



# Climate Model Test Discussion An Observational Perspective

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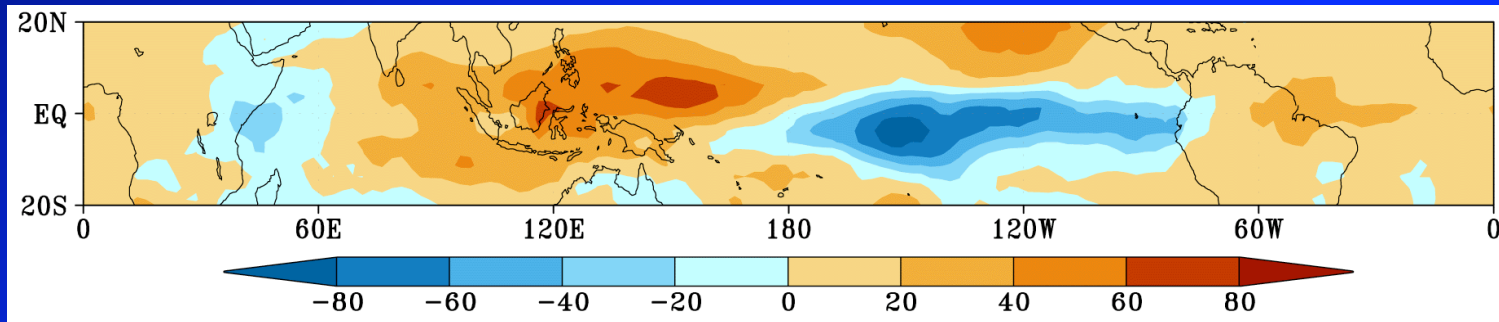
**Bruce Wielicki**  
**NASA Langley Research Center**

**CERES Science Team Meeting**  
**NCAR, March 29-31, 2004**

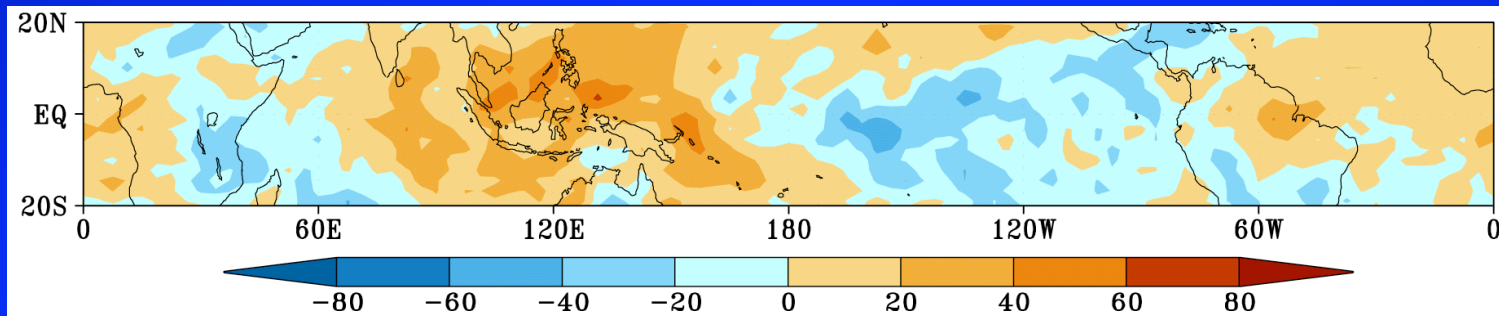


# Jan/Feb 98 El Nino TOA LW Flux Anomalies (relative to ERBE 1985-1989 average)

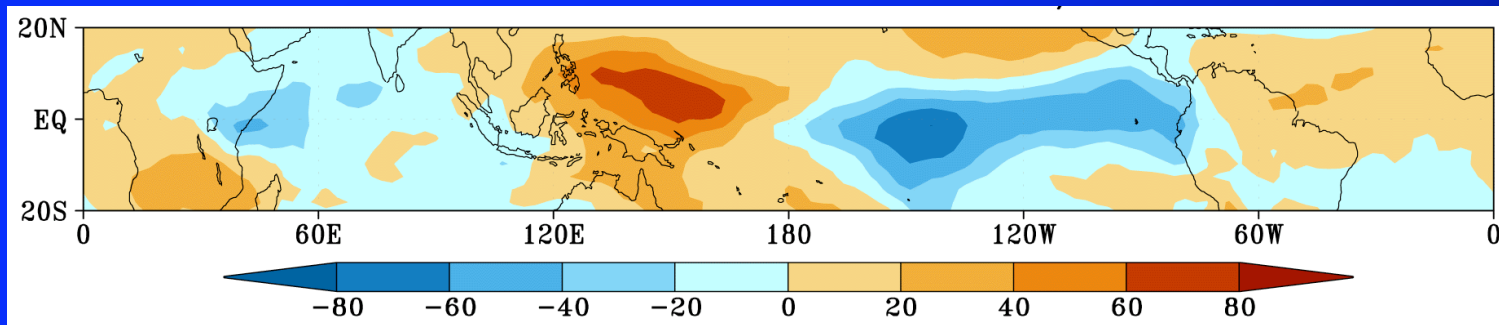
## CERES ERBE-Like LW Flux Observations



## NOAA GFDL Standard Climate Model

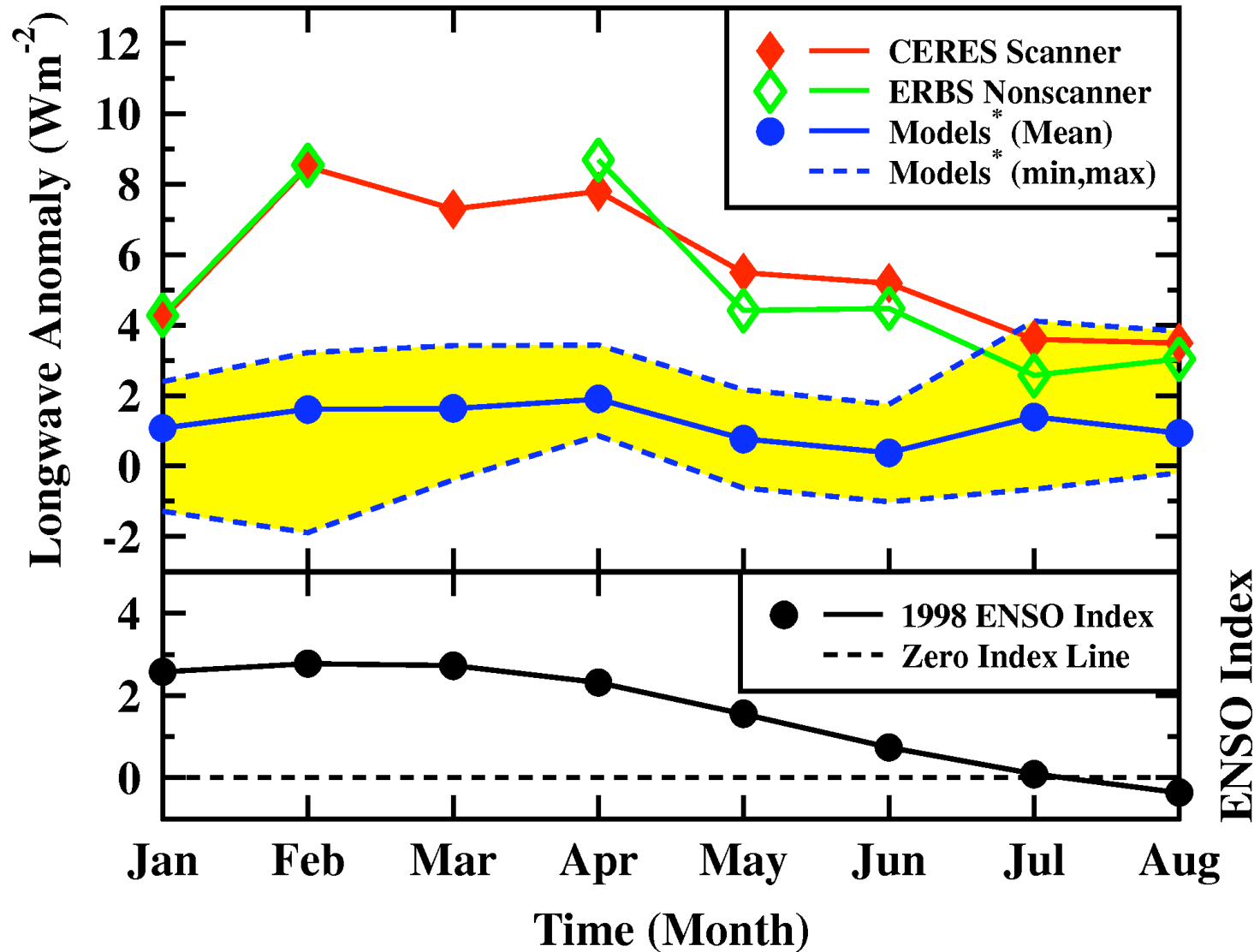


## NOAA GFDL Experimental Prediction Model



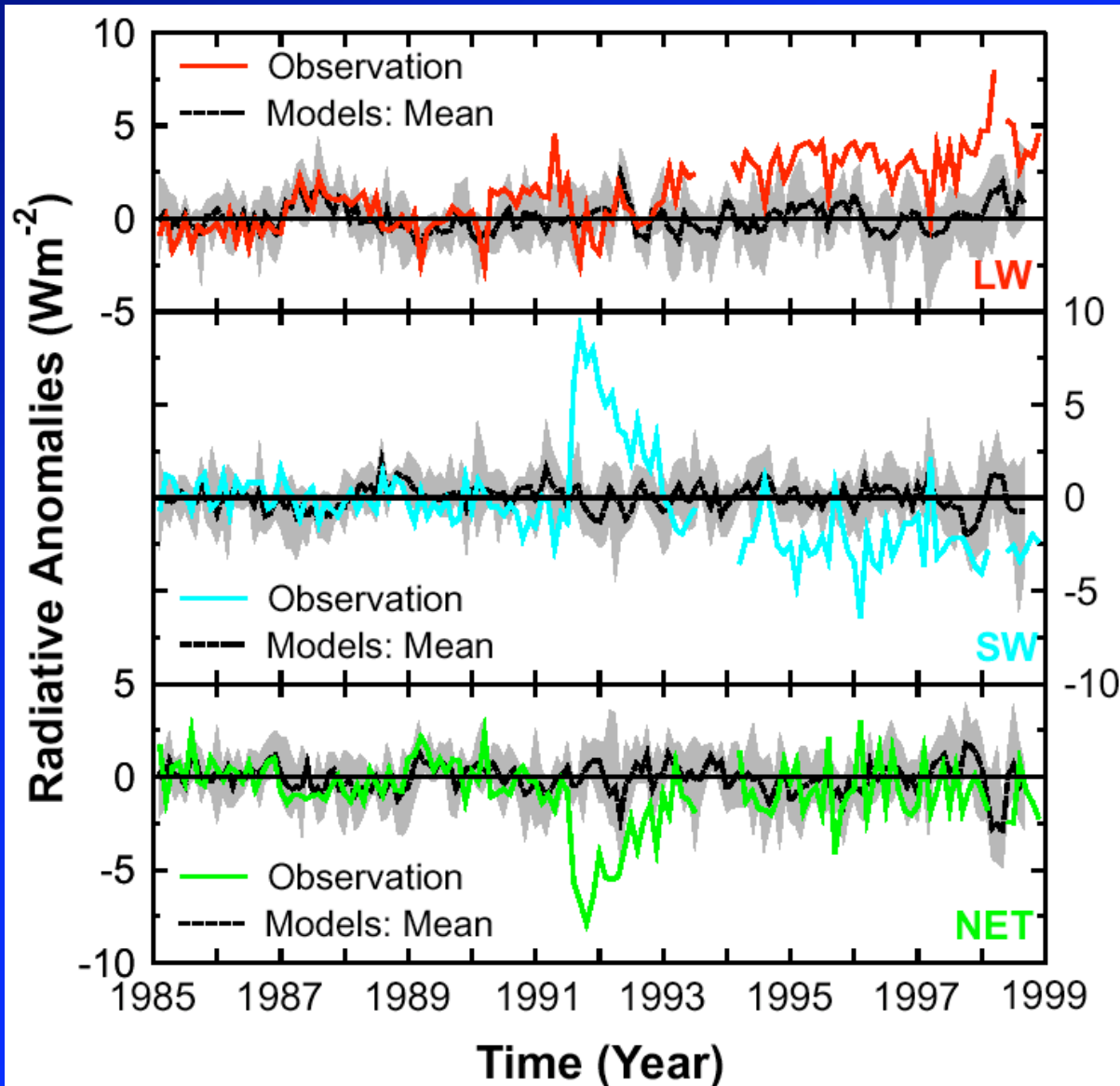
# 1998 El Nino Tropical Mean (20S - 20N) Longwave Flux Anomalies

(Anomalies Referenced to 1985 through 1989 Baseline)



\* 5 Climate Models and NCEP Re-analysis; All used observed SSTs; Climate Models: NCAR-CSM (Kiehl) UKMO (Allan, Slingo), GFDL and GFDL-EP (Soden, Gordon), CSU (Randall)

# Comparison of Observed Decadal Tropical Radiation Variation with Current Climate Models



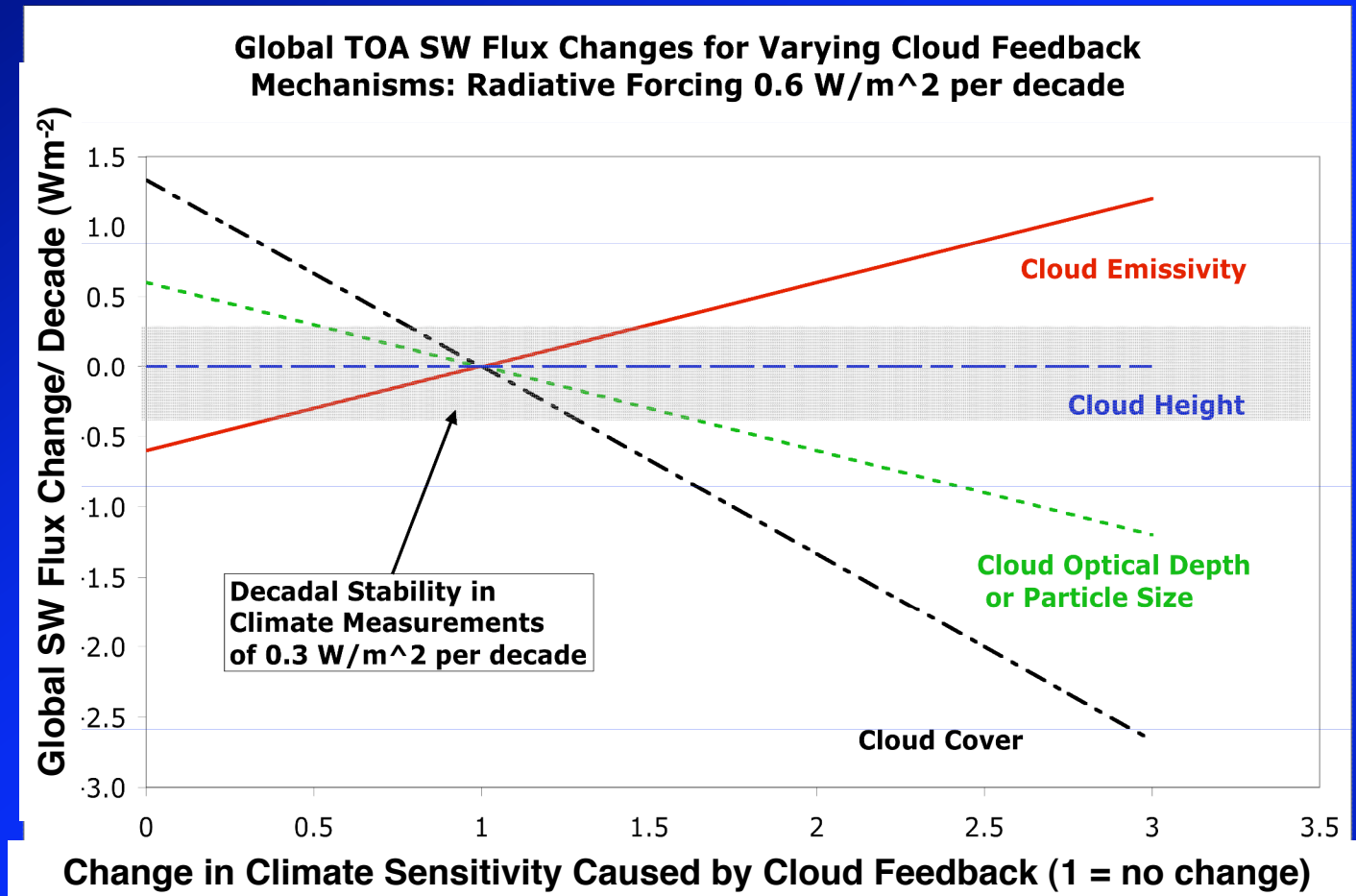
**LW:**  
Emitted Thermal  
Fluxes

**SW:**  
Reflected Solar  
Fluxes

**Net:**  
Net Radiative Fluxes

*Models less variable  
than the observations:*  
- missing feedbacks?  
- missing forcings?  
- clouds physics?

# How accurate to constrain equilibrium global cloud feedback?



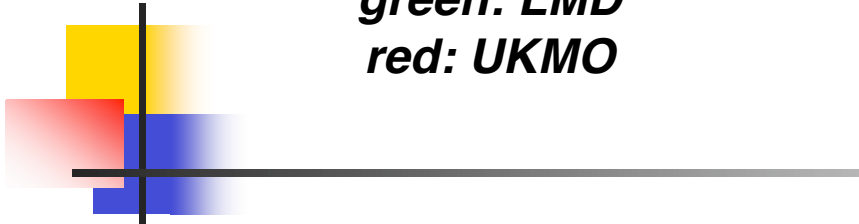
- Regional changes will be larger: but no regional “constraint” and global mean still must be accurately known for global feedback.
- UKMO ensemble climate noise for annual tropical mean SW and LW fluxes  $\sim 0.3 \text{ Wm}^{-2}$ : this might be a reasonable lower limit on accuracy.

*white: ERBE/NCEP, ERA, DAO*

*blue: ECMWF*

*green: LMD*

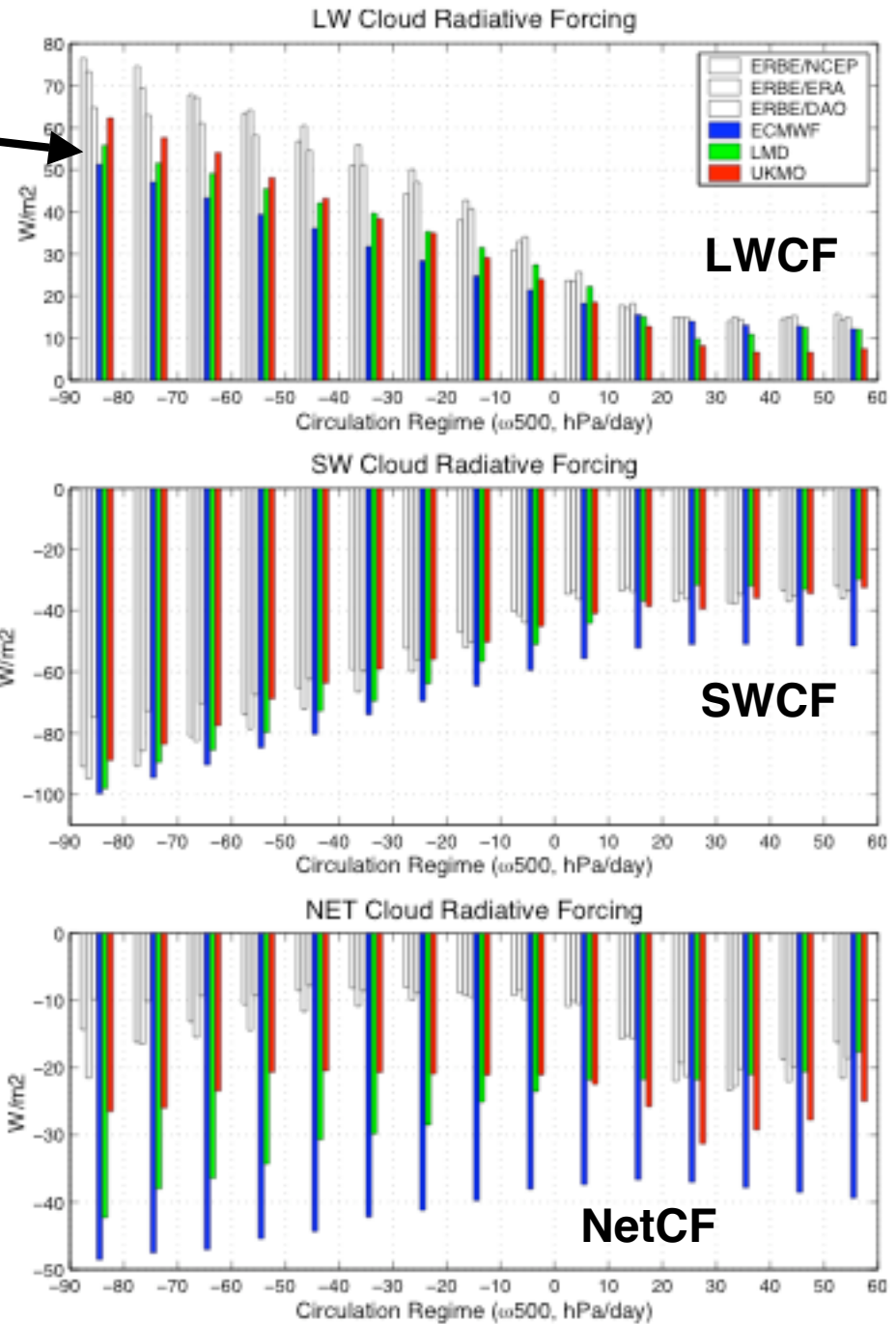
*red: UKMO*

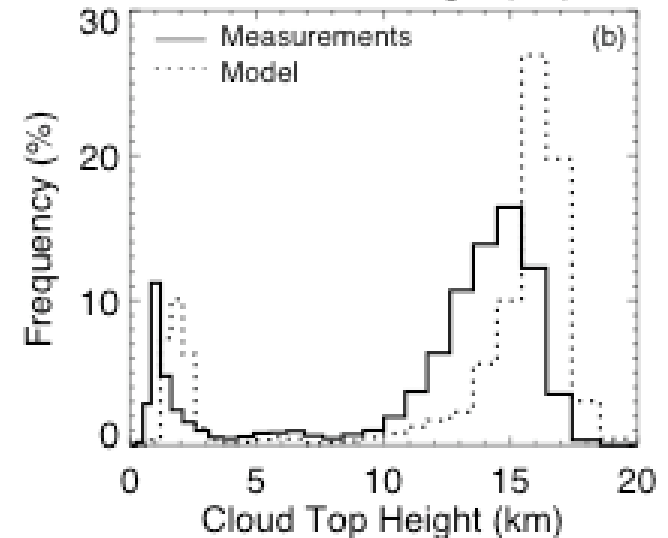
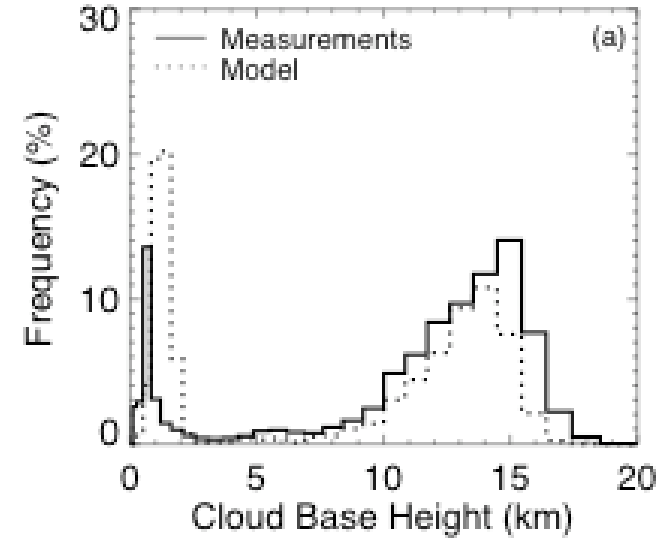
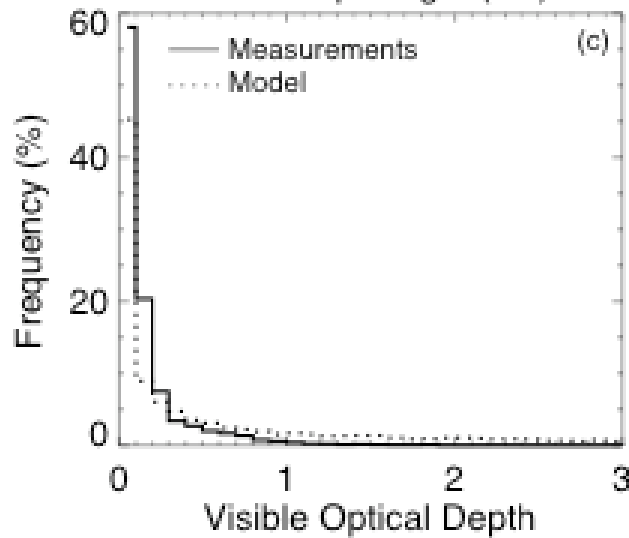
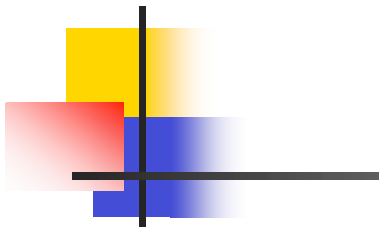


# Model vs Data Intercomparisons by Dynamic Regime:

## Vertical Velocity

(Bony et al., 2003)

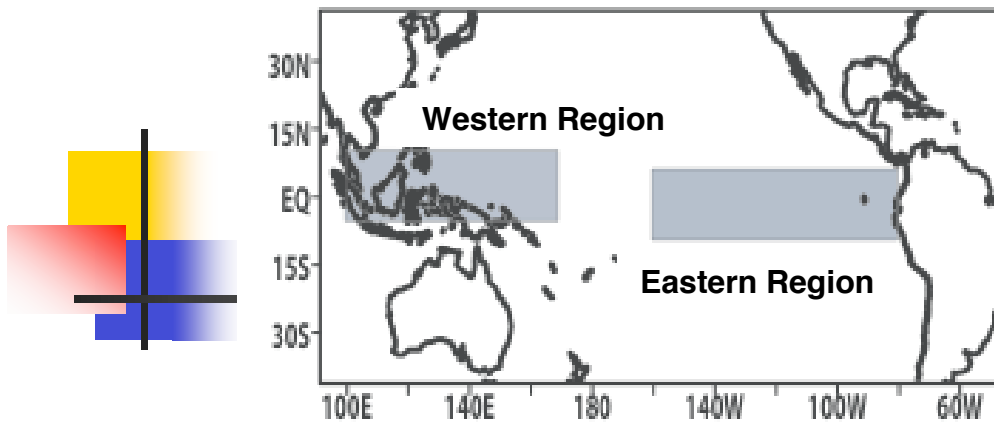




## Model vs Data Intercomparisons

**Pdfs at a Surface Site:  
Cloud Top Height/Base & Tau**

**(Comstock and Jacob  
submitted GRL, 2004**

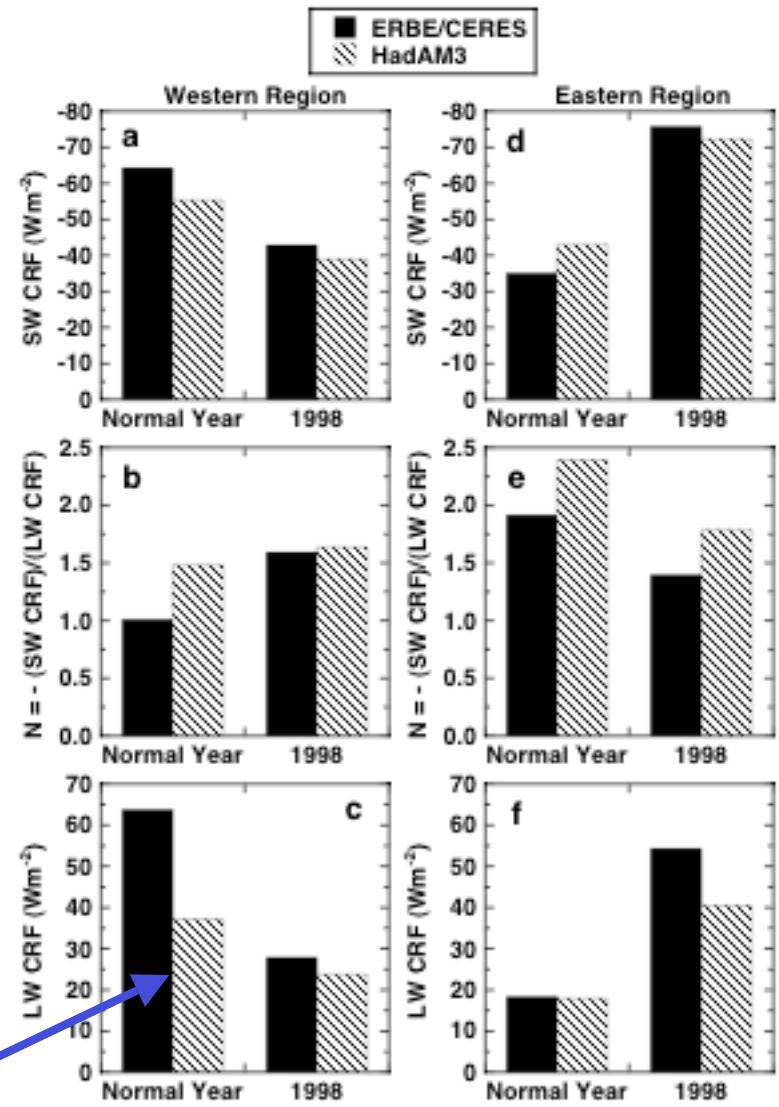


## Model vs Data Intercomparisons

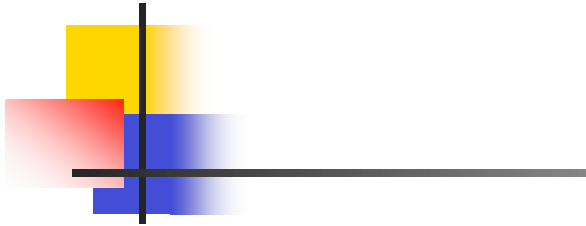
### Cloud Forcing/Ratio Response to El Nino

(Lu, Dong, Cess, Potter, 2004)

*How close should models agree for a given feedback uncertainty?*

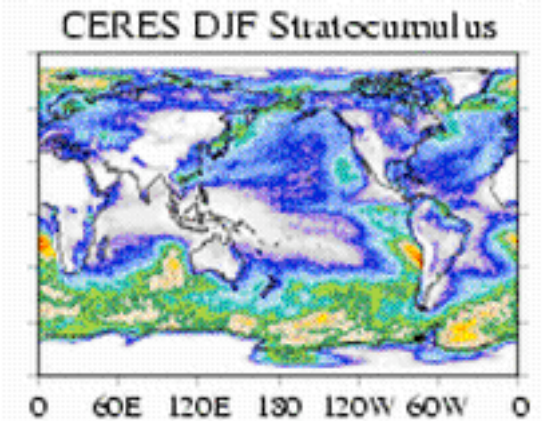
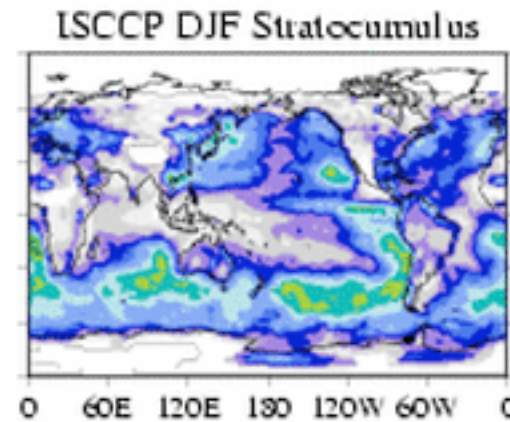
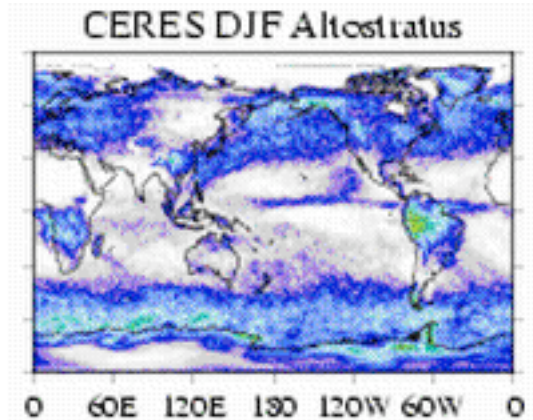
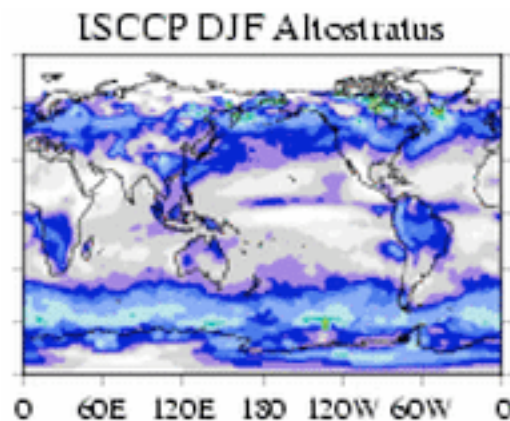
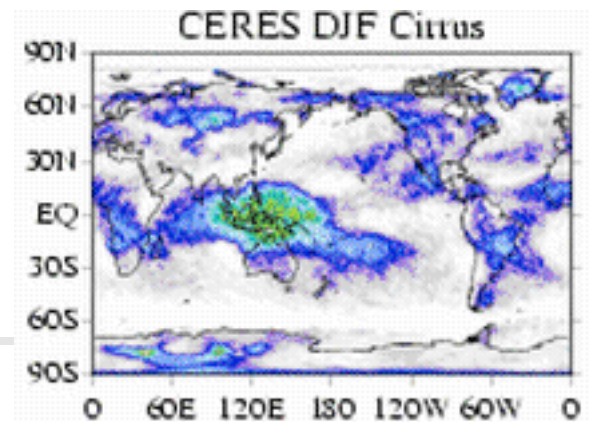
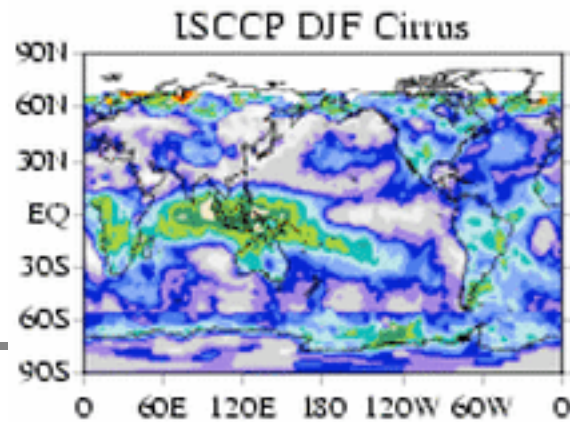




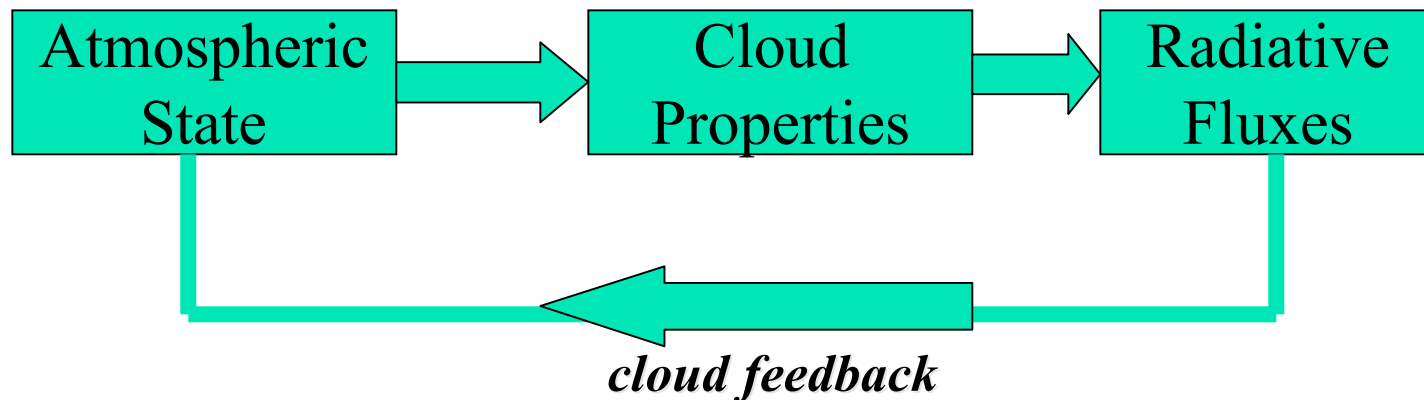


**ISCCP vs. CERES  
Cloud Type  
Frequency of Occurrence  
Wang, Loeb, Minnis 2004**

***GEWEX Radiation Panel  
Cloud Property Data  
and Radiative Flux Data  
Assessments begins  
late 2004***

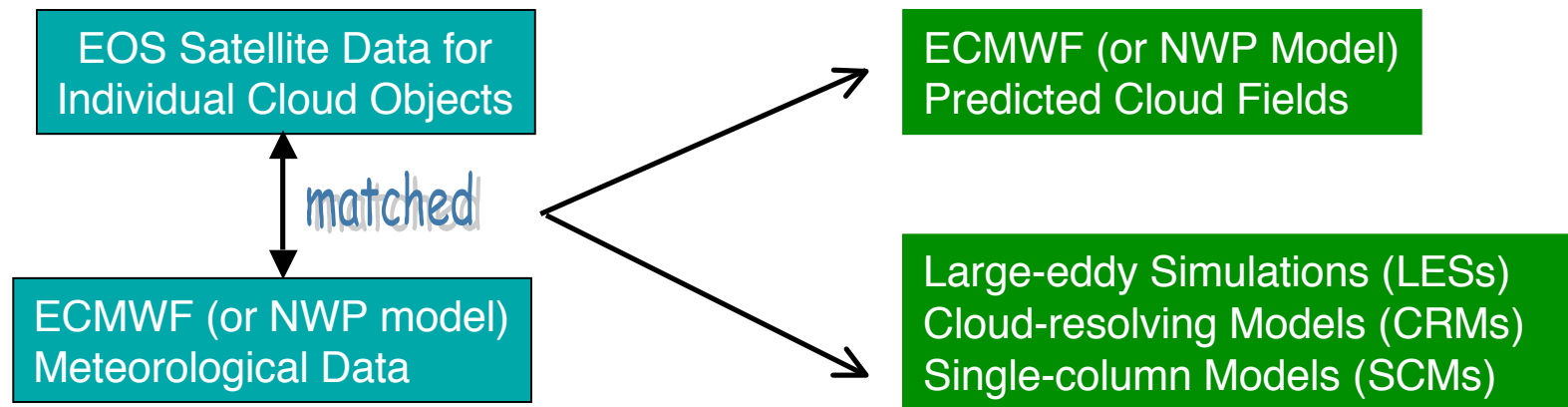


# Motivation



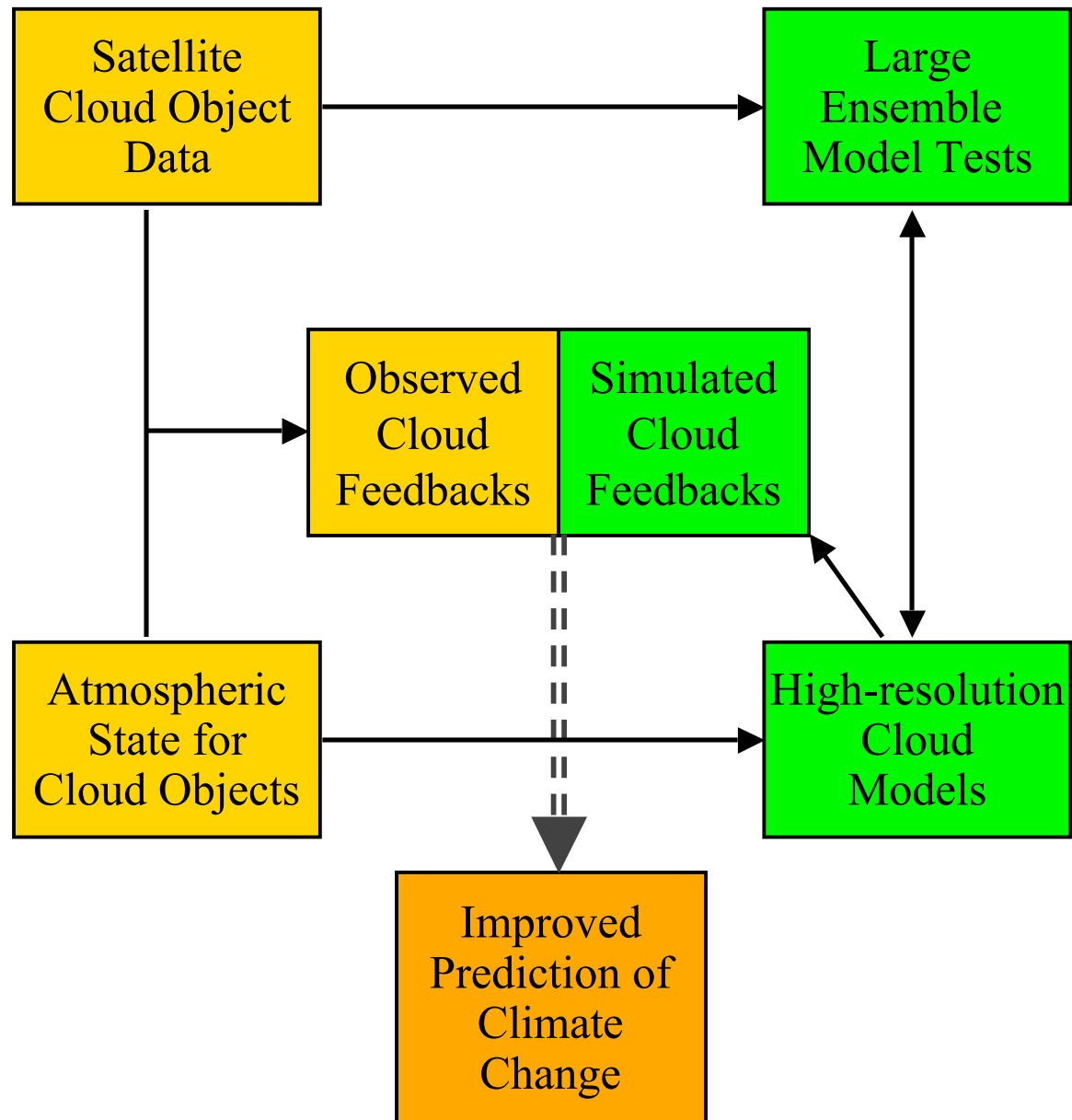
- ♣ Nonlinearity of cloud processes requiring observations on all relevant modeling scales (in space and in time)
- ♣ Existing methods of cloud model evaluation are incomplete

# Using satellite cloud object data for evaluating and improving CRMs and cloud parameterizations



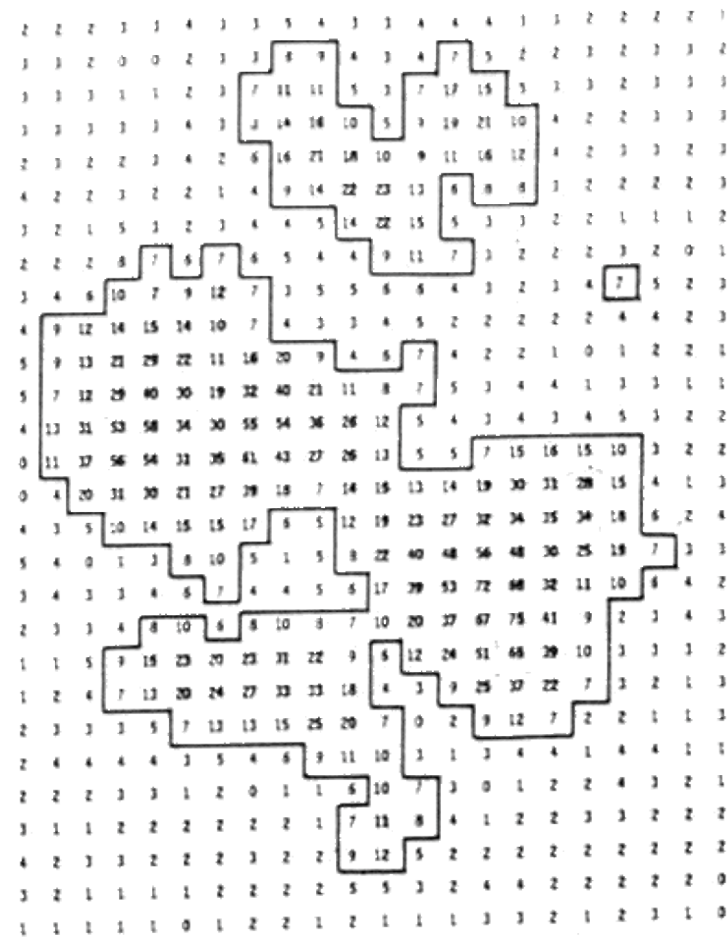
- Analyze the statistics of subgrid characteristics (PDFs) of satellite-observed cloud objects, *not* GCM gridbox means
- Match the CERES SSF (Single Scanner Footprint) cloud and radiation data with ECMWF meteorological data (T, q, u, v and advective tendencies)
- Perform cloud model simulations driven by ECMWF advective tendencies; an iterative process of improvement and evaluation of cloud models
- Also evaluate the ECMWF parameterization using its predicted cloud fields

# A cloud modeling strategy



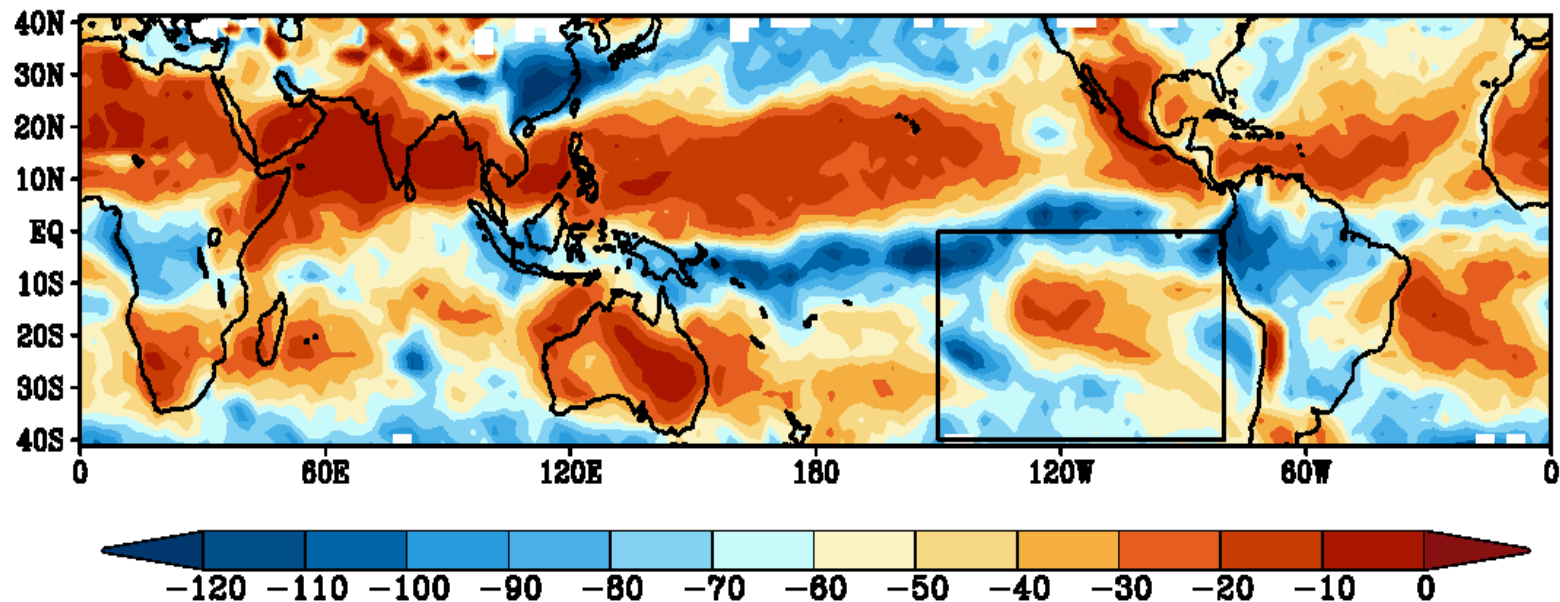
# Satellite data analysis method

- ♣ Define a cloud system as a contiguous region of the Earth with a **single dominant** cloud type (e.g. stratocumulus, stratus, and deep convection)
- ♣ Determine the shapes and sizes of the cloud systems by the satellite data and by the cloud property selection criteria (e.g. Wielicki and Welch 1986)



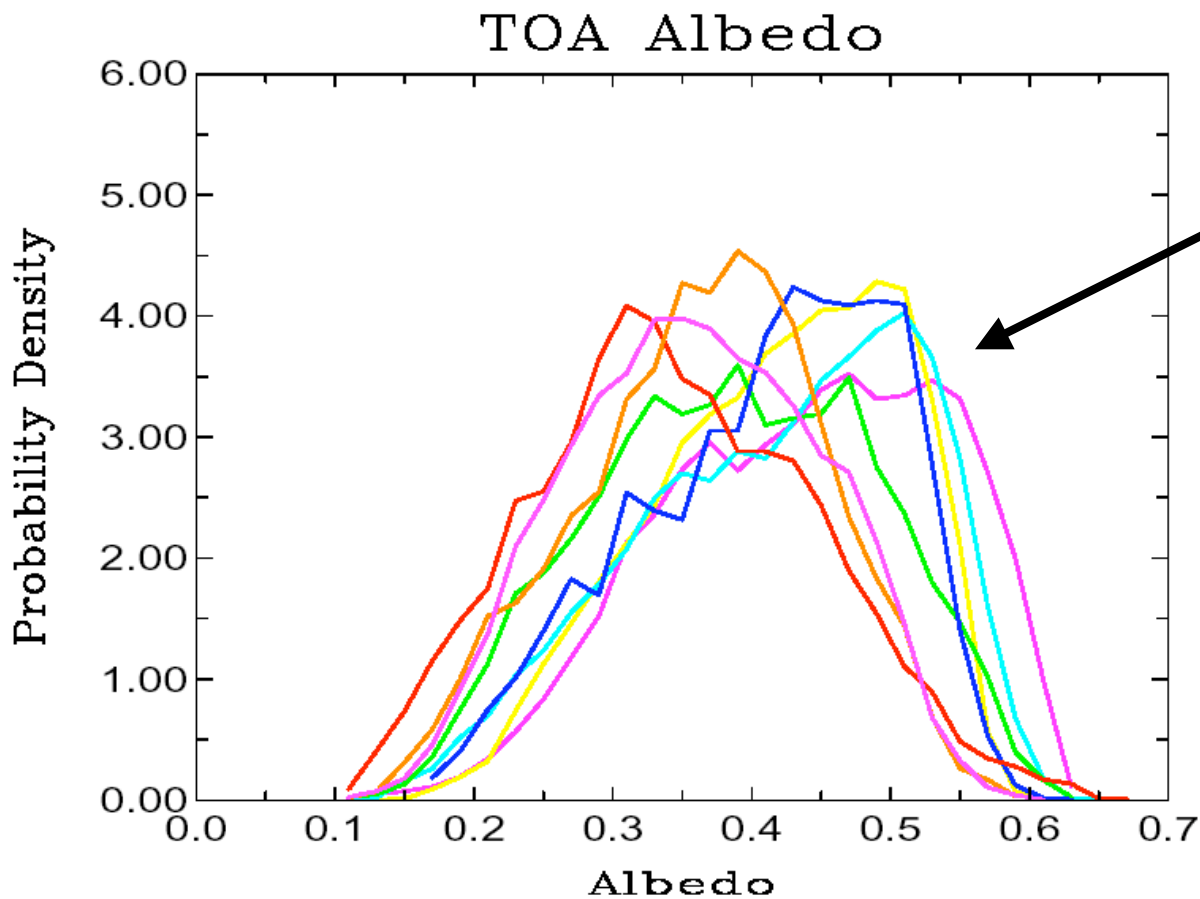
# Boundary Layer Cloud Object Region, Southeast Pacific, March 1998

CERES/TRMM Shortwave Cloud Radiative Forcing ( $\text{Wm}^{-2}$ ), March 1998



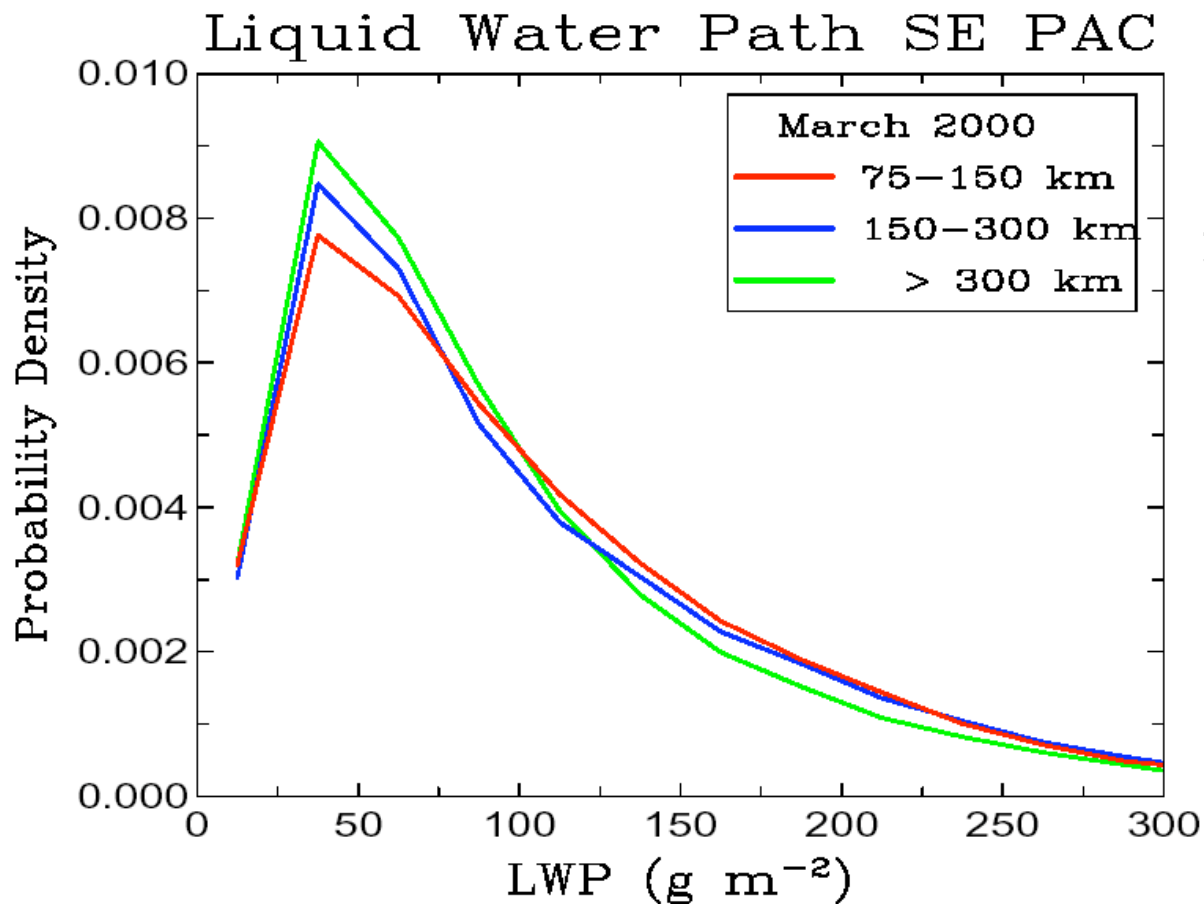
<i>Solid Stratus:</i>	<i>0.99 - 1.00 cloud fraction,</i>	$Z_{cld} < 3 \text{ km}$
<i>Stratocumulus:</i>	<i>0.40 - 0.99 cloud fraction,</i>	$Z_{cld} < 3 \text{ km}$
<i>Cumulus:</i>	<i>0.10 - 0.40 cloud fraction,</i>	$Z_{cld} < 3 \text{ km}$

# Overcast Boundary Layer: Observed CERES Cloud Object Pdfs for March, 1998



Sample individual pdfs for just 8 of the stratus cloud systems (CERES SSF TOA albedo)

# Overcast Boundary Layer: Observed CERES LWP Pdfs for March, 2000



Stratus:

Cloud Fraction = 1

Z<sub>cloud</sub> < 3 km

Water phase

LWP from  $\tau(\text{vis})_{\text{reff}}$

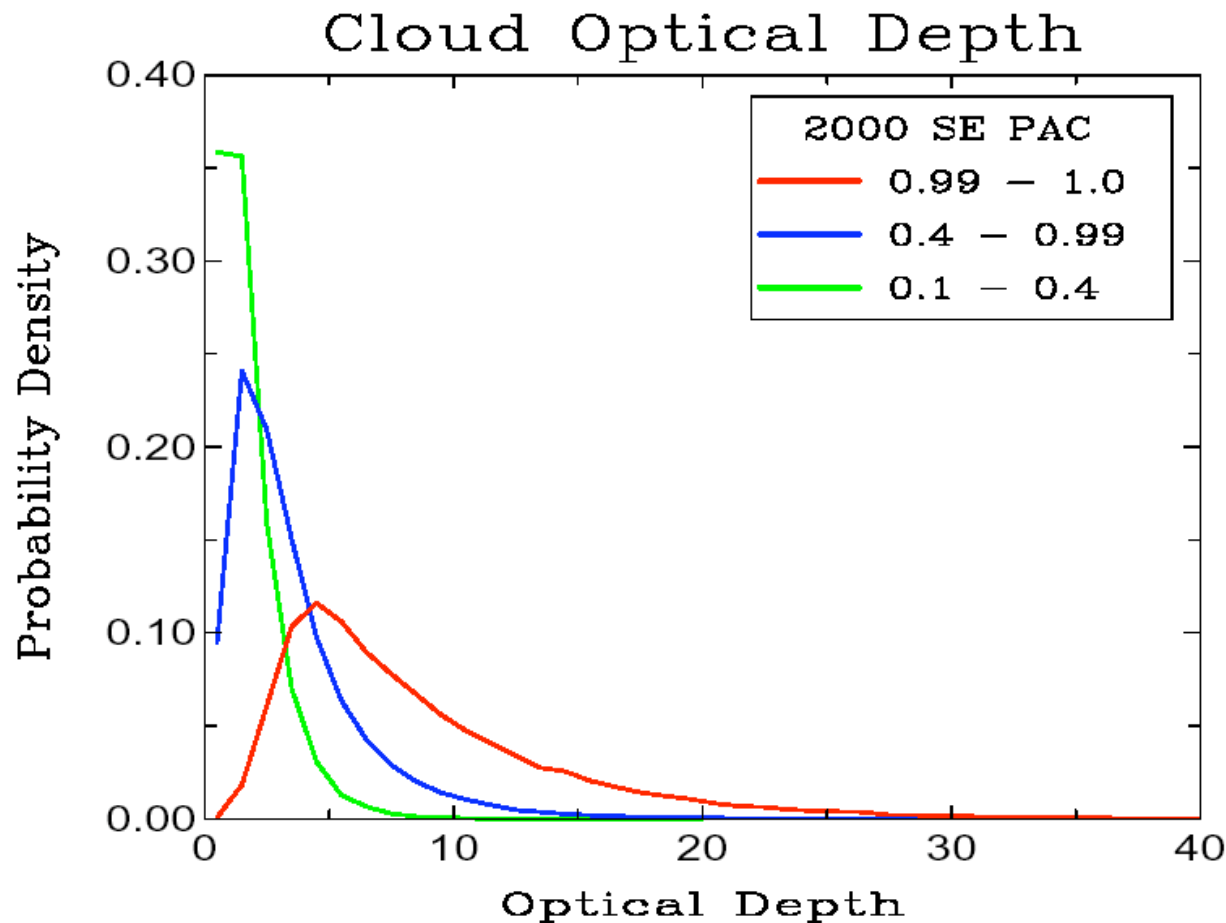
CERES SSF cloud

retrieved using VIRS imager

*Surprisingly, larger stratus decks do not have larger LWP amounts*

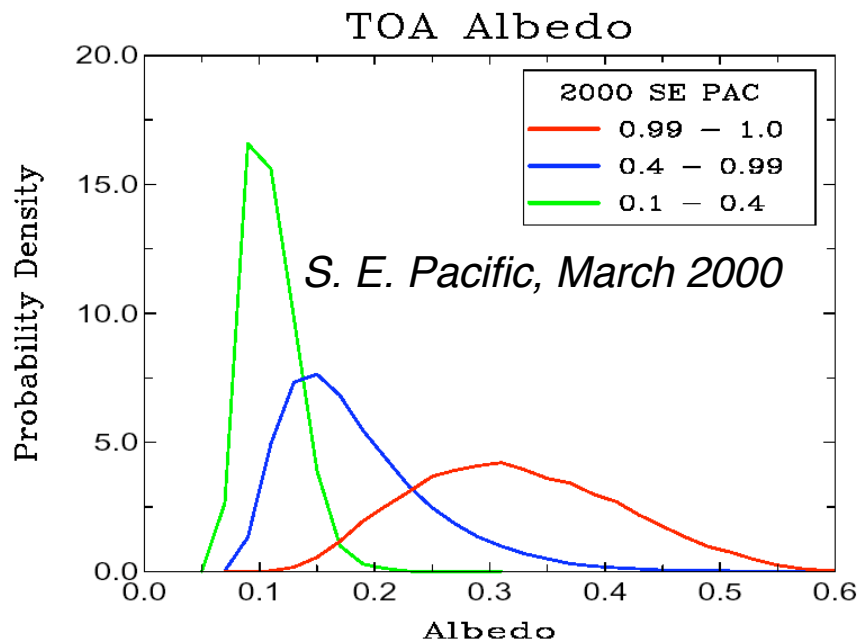


# Boundary Layer: Observed CERES Visible Optical Depth Pdfs for March, 2000



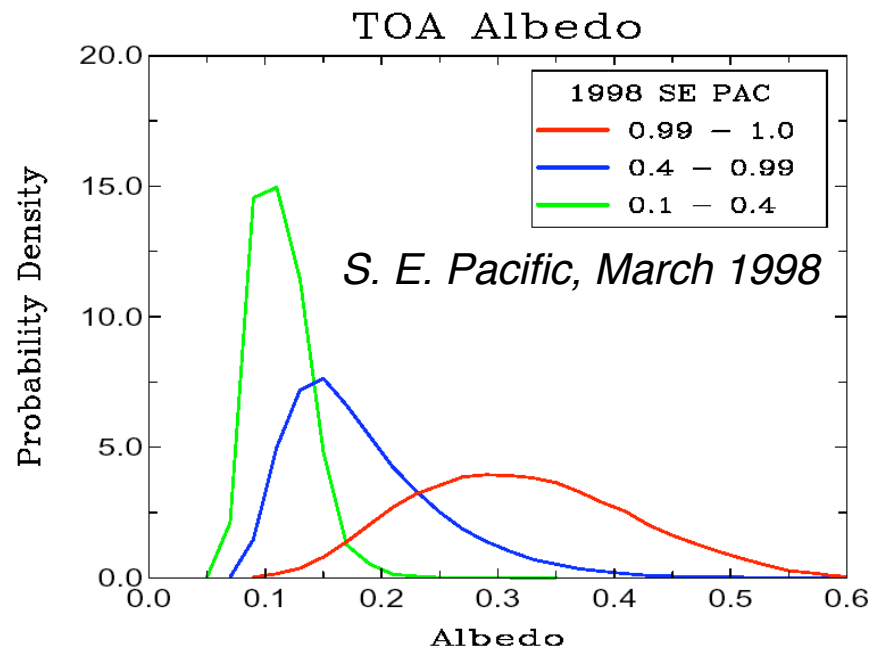
*Similar to  
Landsat Pdfs  
but from a large  
ensemble  
of boundary  
layer cloud  
systems using  
10 to 20km fov  
spatial scale:  
skewed  
distributions  
remain....*

# Boundary Layer: Observed CERES TOA Albedo Pdfs for March, 2000 vs March, 1998

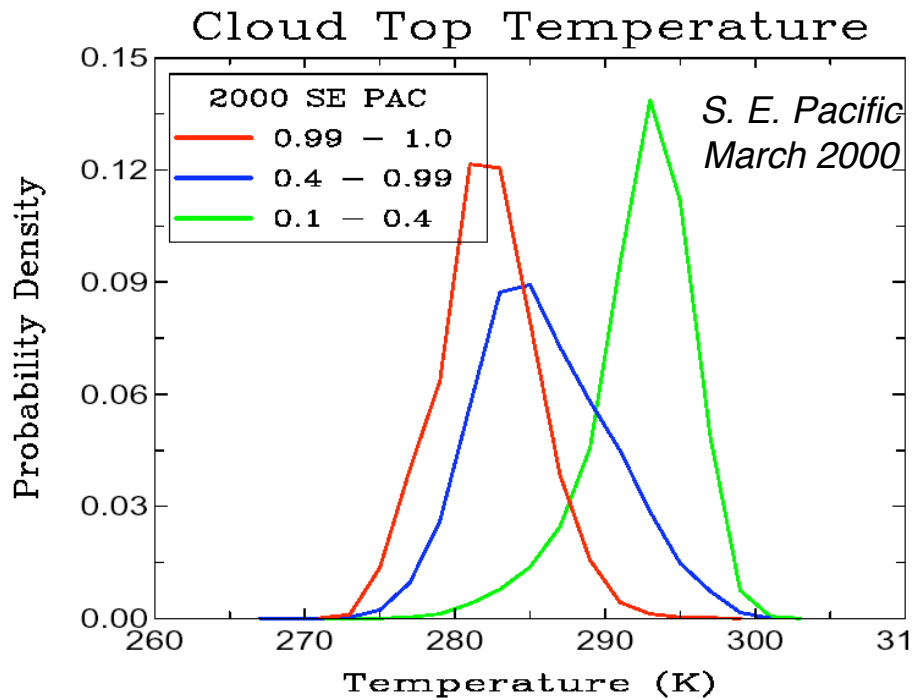


*Suggests stable properties by cloud type: next step to quantify how stable....*

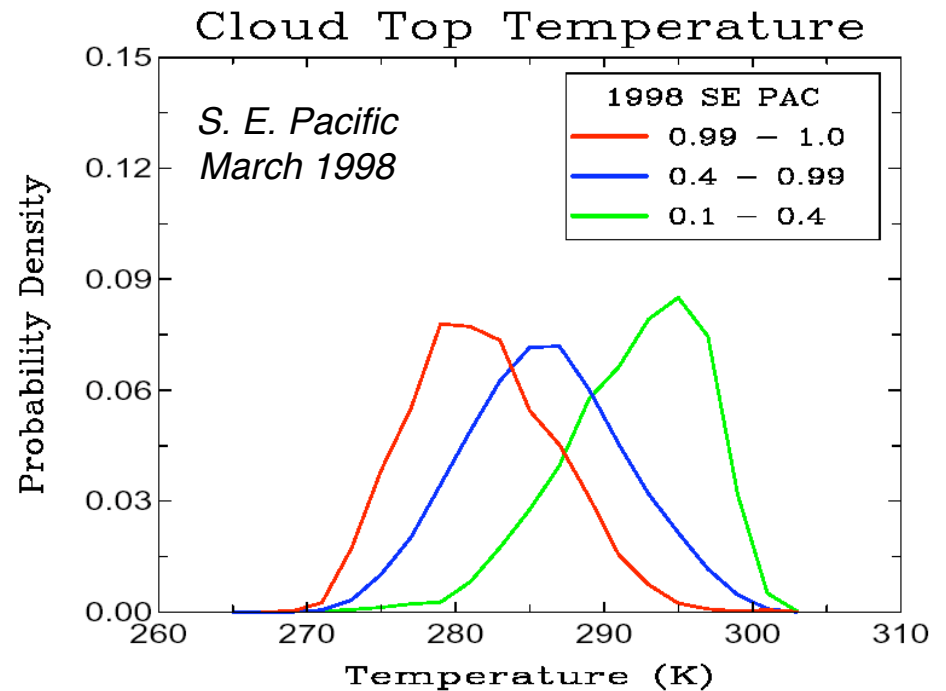
*No apparent difference in the S.E. Pacific, even though the Walker Cell strength reduced, Hadley cell strengthened...*



# Boundary Layer: Observed CERES Cloud Top Temperature Pdfs for March, 2000 vs March, 1998

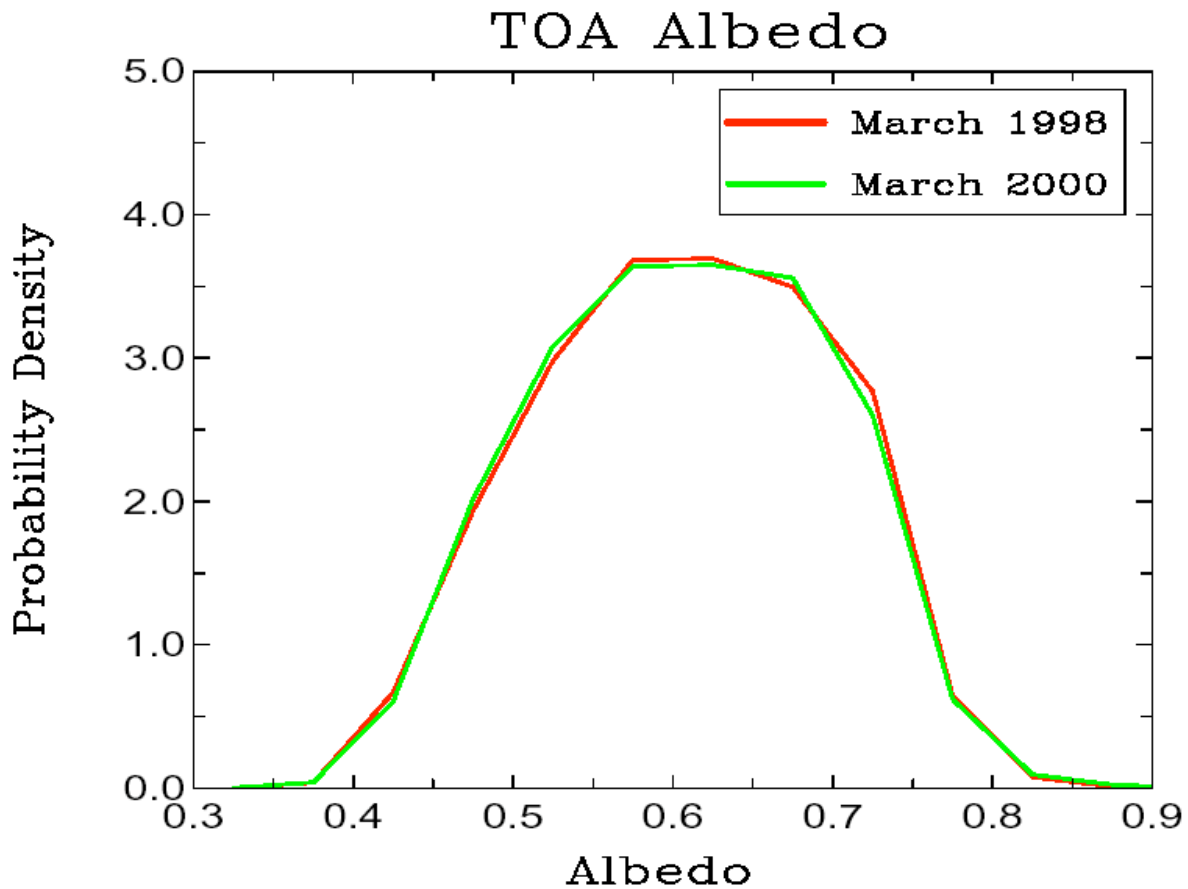


*March 2000: Colder SST (La Nina) & Colder Cloud Top Temperature, but Narrower Frequency Distribution*



# Large Deep Convective Systems:

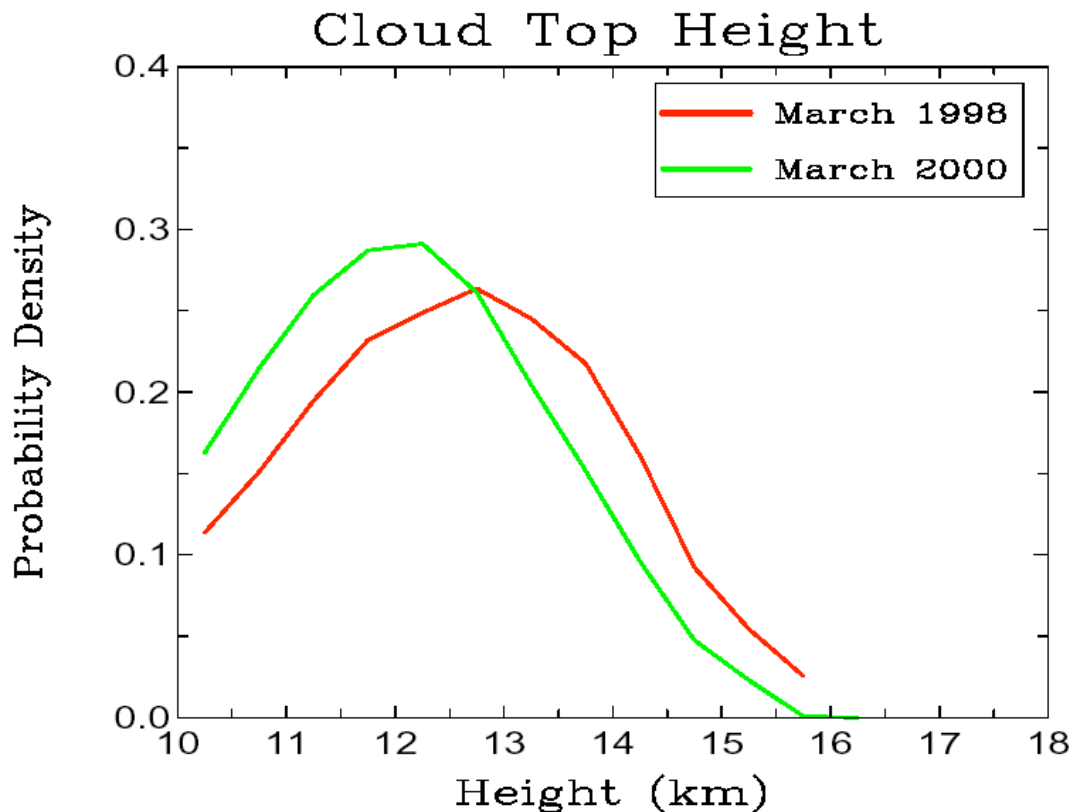
$Z_{\text{cld}} > 10\text{km}$ ,  $\tau > 10$ ,  $C_f = 1$ , Diameter  $> 300\text{km}$   
CERES TOA Albedo



Across the tropics (25N to 25S) large convective systems appear invariant between the 98 El Niño and 2000 La Niña phases of ENSO for TOA albedo pdf.

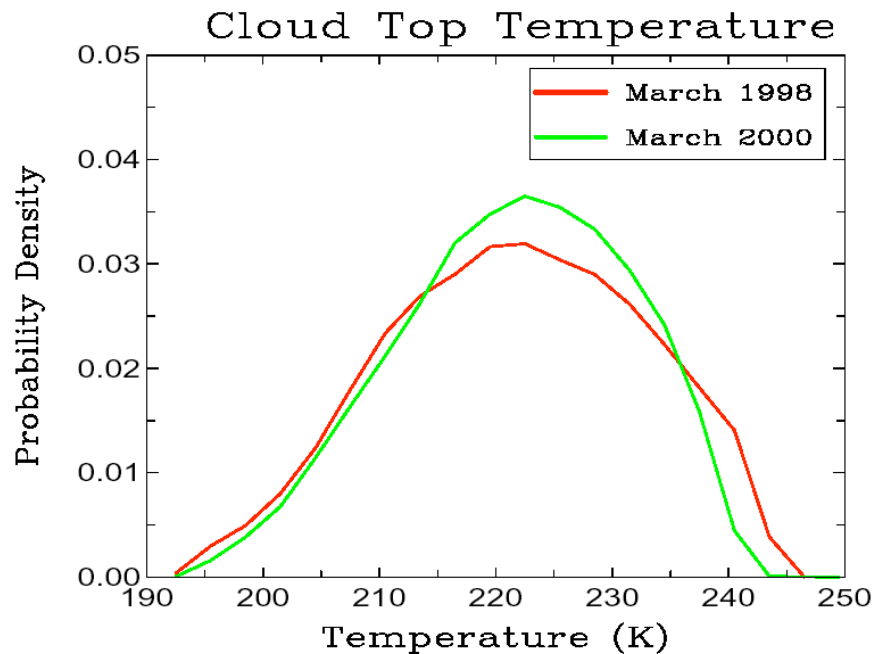
# Large Deep Convective Systems:

$Z_{\text{cld}} > 10\text{km}$ ,  $\tau > 10$ ,  $C_f = 1$ , Diameter  $> 300\text{km}$   
CERES Cloud Height using MODIS



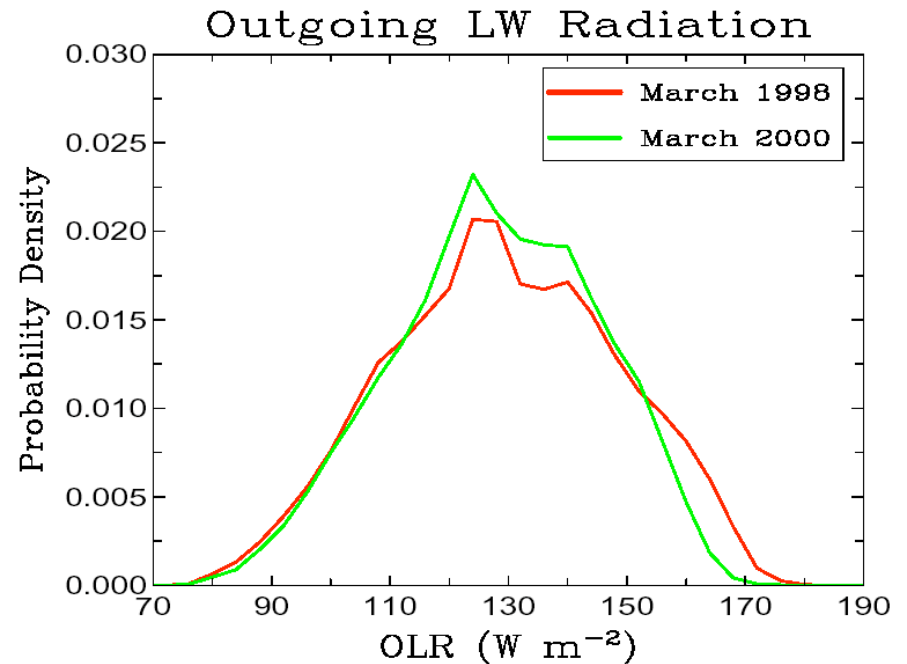
Across the tropics (25N to 25S) large convective systems, however appear to increase cloud height by about almost 1 km during the 1998 El Nino

# Large Deep Convective Systems: $Z_{\text{cld}} > 10\text{km}$ , $\tau > 10$ , $C_f = 1$ , Diameter $> 300\text{km}$ CERES TOA LW Flux and Cloud Eff Temp using MODIS



*Or just the dynamics of these  
large convective complexes?*

*Cloud height changes but much  
smaller cloud temperature and  
TOA LW flux changes:  
Hartmann hypothesis on  
radiative control of tropics?*





## So what do models predict?

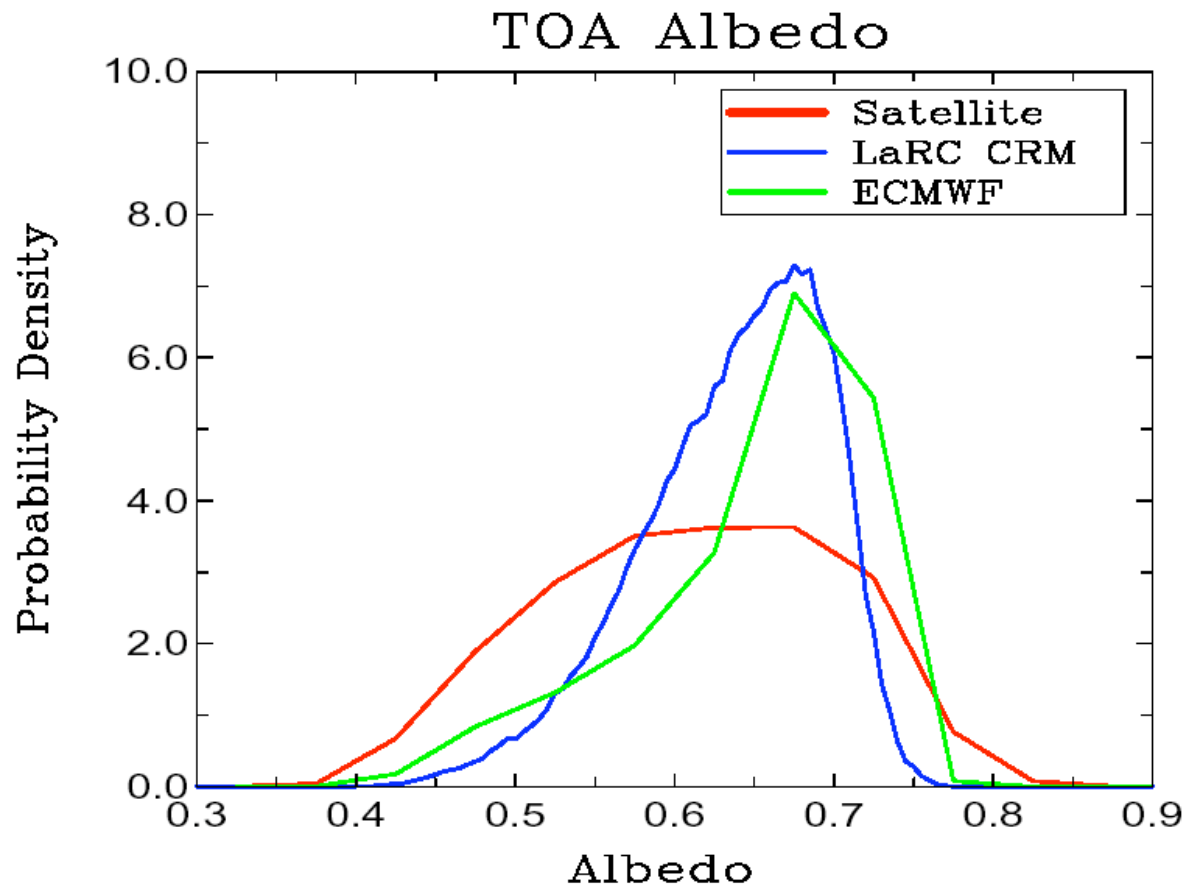
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- √ ECMWF: 0.5 degree, 6-hourly assimilation analysis, including clouds
- √ CRM: LaRC2d CRM: 1 km resolution 2-D, 3rd order turb. closure (UCLA/CSU; Krueger 1988; Xu and Randall 1995)
- √ 29 cases of tropical convective systems with diameters greater than 300 km for March 1998: *Z<sub>cld</sub>>10km, tau>10, ice phase, overcast*

# Large Deep Convective Systems:

$Z_{\text{cld}} > 10\text{km}$ ,  $\tau > 10$ ,  $C_f = 1$ , Diameter  $> 300\text{km}$

March, 1998, 25N to 25S, 29 cloud systems



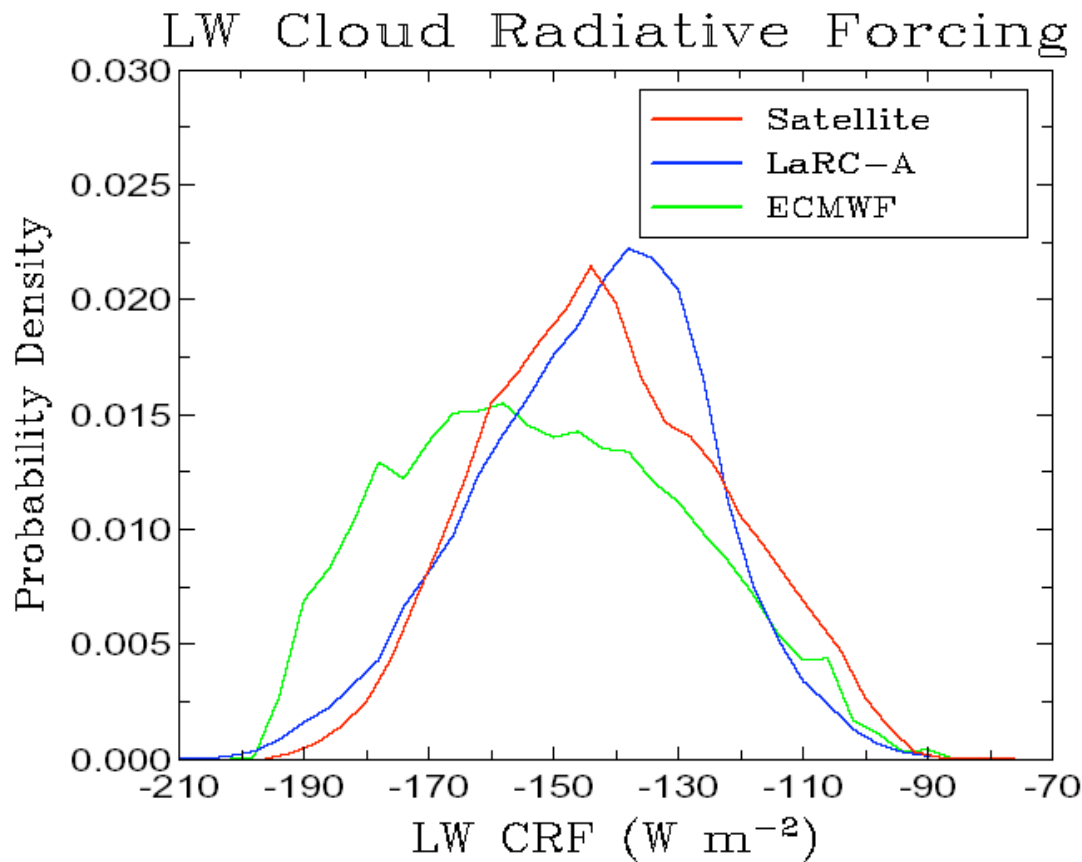
## *TOA Albedo*

*differences are large  
ECMWF clouds are  
too optically thick, with  
insufficient variability.  
CRM is an improvement  
but still needs  
substantial improvement:  
CRM and especially  
ECMWF will overestimate  
cloud surface cooling*



# Large Deep Convective Systems:

$Z_{\text{cld}} > 10\text{km}$ ,  $\tau > 10$ ,  $C_f = 1$ , Diameter  $> 300\text{km}$   
March, 1998, 25N to 25S, 29 cloud systems



*LW Cloud Radiative Forcing  
ECWMF clouds too thick  
and cold. CRM a much  
better prediction of the  
LW cloud radiative effects.*

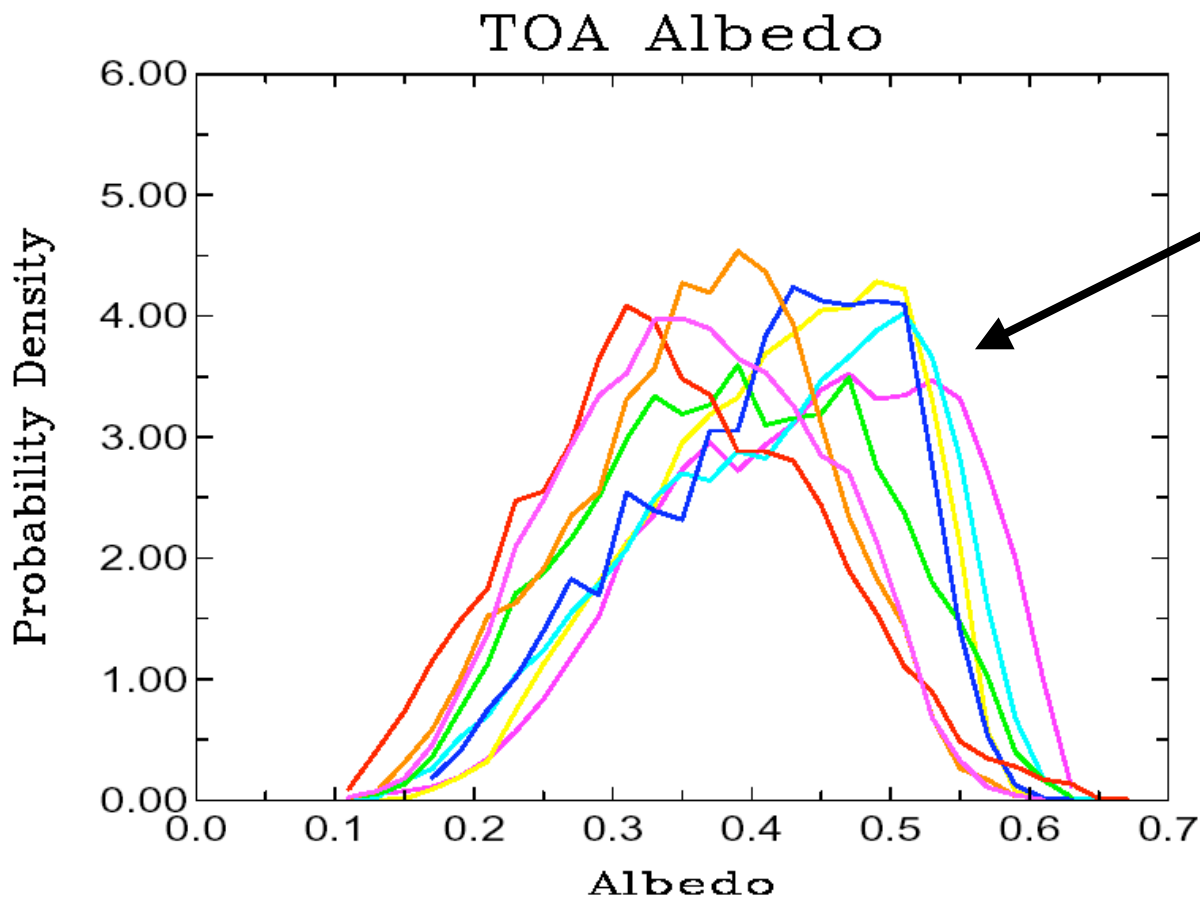


# Conclusions

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- *Cloud objects useful for examining cloud changes by cloud type*
- *Climate change can be separated into:*
  - *changing frequency of cloud type (dominant?)*
  - *changing properties of a cloud type (secondary?)*
  - *test how well models do each cloud change*
  - *with larger ensembles, separate by meteorological state*
    - *e.g. SST, stability, vertical velocity, wind shear, etc*
  - *do models handle the partial derivative of cloud properties versus atmospheric state change? key for cloud feedback*
- *How accurate should models and data agree?*
  - *statistical noise: can beat down with larger samples*
  - *new radiative flux ensemble errors by cloud type very small*
  - *what level differences are key to climate change? critical TBD!*
  - *errors in atmospheric input state: evolve over time, test sensitivity*

# Overcast Boundary Layer: Observed CERES Cloud Object Pdfs for March, 1998



Can we predict why these stratus systems differ as a function of dynamic state?

# Model/Data Comparison Methods

- **Radiation/Cloud Focused**
  - Classic: Monthly Mean Fluxes, Clouds, Aerosols
  - ISCCP 2-D Histograms of Cloud Height/Optical Depth/Frequency of Occurrence
  - By 2-D histogram principle components: the Jacob approach
  - By cloud type: The Xu approach
- **Atmospheric Dynamics Focused**
  - By vertical velocity: the Bony approach
  - By low/high surface pressure and fronts: the Tseuloudis approach.
- **dRadiation/dDynamics and dCloud/dDynamics**
  - Partial derivatives of cloud type/atmosphere state
  - Nonlinear approaches such as neural net

# Model Test Types

- **Fully Coupled Ocean/Atmosphere/Land/Cryosphere**
- **AMIP style specified SST and Sea Ice**
  - Normal atmosphere GCM with SCM clouds
  - MMF atmosphere GCM with CRM clouds
- **Weather Prediction mode (initial condition large scale dynamics)**
  - SCM global cloud predictions
  - MMF CRM global cloud predictions
  - Regional SCM and CRM predictions by cloud type
  - Regional SCM and CRM predictions over surface sites (e.g. ARM)

# Model Tests: Signal to Noise Issues

- **Internal climate noise: strong function of space/time scale**
  - i) red spectrum at weather scales
  - ii) blue spectrum at large time/space scales
  - iii) estimate using models and observations
  - iv) models a lower bound on climate noise?
  - v) results are a function of comparison type:  
grid box versus cloud type (eulerian/lagrangian)
  - vi) Is there a climate prediction limit analogous to weather prediction?
- **Data Errors**
  - i) strong function of time/space scale
  - ii) strong function of physical variable and cloud condition
  - iii) can also be a function of remote sensing conditions  
(e.g. viewing angle, solar zenith)
  - iv) function of comparison strategy: grid box versus cloud type very different
  - v) how can we better use models to set observing requirements?

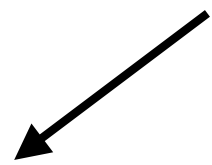
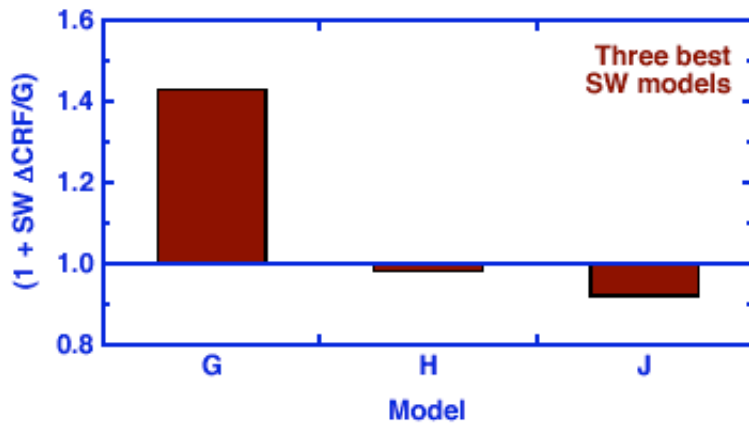
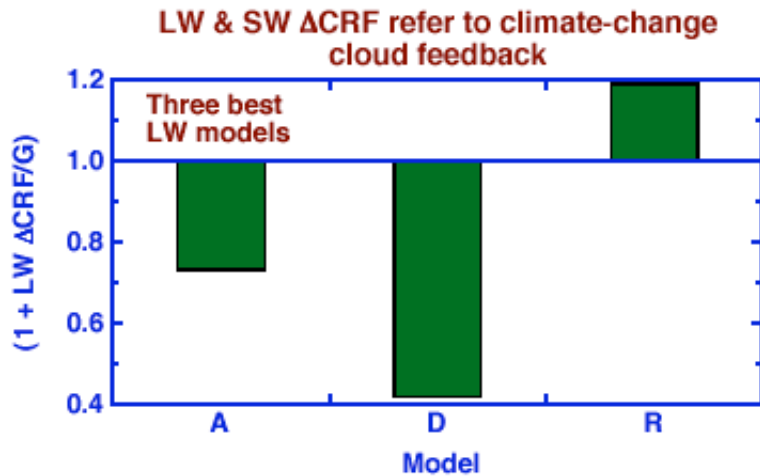
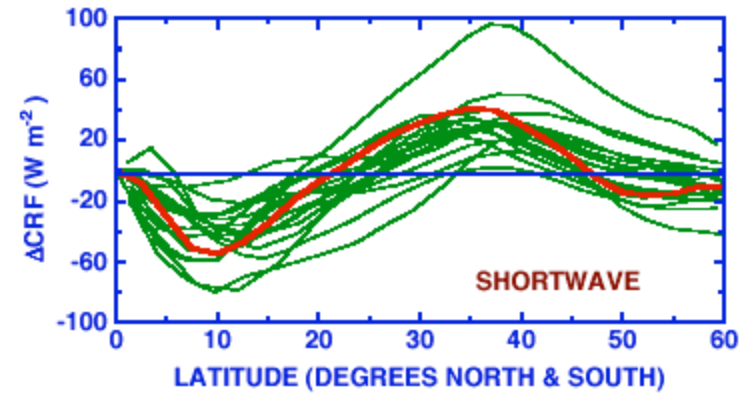
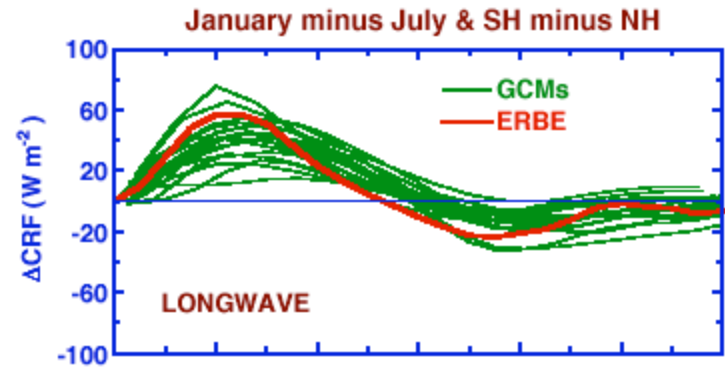
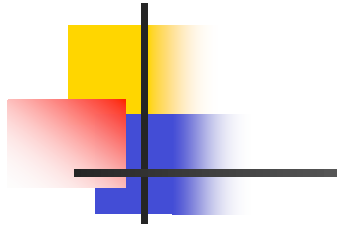
# Key Comparison Issues

- **Accuracy requirements tests to constrain feedback to +/- X%?**
  - can limiting case thought experiments help?
  - can we use different climate models as “Earth plus N other Planets” to test comparison logic on how to tie model comparison accuracy to climate sensitivity accuracy?
  - What collection of tests might be sufficient? Is closure possible?
  - If aliens gave us 10 climate models and said one was perfect: how would we know which one? Could we tell if they were lying?
- **How do we verify completeness of tests? How do we measure it?**
  - completeness of dynamic states tested
  - completeness of cloud types tested
  - completeness of aerosol types tested

# Model vs Data Intercomparisons

## Seasonal Change Tests

(Cess et al)



**Conclusion: Models Best at Seasonal Climate Change don't agree on century scale climate change: Necessary but not sufficient condition.**

*What collection of tests is sufficient?*



# Key Comparison Issues

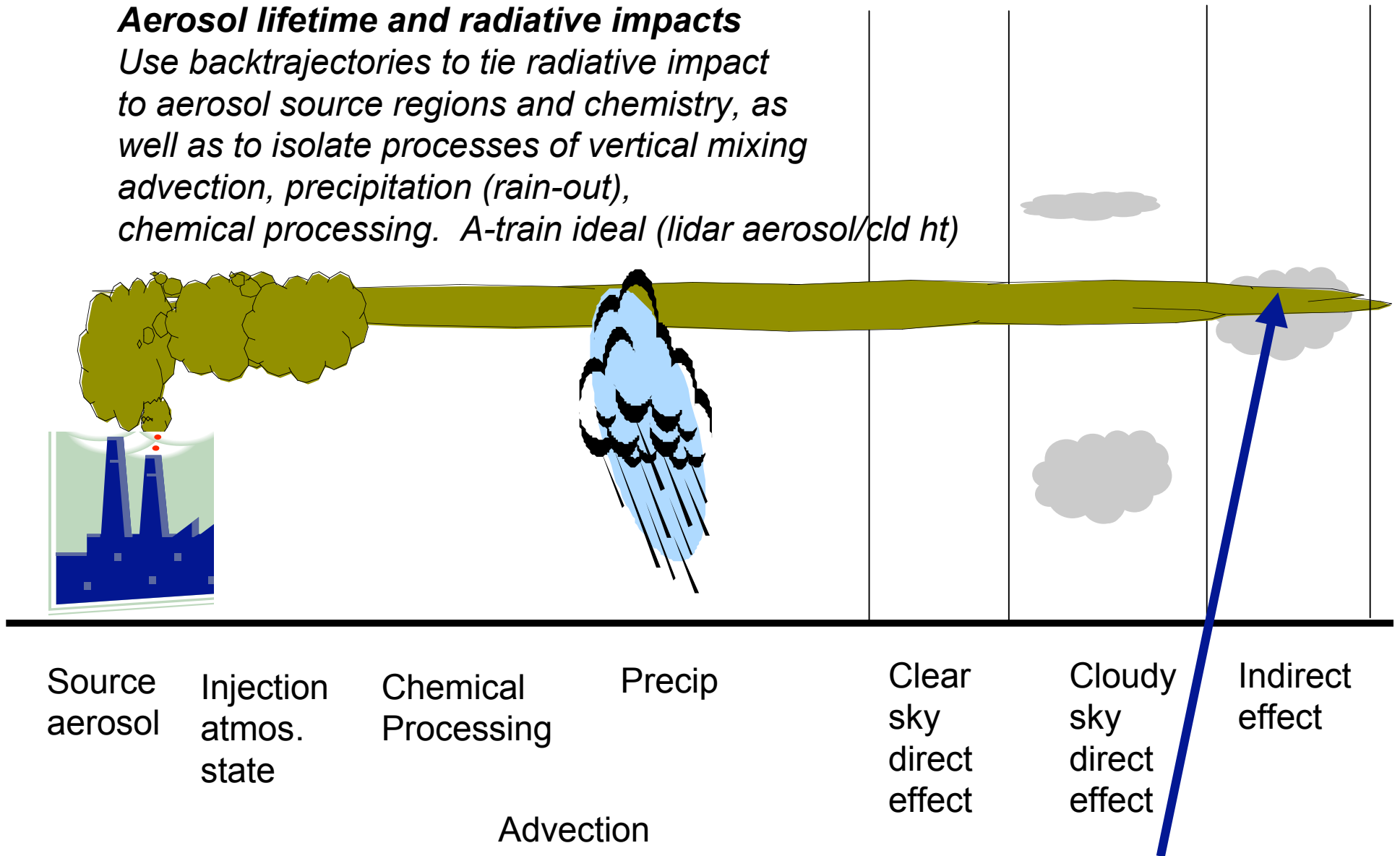
- **Can we predict the Cause/Effect power of tests? i.e. rapidity of forced “Eureka’s”**
- **For aerosol indirect effects with clouds: must decouple dynamics changes in clouds from aerosol changes: for any location/source two are strongly linked.**
  - example: azores with and without European aerosols in boundary layer.
  - this means aerosol “history” must be known for days prior to comparison.
- **Do cloud comparisons need past history of CCN for feedback? or only for forcing?**
- **Recent Xie article in BAMS: dcloud/dSST changes sign with spatial scale:**
  - positive at 1 to 10 km scale, but negative at hundreds of km scale (Klein and Hartmann).

# Key Physical Processes

- **Vertical Velocities**
- **Microphysics parameterizations**
- **Closure for boundary layer clouds even in CRMs: ultimately LES in CRM?**
- **How critical are 3-D radiation effects?**
- **Small scale clear/cloudy radiation are critical (recent MMF tests show this)**

## ***Aerosol lifetime and radiative impacts***

*Use backtrajectories to tie radiative impact to aerosol source regions and chemistry, as well as to isolate processes of vertical mixing advection, precipitation (rain-out), chemical processing. A-train ideal (lidar aerosol/cld ht)*



***Must unscramble cloud fluxes/properties and dynamic state in order to isolate cloud indirect effect....***

# Backup Slides



# Analysis of ECMWF predicted cloud fields

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- ♣ ECMWF meteorological data

- $0.5^\circ \times 0.5^\circ$  gridded, six hourly analysis from data assimilation
- temperature, specific humidity, horizontal wind components

- ♣ ECMWF predicted cloud fields (prognostic parameterization)

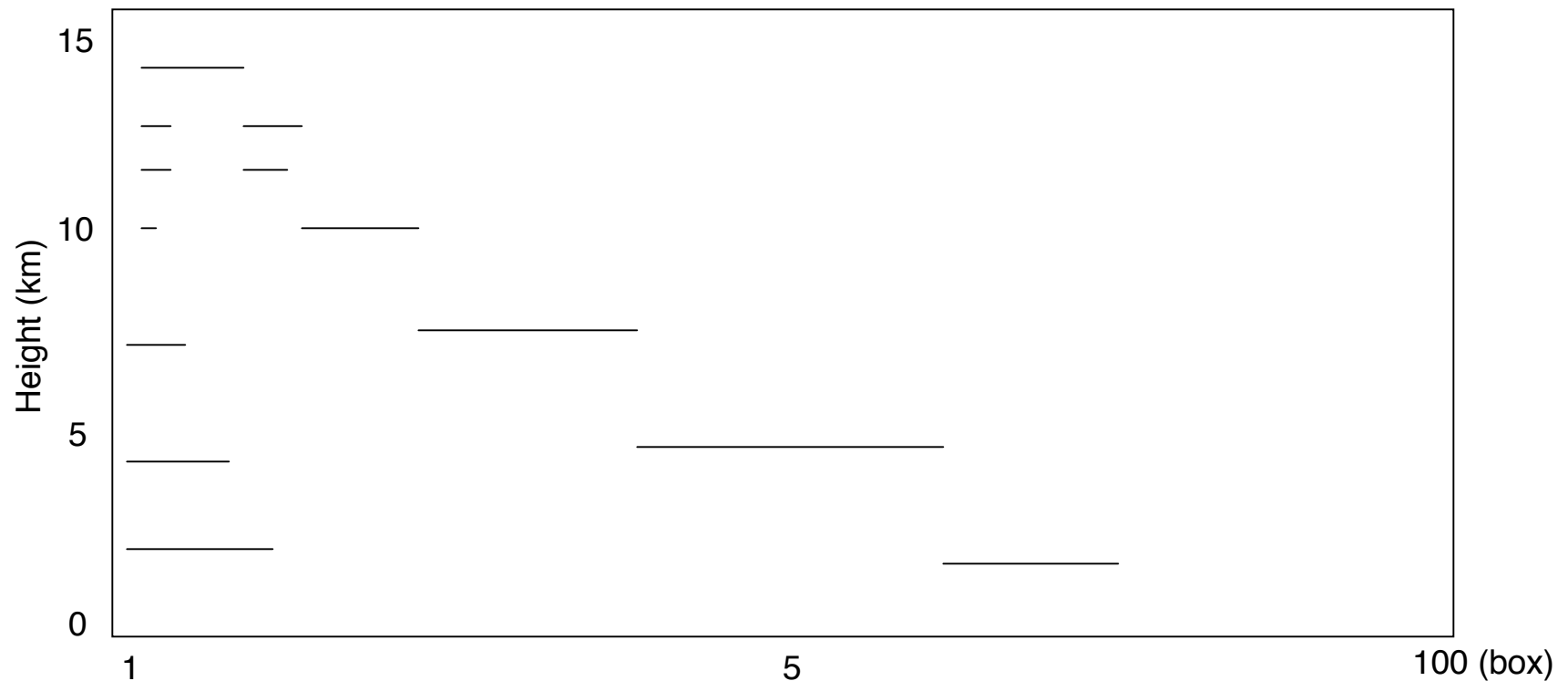
- $0.5^\circ \times 0.5^\circ$  gridded, six-hour predictions
- cloud liquid water content
- cloud ice water content
- cloud cover

- ♣ ECMWF grids are much bigger than some CERES SSF fovs (CERES TRMM range from  $\sim 10$  to  $20$  km diameter)

- ♣ ECMWF does not provide cloud optical properties; we need to use the Fu-Liou radiation code, but it does not treat partially cloudy columns

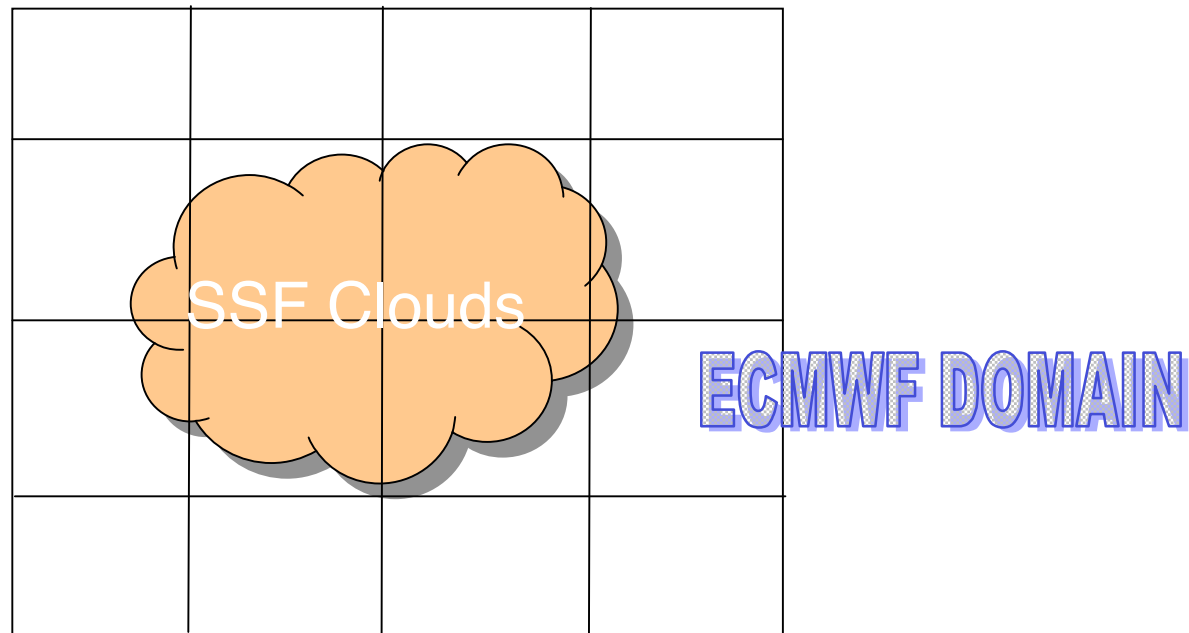
# Analysis of ECMWF predicted cloud fields (cont.)

- Divide an ECMWF grid box into 30 subgrid boxes (~10km CERES flux scale)
- Use the maximum/random overlap assumption (Klein & Jacob 1999)
- Use the Fu-Liou radiation code to obtain cloud optical properties and radiative fluxes for each subgrid box



# Comparison of SSF with ECMWF

- Only subgrid boxes with cloud top height  $> 10$  and cloud optical depth  $> 10$  are selected for statistical analysis
- Cloud top is defined as infrared absorption optical depth 1 into the cloud to be similar to satellite effective radiating cloud top
- Clouds within the near vicinity of the observed cloud systems are also included





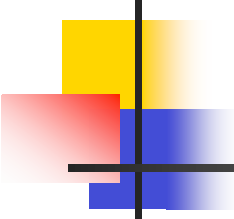
# Cloud resolving model simulation:

## What is a cloud-resolving model (CRM)?

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- ✓ Sufficient spatial and temporal resolution to resolve individual cloud elements ( $\sim 1$  km)
- ✓ Sufficient large domain and long time scale for statistical analyses of cloud systems
- ✓ Explicitly resolve cloud-scale and mesoscale dynamical processes
- ✓ Need to parameterize turbulence, cloud microphysics and radiative transfer
- ✓ Often used as a tool for cloud parameterization development for GCMs
- ✓ Used as a “Super-Parameterization” inside GCM grid boxes.





# Cloud-resolving model simulation: Description of the models

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LaRC2d CRM (UCLA/CSU; Krueger 1988; Xu and Randall 1995)

1. Two-dimensional, anelastic dynamics (no sound waves)
2. Third-moment turbulence closures (35 prognostic equations and one diagnostic equation)
3. Three-phase cloud microphysics parameterization (Lin et al. 1983; Krueger et al. 1995)
4. Harshvardhan et al. (1987) radiative transfer parameterization

LaRC3d CRM (Advanced Regional Prediction System; Xue et al. 2000)

1. 2-D or 3-D fully compressible dynamics
2. Prognostic turbulent kinetic energy (TKE) closure
3. Three-phase cloud microphysics parameterization (Lin et al. 1983)
4. Chou (1990, 1992) and Chou and Suarez (1994) radiative transfer parameterization



# Cloud resolving model simulation: Design of simulation

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- √ 2-D (x-z), horizontal grid size is 2 km
- √ Prescribe large-scale advective tendencies that are calculated from ECMWF data and averaged over an square area three times as great as the satellite observed cloud system
- √ The advective tendencies are assumed to be quasi-steady
- √ Simulation lasts for 24 h
- √ Only the last 12 h is analyzed