Y 25 140 130 Zr 124 (d) Nb Мо Ru Rh Pd ppb Ag ppb Cd ppb In ppb Sn ppb Sb ppb Te ppb Cs ppm 48 41.61 (b) Ва La Ce Pr Nd 7.518 (b) Sm 2.52 (b) Eu Gd Tb Dy Ho Er Tm 4.3 Yb 0.255 Lu (b) Ηf 2 (b) 3.2 3.26 (d) Та W ppb Re ppb Os ppb

technique (a) XRF, (b) IDMS, (c) AA, colormetric, (d) INAA, (e) RNAA, (f) various, see paper

Kaplan 76 Gibson 75 Chyi and Ehmann 73 Longhi 72

Murthy 72

(b)

(b)

Table 1b. Chemical composition of 15555.

reference Maxwell 72 Mason 72 Unruh 84 Birck 75

0.5 g

weight

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Ir ppb Pt ppb Au ppb Th ppm U ppm

reference	Ryder 20	duplicate		Ryder 20	001		Ryder	2001		Chappel 7	3	
SiO2 % TiO2 Al2O3 FeO MnO MgO CaO	4.014 44.5 2.3 8.21 22.75 0.282 11.16 9.22	4.001 45 2.02 9.16 21.49 0.275 11.32 9.47	(a) (a) (a) (a) (a) (a) (a)	23	21.3	(b)	43.7 2.45 8.3 22.6 0.28 11.1 9.2	44.8 2.06 8.6 22 0.28 11.2 9.5	(c) (c) (c) (c) (c) (c) (c)	44.75 2.05 9.01 21.68 0.3 11.39 9.62	43.82 2.63 7.45 24.58 0.32 10.96 9.22	 (a) (a) (a) (a) (a) (a) (a)
Na2O K2O P2O5 S % sum	0.228 0.042 0.065	0.234 0.036 0.053	(a) (a) (a)	0.24	0.259	(b)	0.22 0.04 0.1	0.26 0.04 0.1	(c) (c) (c)	0.27 0.04 0.06 0.04	0.24 0.04 0.07 0.06	(a) (a) (a) (a)
Sc ppm V				41.4	39.1	(b)						
Cr Co Ni	4592 62	4620 67	(a)	4600 57.4	4460 55.4 62	(b) (b) (b)	4387	3451	(c)	3968	4037	(a)
Cu Zn	6	3	(a) (a)	04	02	(0)						
Ga Ge ppb As										2.7		(a)
Se Rb Sr V	4 90 23	3 90 20	(a) (a)	109	99	(b)				0.54 92.2 18	0.76 90.7	(a) (a) (a)
Zr Nb Mo Ru Rh Pd ppb Ag ppb Cd ppb In ppb Sn ppb Sb ppb Sb ppb Te ppb	88 12	20 70 8	(a) (a) (a)							69 5		(a) (a) (a)
Cs ppm Ba La Ce				47 5.14 13.8	39 3.88 11.6	(b) (b) (b)						
Pr Nd Sm Eu				9 3.64 0.86	8 2.78 0.78	(b) (b) (b)						
Ga Tb Dy Ho Fr				0.77	0.6	(b)						
Tm Yb Lu Hf Ta W ppb Re ppb Os ppb				2.28 0.31 2.8 0.4	1.77 0.25 2.03 0.29	(b) (b) (b) (b)						
Ir ppb Pt ppb Au ppb Th ppm U ppm				0.41	0.29	(b)						

Table 1c. Chemical composition of 15555.

technique (a) XRF, (b) INAA, (c) fused bead, elec. probe

Table 2. Additional trace element data for 15555									
	Rb ppm	Sr ppm	U ppm	Th ppm	Κ%				
Chappell et al. 1971	0.68	89.7							
	0.72	91.7							
Murthy et al. 1971	0.7	85.32							
	0.538	74.11							
Tatsumoto et al. 1972	0.874	92			0.0538				
			0.1264	0.4596					
			0.1173	0.4296					
Mark et al. 1973	0.675				0.0313				
Compston et al. 1972	0.63	89.9							

Other Studies on 15555

Boyd 1972 Bell and Mao 1972 Michel-Levy and Johann 1973 Nord et al. 1973 Heuer et al. 1972 Crawford 1973 Czank et al. 1973 Wenk et al. 1973 Wenk and Wild 1973 Meyer et al. 1974 Blank et al. 1982 Brunfelt et al. 1973 Roedder and Weiblen 1972 Weeks 1972 Burns et al. 1972, 1973 Huffman et al. 1972, 1974, 1975 Simmons et al. 1975 Cukierman et al. 1973 Mark et al. 1973 Husain 1974 Friedman et al 1972 Eisenstraut et al. 1972 Gibson et al 1975 Kaplan et al. 1976 DesMarais et al. 1978 Allen et al 1973 Rosholt 1974 Fleischer et al. 1973 Megrue 1973 Fireman et al. 1972 Collinson et al. 1972, 1973 Pearce et al. 1972, 1973 Dunn and Fuller 1972 Hargraves et al 1972 Nagata et al. 1972, 1973 Schwerer and Nagata 1976 Chung and Westphal 1973 Schwerer et al. 1973, 1974 Schwerer et al. 1973 Tittmann et al. 1972 Warren et al 1973 Chung 1973 Hemingway et al. 1973 Adams and McCord 1972 Charrette and Adams 1975 Brito et al. 1973 Cukiermann et al. 1973 Cukiermann and Uhlmann 1974

topic pyroxene zoning olivine zoning petrogaphy HTEM, microstructure microstructure plagioclase crystallographic details, plagioclase crystallographic details, plagioclase crystallographic details, plagioclase ion microprobe, plagioclase proton microprobe, opaques trace element composition, plagioclase, pyroxene immiscible melt inclusions Mossbauer spectra microscopic spectra Mossbauer spectra microcracks recrystallization age dating age dating Pyrolysis, H, C isotopes GC Combustion Combustion, S, C isotopes Combustion, C isotopes INAA, Pb etc. Th isotopes tracks laser probe, rare gases solar wind rare gas magnetic data magnetic data magnetic data magnetic data magnetic data magnetic data dielectric data electrical conductivity Mossbaurer spectra seismic wave velocity seismic wave velocity, pressure seismic wave velocity, pressure specific heat reflectance spectra reflectance spectra thermoluminescence studies viscosity

viscosity



Figure 10: Argon 39/40 plateau ages for plagioclase and whole rock 15555 from Podosek et al. (1972).



Figure 12: Rb/Sr internal mineral isochron for 15555 (from Chappell et al. 1971).



Figure 14: Rb/Sr internal mineral isochron for 15555 (from Birck et al. 1975).



Figure 11: Rb/Sr internal mineral isochron for 15555 from Wasserburg and Papanastassiou (1971).



Figure 13: Rb/Sr internal mineral isochron for 15555 (from Murthy et al. 1971).

and accuracy that are claimed. In this way, 15555, becomes a sort of control sample for geochronology.

Cosmogenic isotopes and exposure ages

Marti and Lightner (1971) determined the cosmic ray exposure age as 81 m.y. by ⁸¹Kr . Podosek et al. (1971) and York et al. (1972) determined exposure ages of 90 m.y. and 76 m.y., respectively, by ³⁸Ar. This age in not associated with any local crater (Arvidson et al. 1975).

The lunar orientation and history of surface exposure is not well documented for 15555 and it has not often been allocated for depth profiles of cosmic ray, solar flare radionuclide studies *(however, see Fireman below)*. The large impact (figure 1) would have caused the rock to jump or role! Behrmann et al. (1972)



Figure 15: U-Pb data for leaches and "whole rock" splits. Line is drawn thru whole rock data and intersection at 3.3 (the age determined by Rb-Sr) (from Tera and Wasserburg 1971).

determined a track age of 34 m.y. and calculated that the erosion rate was about 1 mm per m.y. On the other hand, Poupeau et al. (1972) also determined the track density and concluded that the sample had been buried beneath the regolith. Bhandari et al. (1972) determined the track density (and suntan age).

Other Studies

Lunar sample 15555 was allocated for many other studies, including physical properties, spectroscopy, thermoluminescence, isotopic analysis (C, S, Th, etc. see table). Supplemental data was collected on companion samples.

Marti and Lightner (1971), Mergue (1973) and Husain (1974) reported the content and isotopic ratios for rare



Figure 16: Sm-Nd internal mineral isochron for 15555 from Nyquist et al. (1991). Linear arrays are also determined for heat treated samples, but with lower ages!

gasses in 15555. Fireman et al. (1972) used measurements of ³H (tritium), ³⁷Ar and ³⁹Ar from different depths in 15555 (and other Apollo samples) to determine the intensity of recent and long term solar flare activity.

Numerous experimental studies have been carried out on 15555 powder and/or synthetic mix *(but what composition should be used?)*. Humphries et al. (1972), Longhi et al. (1972), Kesson (1975) and Walker et al. (1977) determined the mineral phases present at various temperatures and pressures (figures 17 and 18). If, and only if, the composition is correct and the source region contains olivine, pyroxene and plagioclase, then the depth of origin (~200 km) can be concluded from these phase diagrams.

Summary of Age D	ata for 15555				
	Rb/Sr	Ar/Ar	U/Pb	Pb/Pb	Sm/Nd
Chappell et al. 1971	3.54 ± 0.13 b.y.				
Wasserburg, Pap 1971	3.32 ± 0.06				
Alexander et al. 1971		3.33 ± 0.05			
Husain et al. 1971		3.28 ± 0.06			
Murthy et al. 1971	3.3 ± 0.08				
Podosek et al. 1971		3.22 ± 0.03			
plagioclase		3.31 ± 0.03			
York et al. 1972		3.31 ± 0.05			
Cliff et al. 1972	3.34				
Papanastassiou, W 1973	3.32 ± 0.04				
Birck et al. 1975	3.34 ± 0.09				
Tatsumoto et al. 1972			3.3 (and 4.65)		
Tera and Wasserburg 1974			3.3 (and 4.42)		
Andersen and Hinthorne 19	973			3.36 ± 6	0.06
				3.46 ± 6	0.09
Nyquist et al. (1991)					3.32 ± 0.04
(heated)					3.23 ± 0.02
Caution: These ages have	e not been undated	using new decay d	constants.		

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Figure 17: Experimental phase diagram for 15555 from Walker et al. (1977).



Figure 18: Experimental phase diagram for 15555 by Kesson (1975).

Processing

Two slabs, cut at right angles, were made from this rock for allocations (figures 19-21). Slab ,46 is illustrated in figure 21. Slab ,57 was cut from ,48 and used for many allocations.

This large rock has been used to prepare 13 lunar sample displays, one of which is illustrated in figure 22. These are located in Edmonton, Geneva, Oakland, Yorba Linda, Denver, Washington D.C., Illinois, Kansas, Boston, Michigan, Philadelphia, Austin and Utah. Three thin sections of 15555 are also on display.

List of NASA	photo #s for	15555
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S71-43390-43394	color, dusty
S71-43952-43954	dust free
871-51781-51795	TS
S71-52201	TS
\$71-52213	TS
S71-57110	after slab cut, B&W
S71-57987	exploded parts, slab
S74-23072	TS
874-31406-31413	,461 - ,463
875-33416-33421	,56
S79-27098-27100	set of thin sections
S85-29591-29600	,791 - ,463
\$90-37023	,160
893-45953-45962	,838
S96-09087	,880
S97-16866	,880



Figure 19: First saw cuts of 15555. Photo number S71-57110. Cube is 1 inch.



Figure 20: Sawing slab. Photo # S71-57094.



Figure 21: Parts diagram for slab 15555,46. *Photo* # *S*71-57987. *Cube is 1 inch.*



Figure 22: Display case with 15555,160. Case is made from optical glass and filled with dry nitrogen.

Figure 23: Location of samples studied by Fireman et al. (1972) in second slab (,57) cut from 15555,45.

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