International Workshop on the Assimilation of Satellite-Observed Cloud and Precipitation Observations in NWP Models 1-3 May 2005

Rainfall modelling and assimilation into mesoscale NWP models S. Davolio

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BOLAM MODEL

- Hydrostatic, Limited Area Model
- Horizontal resolution 30 km 7 km
- Multiple Nesting over specific area
 - Parameterized deep convection (Kain-Fritsch)

MOLOCH MODEL

- Non hydrostatic, fully compressible
- Horizontal resolution 1-3 km
 - Convection explicity resolved



Data assimilation -OI (SYNOP and TEMP) -Development of EKF methods based on the dynamical stability of the system (A.Trevisan) -Nudging for rainfall assimilation

 The <u>Mesoscale Alpine Programme (MAP)</u>, which focused on the region around the Alps, has been (and will be) a good opportunity for investigate the mechanisms responsible for severe precipitation events using state-of-the-art models

- •The application of operational and experimental models at increasing resolution has allowed to better simulate small scale processes as:
 - orographic forcing
 - microphysics
 - convection



Results from many case studies of heavy precipitation events south of the Alps (in connection or not with MAP)

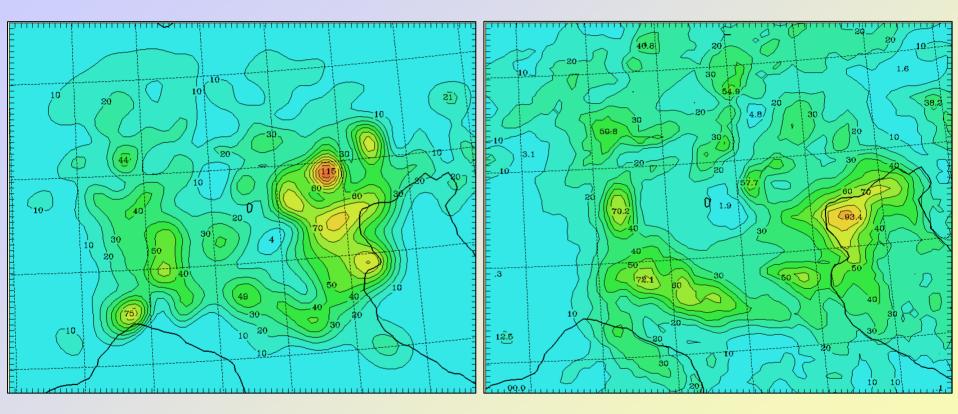


- Identification of "large-scale" precursor
- Identification of mesoscale *ingredients*:
 - pre-frontal low level jets;
 - orographic flow deviations (barrier winds) and convergence;
 - convective/precipitating bands over the sea.
- Identification of physical *processes* and *factors*:
 - orography
 - latent heat (condensation and evaporation of precipitation)
 - sensible and latent fluxes from the Mediterranaean sea, etc...



MAP-IOP15

24 hour acc. precipitation, 07 Nov 1999, 06 UTC Contour interval 10 mm



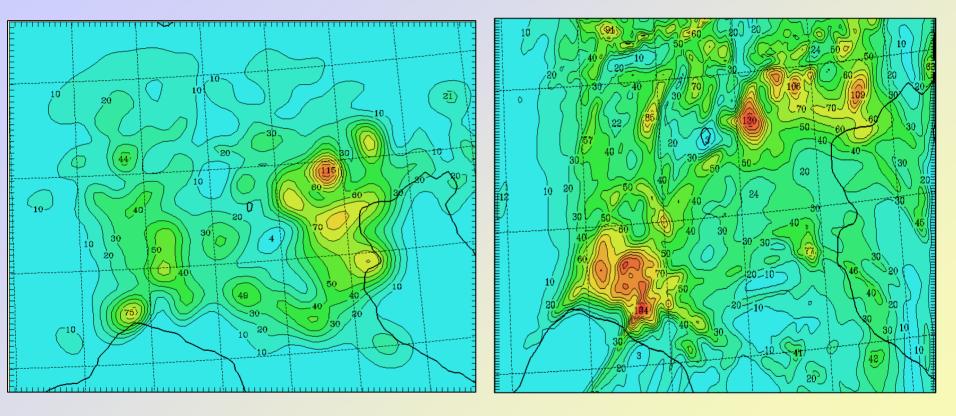
Observed

BOLAM (7km)



MAP-IOP15

24 hour acc. precipitation, 07 Nov 1999, 06 UTC Contour interval 10 mm



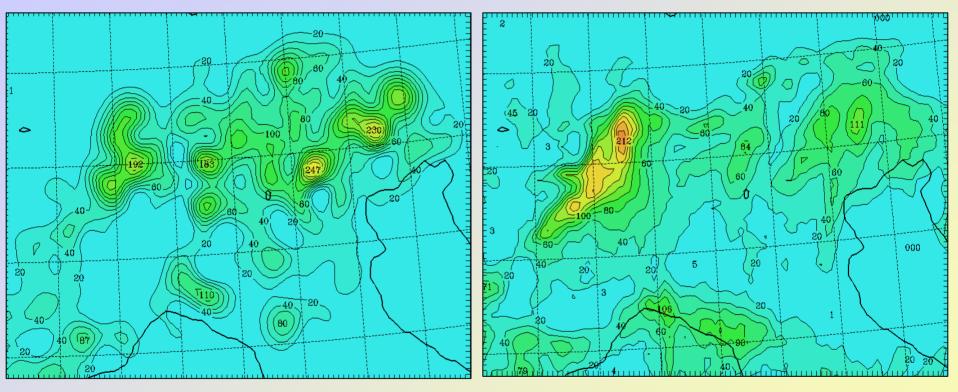
Observed

MOLOCH (2km)



MAP-IOP2b

24 hour acc. precipitation, 21 Sep 1999, 00 UTC Contour interval 10 mm



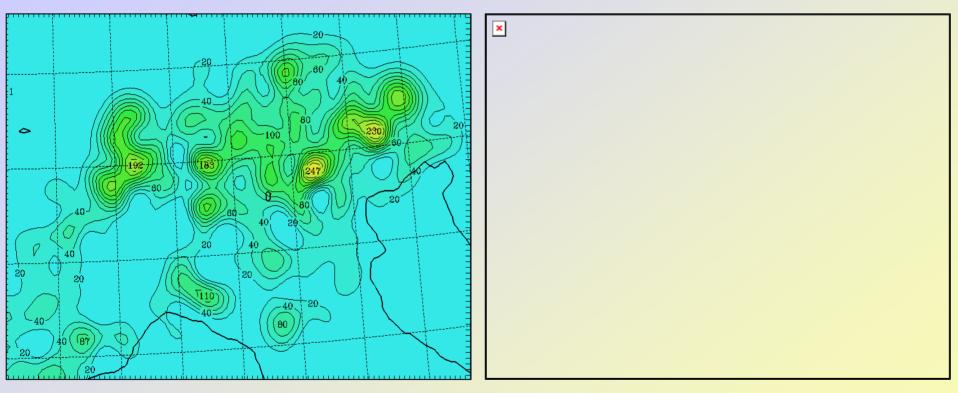
Observed

BOLAM (7km)



MAP-IOP2b

24 hour acc. precipitation, 21 Sep 1999, 00 UTC Contour interval 10 mm

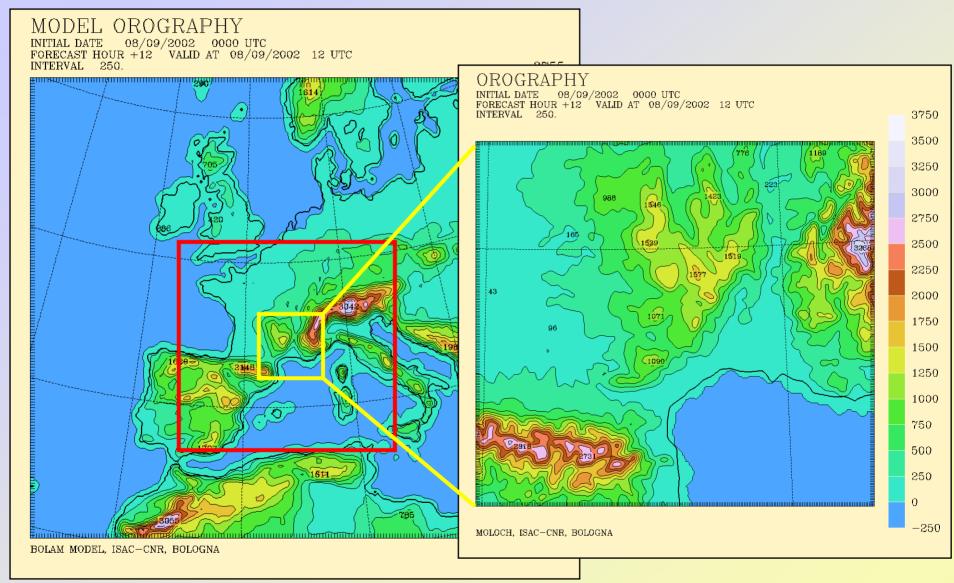


Observed

MOLOCH (2km)

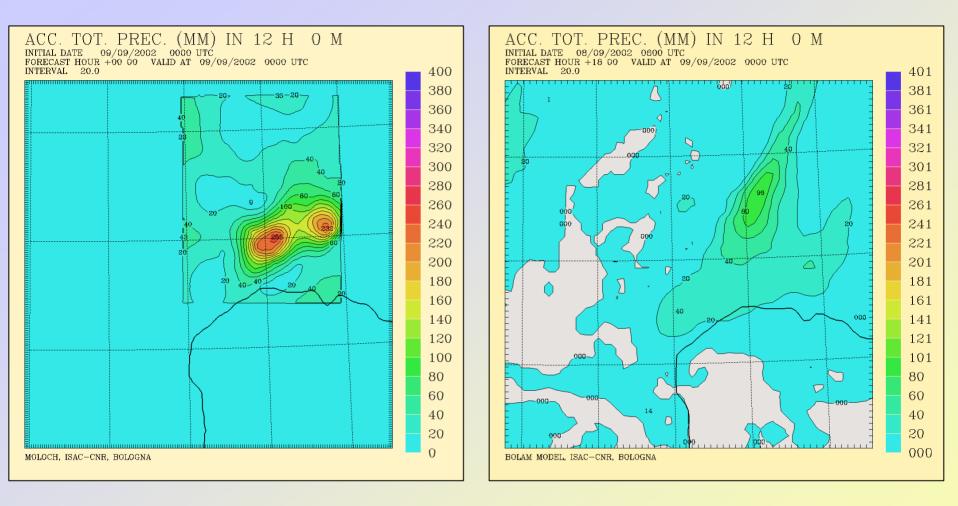


GARD EVENT: 8-9 September 2002





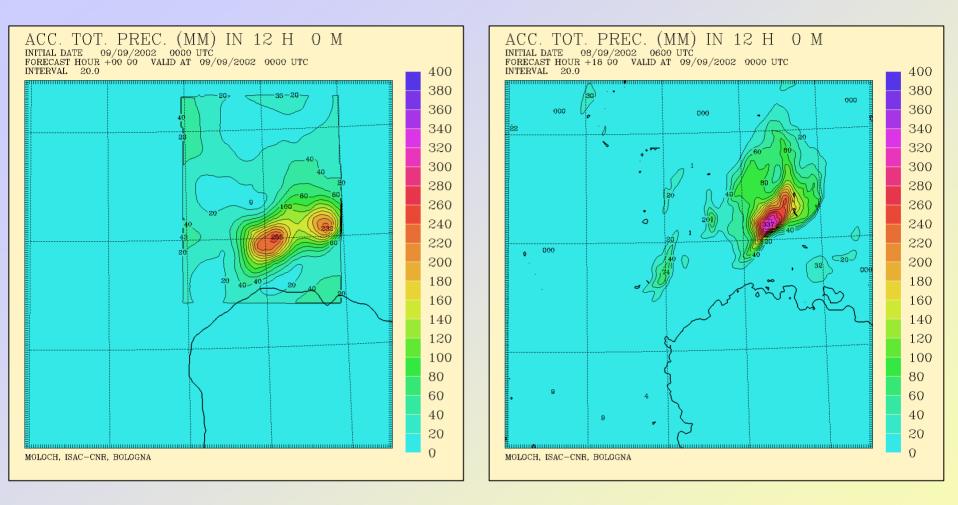
12 hour acc. precipitation, 9 Sep. 2002, 00 UTC



Convective instability is mainly released over the cost. The rainfall max, located over the mountains, is produced by large scale uplift, due to flow interaction with orography.



12 hour acc. precipitation, 9 Sep. 2002, 00 UTC



Trigger and rainfall location are now dissociated. The MCS develops over the coastal area and moves NE, becoming stationary at some distance from the main mountain slopes.



Rainfall Assimilation: the nudging scheme

After comparing RR_t and RR_f **maps** nudging of specific humidity profile

$$\frac{\partial \mathbf{q}(\mathbf{k})}{\partial t} = -\tau_{s,c}^{-1}(\mathbf{k}) \left\{ \mathbf{q}(\mathbf{k}) - \varepsilon_{s,c} \mathbf{q}^{*}(\mathbf{k}) \right\}$$

The forcing is a <u>function of the precipitation type</u> $\varepsilon_{s,c} = \text{over/under saturation coefficient}$ $\tau_{s,c} = \text{relaxation time (vertically modulated)}$

•Stratiform precipitation:

q is changed only in the middle-lower troposphere where large scale condensation takes place.

 $\mathsf{RR}_{\mathsf{f}} < \mathsf{RR}_{\mathsf{t}}$ $\mathsf{q}(\mathsf{k})$ $\mathcal{E}_{s}^{+}(k)q^{*}(k)$ $\mathcal{E}_{s}^{+}(k)=1.1$ $\mathsf{RR}_{\mathsf{f}} > \mathsf{RR}_{\mathsf{t}}$ $\mathsf{q}(\mathsf{k})$ $\mathcal{E}_{s}^{-}(k)q^{*}(k)$ $\mathcal{E}_{s}^{-}(k)=0.8$

•Convective precipitation:

q is changed only in the boundary layer.

 $\mathsf{RR}_{\mathsf{f}} < \mathsf{RR}_{\mathsf{f}}$ $\mathsf{q}(\mathsf{k})$ $\varepsilon_c^+(k)q^*(k)$ $\varepsilon_c^+(k) = 0.9$ $\mathsf{RR}_{\mathsf{f}} > \mathsf{RR}_{\mathsf{f}}$ $\mathsf{q}(\mathsf{k})$ $\varepsilon_c^-(k)q^*(k)$ $\varepsilon_c^-(k) = 0.1$



REMARKS

 Physically based scheme, applied to a mesoscale model with parameterized convection (dependence on the model)

- Coefficients selected after experimenting
- The assimilation scheme allows the assimilation of precipitation also when the rain is not purely convective
- More: Davolio and Buzzi, Wea. Forecasting, 2004.

EXPERIMENTS:

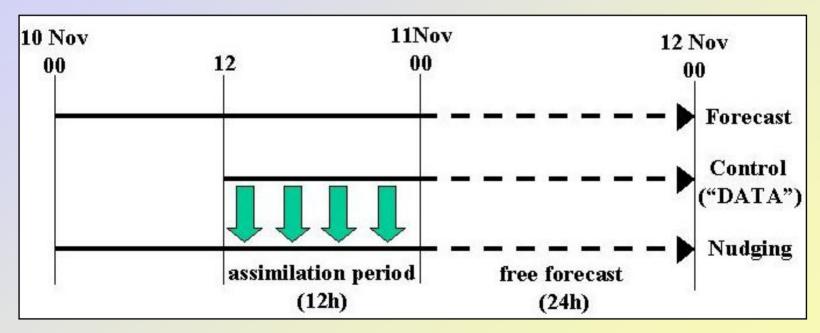
- real case, idealized data
- real case, real data
- idealized experiments



Real case, Idealized Data

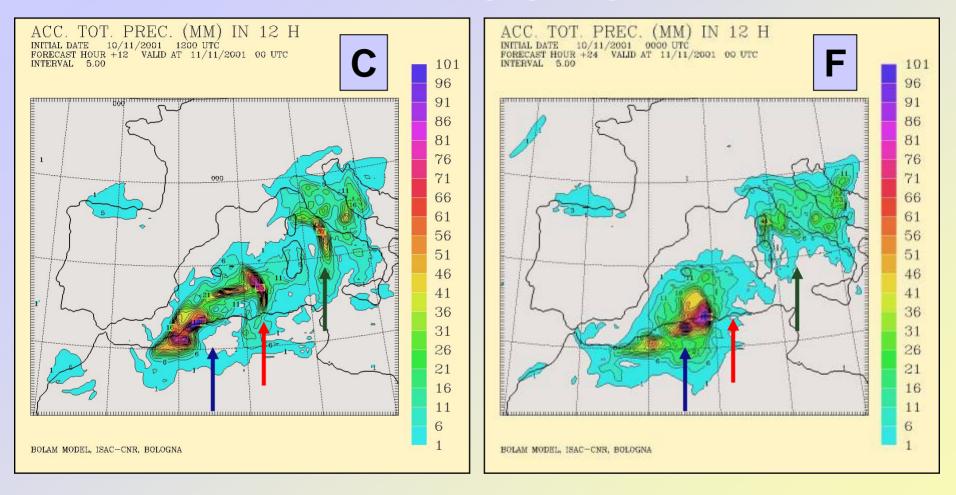
- METHOD: Lagged Forecast scheme (Algeria flood, Nov. 2001) Two different simulations from initial condition 12 hours apart:
 - *"Control Run"*: represents the reference state and provides the target rain rate.
 - *"Forecast Run"*: represents the real forecast to be improved.

Nudging procedure applied for 12 hours to a simulation starting from the same initial condition of the Forecast Run → (*Nudging Run*).





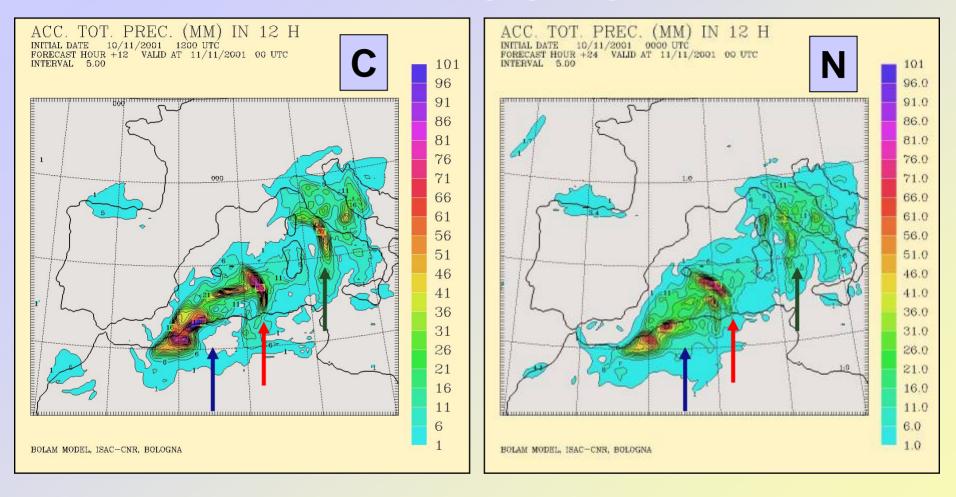
Results at the end of the nudging stage





- Different rainfall patterns over the coast and south of the Balearic Islands
- Rain band missing in the forecast run
 - Rain band and area of light rainfall around Sardinia

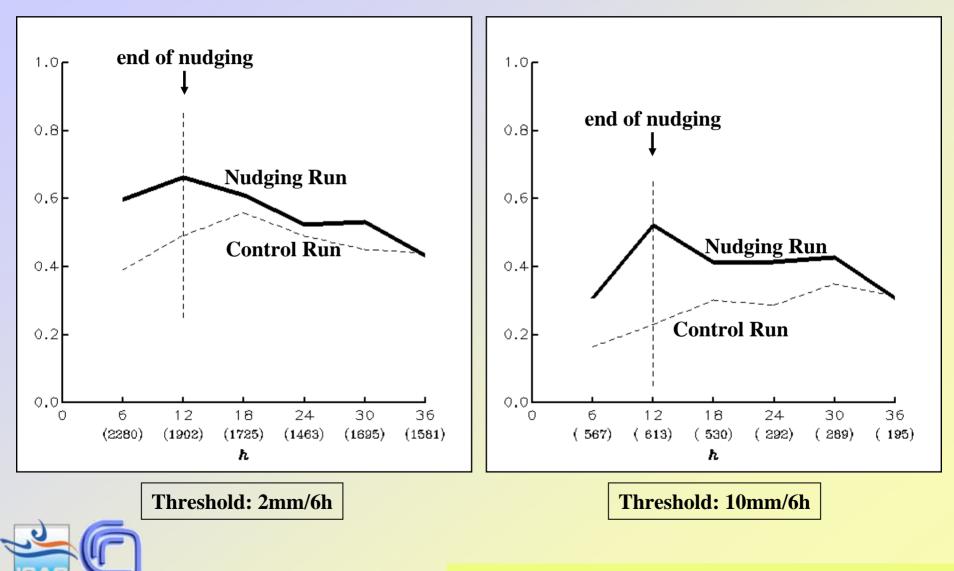
Results at the end of the nudging stage

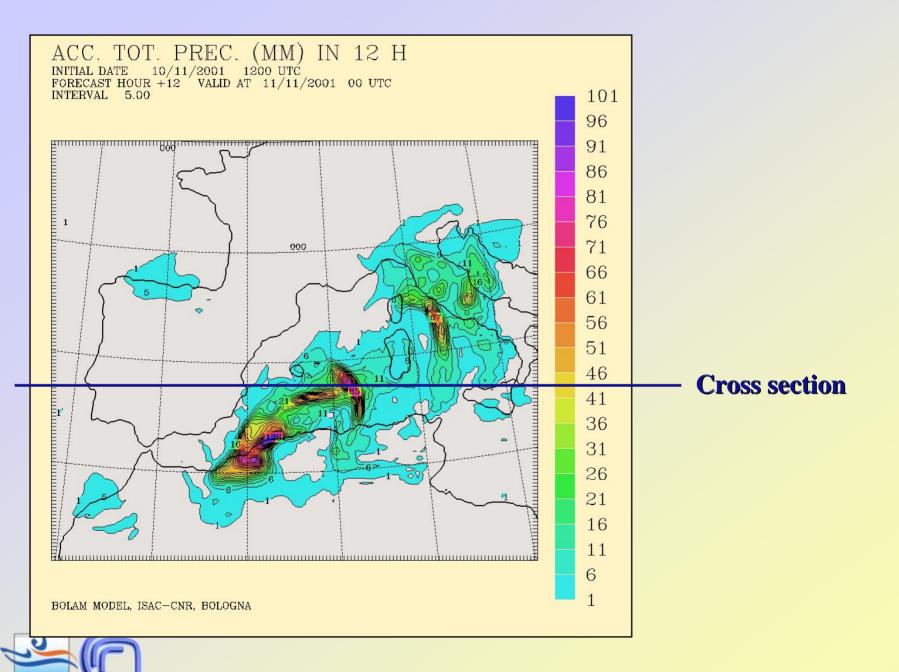


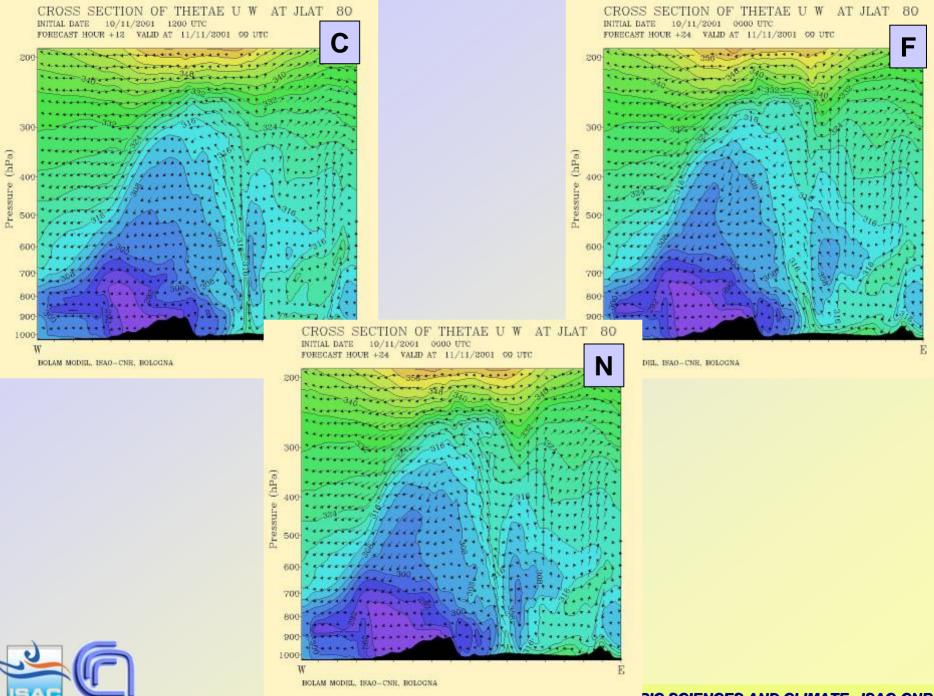


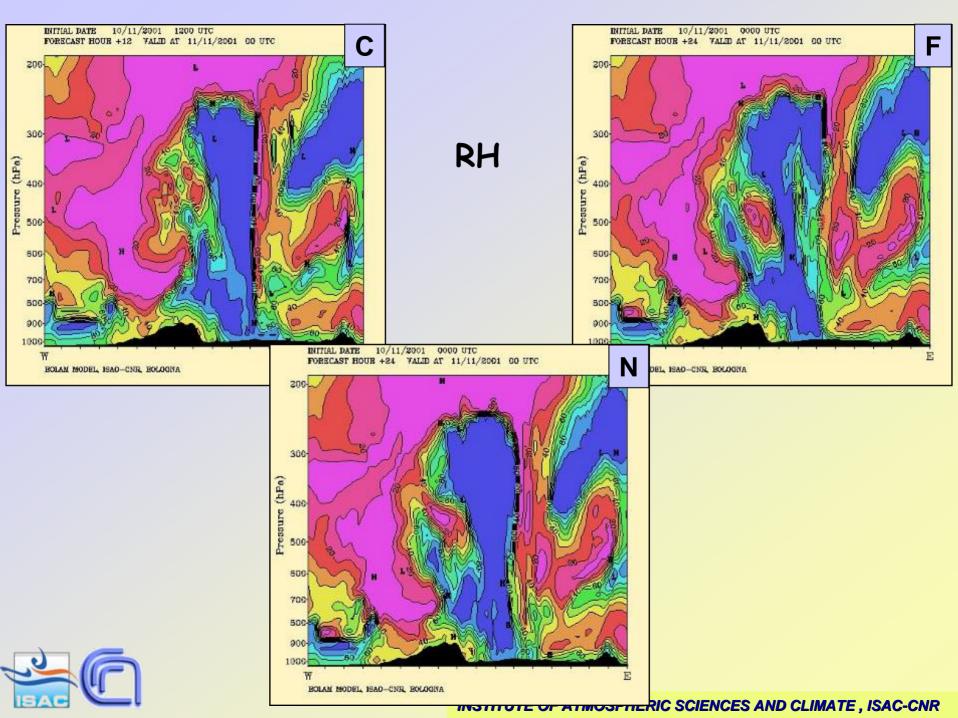
- Rain band slightly shifted westward but correct in intensity
- Rain band in phase but intensity too low (36 mm/12h instead of 67 mm/12h)

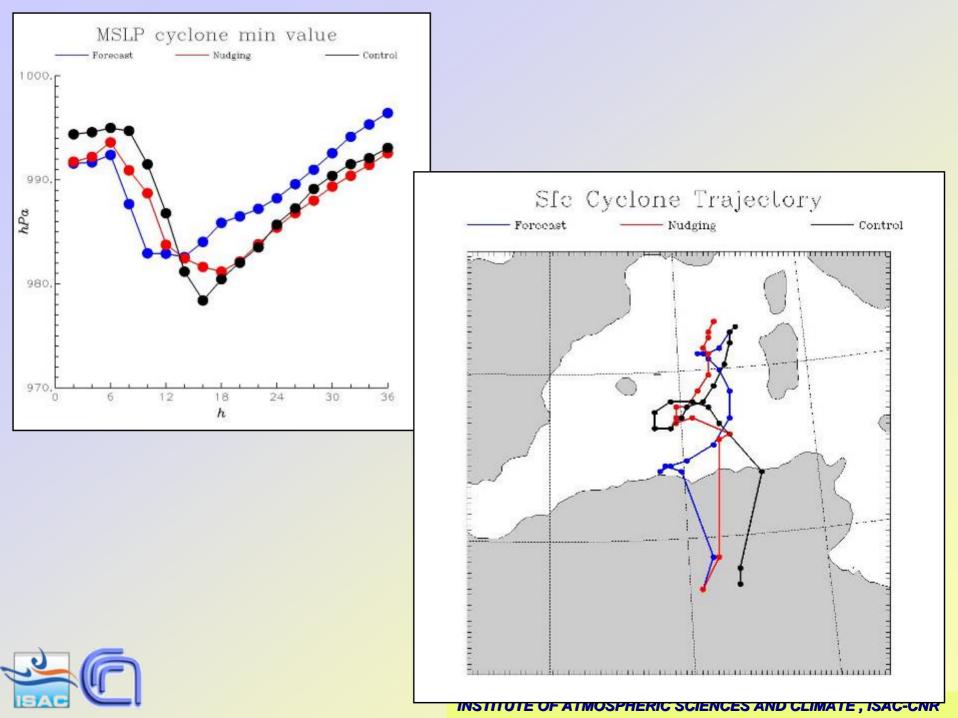
RESULTS after the nudging stage Equitable Threat Score vs simulation time











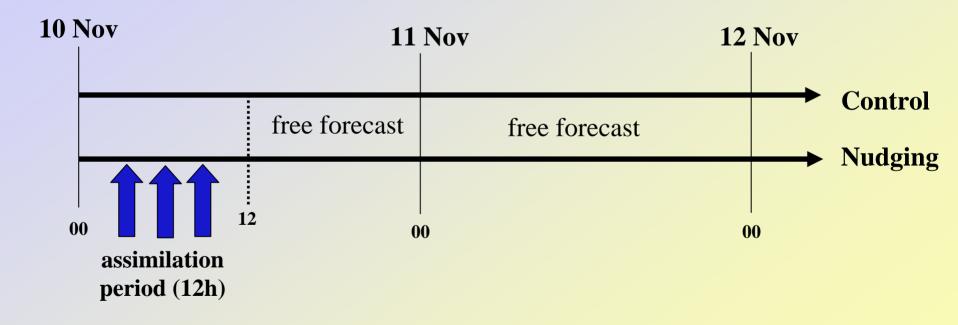
Real case, Real data: assimilation of satellite rainfall data

• 2-h accumulated satellite rainfall

 \succ (seems to be the best interval from sensitivity tests)

• Satellite data: IR/PMW

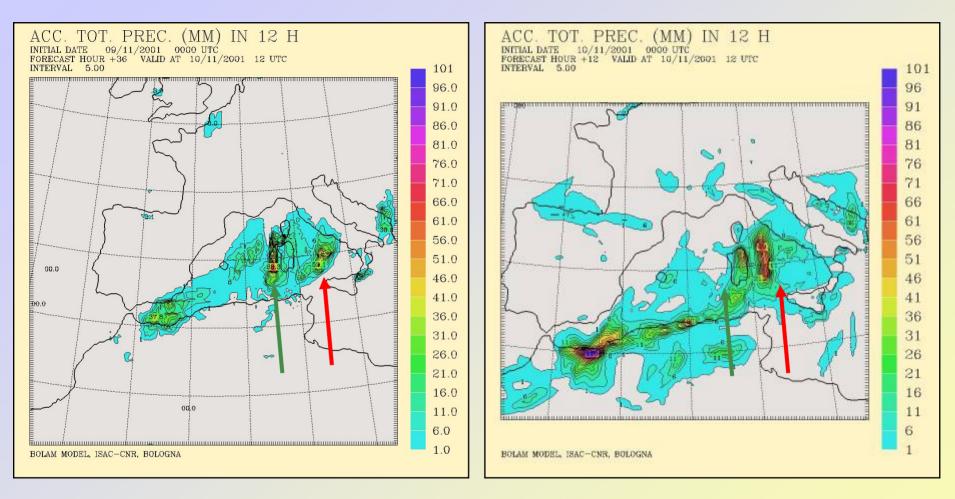
Assimilation ONLY OVER THE SEA





Satellite data

Control Run

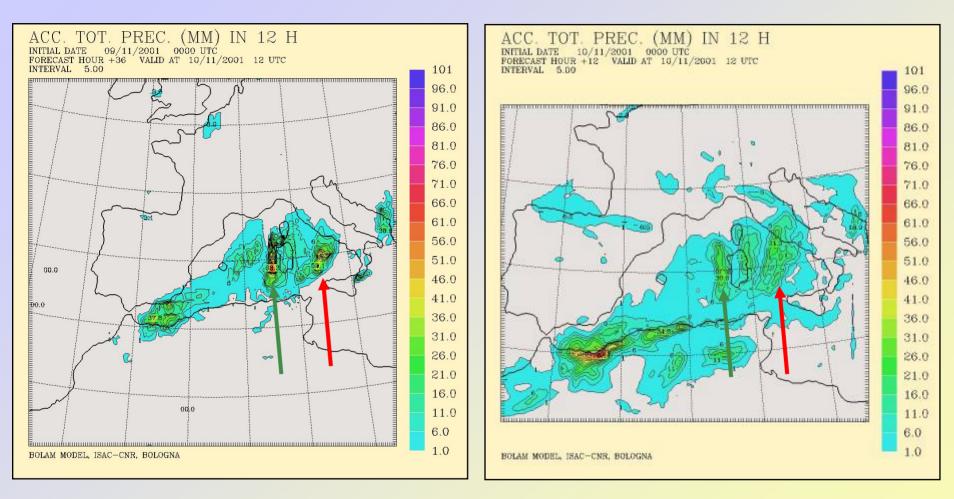


- Heavy precipitation west of Sardinia (satellite data).
- Area of rainfall over Tyrrenian Sea.



Satellite data

Nudging



Precipitation west of Sardinia (position ok, intensity 30 instead of 68 mm/12h).
Improvement over Tyrrenian Sea.



Idealized experiments:

Impact of rainfall assimilation on baroclinic wave development

Motivation:

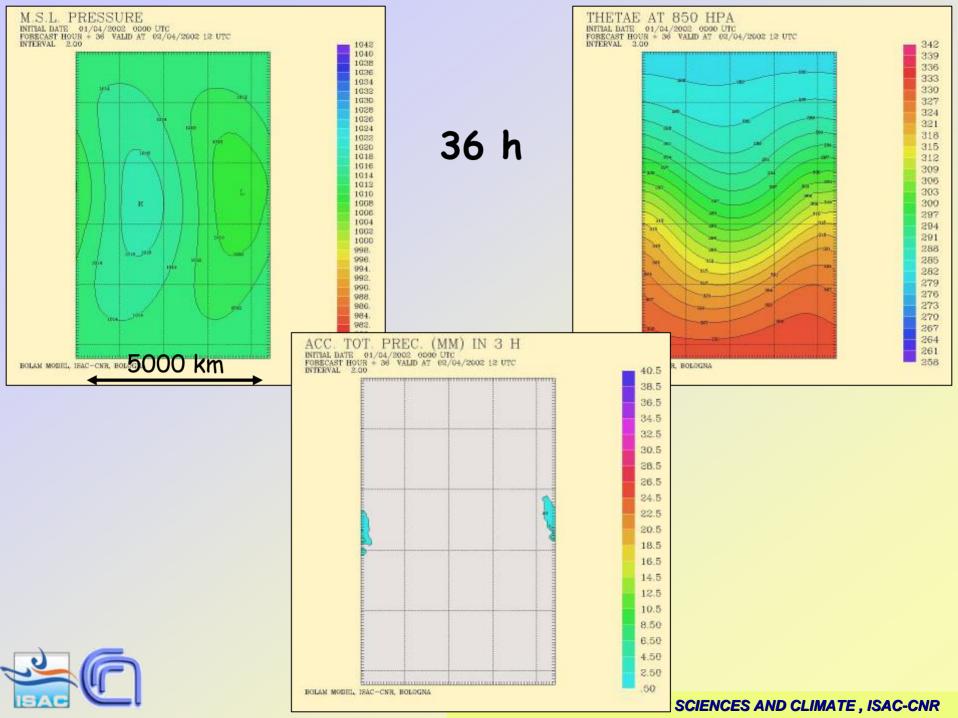
- Investigate the effect of rainfall assimilation using a conceptual/physical model of meteorological phenomena.
- GOOD TOOL to interpret the results obtained in the realistic framework.

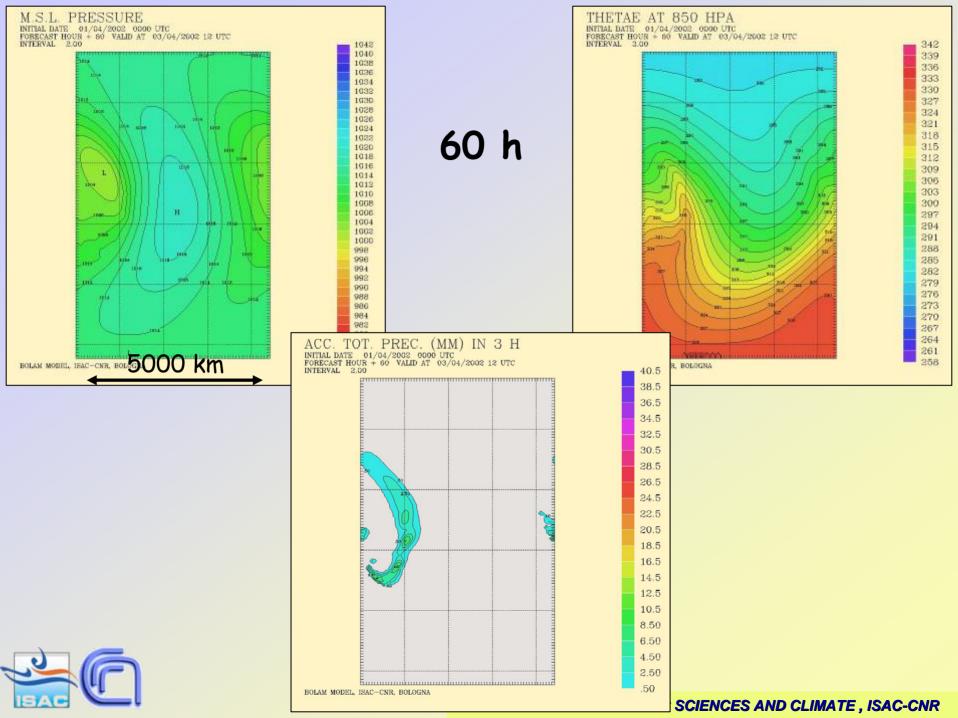
Model:

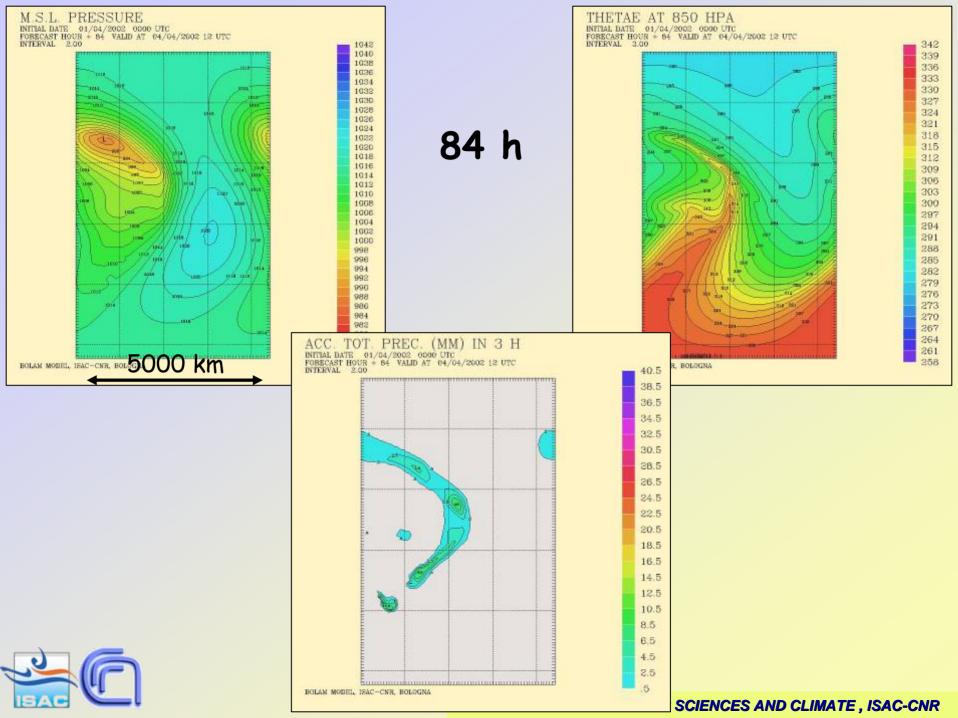
Bolam model in channel version, representing a midlatitude belt, hor.res. \approx 50km

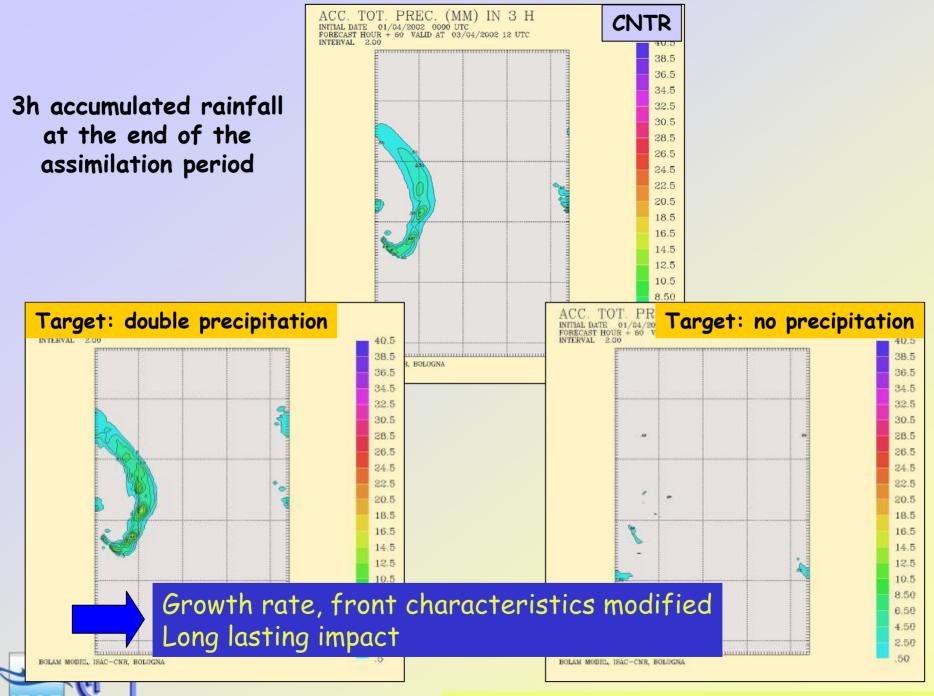
- Initial state: unstable baroclinic zonal jet. A baroclinic wave develops into large amplitude cyclones and anticyclones.
- 1) Nudging to investigate effects of intensity error
- 2) Nudging to investigate effects of phase error

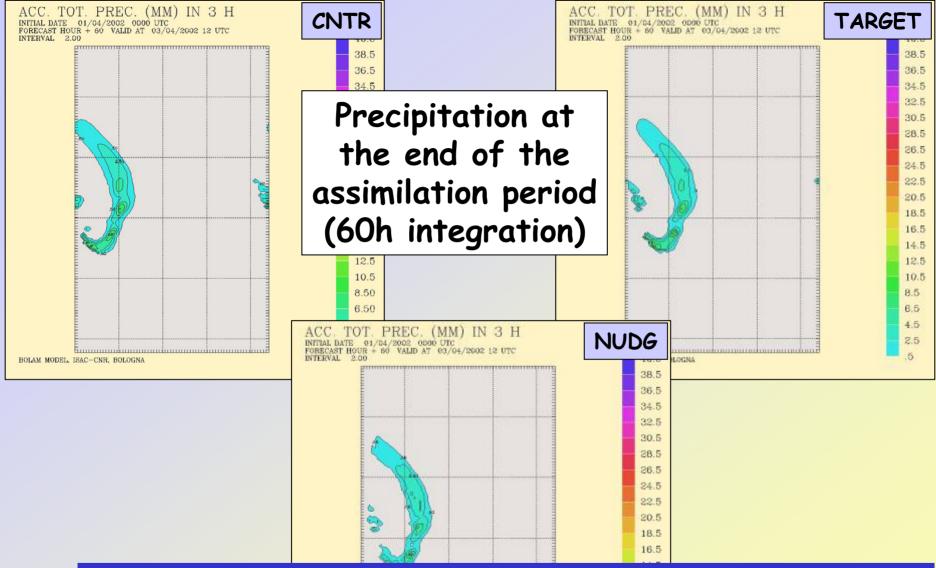












Precipitation pattern and frontal features (T, q, low level jet, cloud cover, area of upward vertical motion) corrected during the nudging, BUT the correction tends to disappear (as in the realistic experiments).

Conceptual model behaviour useful to explain real assimilation results.

• Intensity errors: local modification (rainfall area) affects the global evolution because the forcing (LH) amplifies (reduces) the large scale growing mode.

• Phase errors: local modification that imposes a forcing far from the structure of the growing mode is unable to modify the large scale evolution. For a small scale precipitation system associated with small scale instability (e.g. MCS) it may have impact, but not on large scale dynamical evolution.

In order to effectively modify the large scale, acting locally, it is necessary to localize areas where the leading instabilities are growing (e.g. bred vectors).



OPEN PROBLEMS AND PERSPECTIVE

 Assimilation of real satellite data shows positive impact even if limited to sea areas

- Problems in case of orographic precipitation
- Gard Case Study: assimilation of IR precipitation estimates:
 - data error
 - need for more complex scheme (advection)
- Application of assimilation schemes, for rainfall and clouds, to high resolution CRM
 - no partitioning problem
 - time scales convection/data
 - precipitation alone is not enough. What else? (Cloud information?)

