Global Climate Change

Key Messages:

- Human activities have led to large increases in heat-trapping gases over the past century.
- Over the last 100 years, global average temperature and sea level have increased, and precipitation patterns have changed.
- Numerous independent lines of evidence show that many of the climatic changes of the past 50 years are primarily human-induced.
- Global temperatures will continue to rise over this century; by how much and for how long depends on a number of factors, including the amount of heat-trapping emissions and how sensitive the climate is to those emissions.

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L23 This introduction to global climate change explains very briefly what has been happening to the world's L24 climate and why, and what is projected to happen in the future. While this Report focuses on climate change L25 impacts in the United States, understanding these changes and their impacts necessarily requires an under-L26 standing of the global climate system.





Many changes have been observed in global climate over the past century. The nature and causes of these changes have been comprehensively chronicled in a variety of recent reports, such as those by the Intergovernmental Panel on Climate Change (IPCC) and the U.S. Climate **Change Science Program** (CCSP). This Section does not intend to duplicate these comprehensive efforts, but rather to provide a brief synthesis, and to integrate more recent work with the assessments of the IPCC, CCSP, and others.

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Human activities have led to large increases in heat-trapping gases over the past century.

The Earth's climate depends on the functioning of a large natural "greenhouse effect". The greenhouse effect is the result of gases like water vapor, carbon dioxide, ozone, methane, and nitrous oxide, which absorb heat radiated from the Earth's surface and lower atmosphere and then radiate much of the energy back towards the surface. Without this natural greenhouse effect, the average surface temperature of the Earth would be about 60°F colder. However, human activities release additional heat-trapping gases into the atmosphere, particularly through the burning of fossil fuels (coal, oil, and natural gas). This intensifies the natural greenhouse effect, thereby changing the climate of our planet.

Earth's climate is influenced by a variety of factors, both human-induced and natural. The increase in the carbon dioxide concentration has been the principal factor causing warming over the past 50 years. Its concentration has been building up in the Earth's atmosphere since the beginning of the industrial era, primarily due to the burning of fossil fuels and the clearing of forests. Human activities have also increased the emissions of other greenhouse gases, such as methane, nitrous oxide, and halocarbons². These emissions are thickening theR1blanket of heat-trapping gases in Earth's atmo-R2sphere, causing surface temperatures to rise.R3

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Heat-trapping gases

Carbon dioxide concentration has increased due R6 to the use of fossil fuels in electricity generation, R7 transportation, industrial processes, and space and **R**8 water heating. It is also produced as a by-product R9 during the manufacturing of cement. Deforestation R10 provides a source of carbon dioxide, and reduces its R11 uptake by trees and other plants. Globally, over the R12 past several decades, about 80 percent of human-R13 induced carbon dioxide emissions came from the R14 burning of fossil fuels, while about 20 percent R15 resulted from deforestation. The concentration of R16 carbon dioxide in the atmosphere has increased by R17 roughly 35 percent since the industrial revolution². R18

Methane concentration has increased mainly as a result of agriculture, raising livestock (which produce methane in their digestive tracts), mining, transportation, and use of certain fossil fuels, sewage, and decomposing garbage in landfills. About 70 percent of the emissions of atmospheric methane are now related to human activities.

Nitrous oxide concentration is increasing as a result of fertilizer use and fossil fuel burning.

Halocarbon emissions come from the R31 release of manufactured chemicals to the R32 atmosphere. Examples include chloro-R33 fluorocarbons (CFCs), which were used R34 extensively in refrigeration and other R35 industrial processes before their presence R36 in the atmosphere was found to cause R37 stratospheric ozone depletion. The abun-R38 dance of these gases in the atmosphere is R39 now decreasing as a result of international R40 regulations designed to protect the ozone R41 layer. Continued decreases in halocarbon R42 emissions are expected to reduce their ef-R43 fect on climate change in the future^{2,3}. R44

Ozone itself is a greenhouse gas, and is continually produced and destroyed in the atmosphere by chemical reactions. In the troposphere, the lowest 5 to 10 miles of the atmosphere near the surface, hu-

2,000 Years of Greenhouse Gas Concentrations





L1 man activities have increased ozone concentration 1.2 through the release of gases such as carbon mon-L3 oxide, hydrocarbons, and nitrogen oxides. These IA gases undergo chemical reactions to produce ozone L5 in the presence of sunlight. In addition to trapping heat, excess ozone in the troposphere causes respi-L6 L7 ratory illnesses and other human health problems. In the stratosphere, the layer above the troposphere, 1.8 19 ozone exists naturally and protects life on Earth L10 from exposure to excessive ultraviolet radiation L11 from the Sun. As mentioned previously, halocar-L12 bons released by human activities destroy ozone L13 in the stratosphere and have caused the ozone hole L14 over Antarctica. Changes in the stratospheric ozone layer have contributed to changes in wind patterns L15 and regional climates. L16

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L18 *Water vapor* is the most important and abundant L19 greenhouse gas in the atmosphere. Human activi-L20 ties produce only a small increase in water vapor through combustion processes and irrigation. L21 L22 However, the surface warming caused by human-L23 produced increases in other greenhouse gases leads L24 to a large increase in water vapor, since a warmer L25 climate increases evaporation and allows the atmo-L26 sphere to hold more moisture. This in turn leads to L27 more warming, creating a "feedback loop".

L29 Other human influences

L30 In addition to the global-scale climate effects of L31 heat-trapping gases, human activities also produce L32 additional local and regional effects. Some of these L33 activities partially offset the warming caused by L34 greenhouse gases, while others increase the warming. One such influence on climate is caused by L35 L36 tiny particles called "aerosols" (not to be confused L37 with aerosol spray cans). For example, the burning L38 of coal produces emissions of sulfur-containing L39 compounds. These compounds form "sulfate aerosol" particles, which reflect some of the incoming L40 L41 sunlight away from the Earth, thus leading to local L42 or regional cooling influence. Sulfate aerosols also I.43 tend to make clouds more efficient at reflecting L44 sunlight, causing an additional indirect cooling effect. Another type of aerosol, often referred to I.45 L46 as soot or black carbon, absorbs incoming sunlight L47 and traps heat in the atmosphere. Thus, depending L48 on their type, aerosols can either mask or increase I.49 the warming caused by increased levels of green-L50 house gases. At the global scale, the sum of these

aerosol effects offsets some of the warming caused by heat-trapping gases and, in some locations with large amounts of aerosol particles, can even cause a net cooling.

The effects of various greenhouse gases and aerosol particles on Earth's climate depend in part on how long these gases and particles remain in the atmosphere. After emission, the atmospheric concentration of carbon dioxide remains elevated for many centuries, while the elevated concentrations of aerosols and methane would persist for only days to decades if emissions were reduced. Reductions in some of these shorter-lived gases and particles can thus have relatively rapid and potentially complex effects on climate^{4,5}. In contrast, while the concentrations of carbon dioxide and other longlived gases go up rapidly after their emission, the climate effects of reductions in their emissions will not become apparent for at least several decades.

Human activities have also changed the land surface in ways that alter how much heat is reflected or absorbed by the surface. Such changes include the cutting and burning of forests, the replacement of other areas of natural vegetation with agriculture and cities, and large-scale irrigation. These transformations of the land surface can cause local (and even regional) warming or cooling. Globally, the net effect of these changes has probably been a slight cooling of the Earth's surface over the past 100 years^{6,7}.

Natural influences

Two important natural factors also influence climate: the Sun and volcanic eruptions. Over the past three decades, human influences on climate have become increasingly obvious, and global temperatures have risen sharply. During the same period, the Sun's energy output (as measured by satellites since 1979) has followed its historic 11-year cycle of small ups and downs, but with no net increase⁸. The two major volcanic eruptions of the past 30 years have had short-term cooling effects on climate, lasting 2 to 3 years⁵. Thus, these natural factors cannot explain the warming of recent decades; in fact, their net effect on climate has probably been a slight cooling influence over this period. Slow changes in Earth's orbit around the Sun and its tilt toward or away from the Sun are also a purely

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Global Climate Change Impacts in the United States

natural influence on climate, but are only important on timescales from thousands to many tens of thousands of years.

> The climate changes that have occurred over the last century are not solely caused by the human and natural factors described above. In addition to these influences, there are also purely natural fluctuations in climate (often called "climate noise") that occur even in the absence of changes in human activities, the Sun, or volcanoes. One example is the El Niño phenomenon, which has important influences on many aspects of regional and global climate. Many other modes of natural internal variability have been identified by climate scientists and their

effects on climate occur at the same time as the effects of human activities, the Sun, and volcanoes.

Carbon release and uptake

Once carbon dioxide is emitted to the atmosphere, some of it is absorbed by the oceans and by vegetation on land; about 45 percent of the carbon dioxide emitted by human activities in the last 50 years has been taken up by these natural "sinks". The rest has remained in the air, increasing the atmospheric concentration^{1,2,9}. It is thus important to understand not only how much carbon dioxide is emitted, but also how much is taken up, over what time scales, and how these sources and sinks of carbon dioxide might change as climate continues to warm.



The figure above shows the amount of warming influence (red bars) or cooling influence (blue bars) that different factors have had on Earth's climate over the industrial age (from about 1750 to the present). Results are in watts per square meter. The longer the bar, the greater the influence on climate. The top part of the box includes all the major human-induced factors, while the second part of the box includes the Sun, the only major natural factor with a long-term effect on climate. The cooling effect of individual volcanoes is also natural, but is relatively short-lived (2 to 3 years). The bottom part of the box shows that the total net effect of human activities is a strong warming influence. The thin lines on each bar provide an estimate of the range of uncertainty.

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Global Climate Change

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L1 The rate of rise in global emissions of carbon diox-12 ide has been accelerating. The growth rate increased L3 from 1.3 percent per year in the 1990s to 3.3 percent IA per year between 2000 and 2006¹⁰. The increasing 1.5 emissions of carbon dioxide have clearly contributed L6 to the observed increased concentration of carbon di-L7 oxide in the atmosphere, but are perhaps not the only factor. There is some evidence that a recent decrease 1.8 19 in the rate of uptake of carbon dioxide by the oceans L10 and by land vegetation contributed to the observed L11 increased carbon dioxide concentration in the atmo-L12

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Over the last 100 years, global average L15 temperature and sea level have increased, L16 and precipitation patterns have changed. L17

L19 Temperatures are rising

L20 Global average surface air temperature has been increasing rapidly since 1970¹². The estimated change L21 L22 in the average temperature of Earth's surface is based L23 on measurements made by satellites and at thousands L24 of weather stations, ships, and buoys around the L25 world. These measurements are independently com-L26 piled, analyzed, and processed by different research L27 groups. An important step in the data processing is to identify and adjust for the effects of changes in the L28 1.29 instruments used to measure temperature, the mea-L30 surement times and locations, and the local environ-L31 ment around the measuring site (such as the growth of cities, and the development of so-called "urban heat L32 island" effects) or within a satellite's field of view. L33 L34 A number of research groups around the world have L35 produced estimates of global-scale changes in surface L36 temperature. L37

L38 The warming trend that is apparent in all of these L39 temperature records is confirmed by other indepen-I 40 dent observations, such as the melting of Arctic sea L41 ice, the retreat of mountain glaciers on every conti-L42 nent¹³, reductions in the extent of snow cover, earlier I.43 blooming of plants in spring, and increased melting of L44 the Greenland and Antarctic ice sheets¹⁴.

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L46 Additionally, temperature measurements above the surface have been made by weather balloons since L47 L48 the late 1940s, and from satellites since 1979. These I.49 measurements show warming of the troposphere, consistent with the surface warming^{16,17}. They also L50



Global Temperature and CO₂

Layers of the Atmosphere

Closest to the Earth's Surface

Stratosphere

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Temperature (°F)

The illustration shows the layers of the atmosphere

closest to Earth's surface. The troposphere extends

from the surface up to roughly 6 miles above the surface

and the stratosphere is above that. The colored band

shows the average temperature of the atmosphere at

different altitudes. In the troposphere, temperatures

generally decrease with height, while in the stratosphere

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Troposphere

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temperatures increase with height.

Pressure (millibars)



Global annual average temperature (as measured over both land and oceans). Red bars indicate temperatures above the 1901-2000 average, blue bars are below average temperatures. The black line shows carbon dioxide concentration. While there is a clear longterm global warming trend, each individual year does not show a temperature increase relative to the previous year, and some years show greater changes than others. These year-to-year fluctuations in temperature are due to natural processes, such as the effects of El Niños and La Niñas.

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reveal cooling in the stratosphere¹⁶. This pattern of tropospheric warming and stratospheric cooling agrees with our understanding of how atmospheric temperature would be expected to change in response to increasing greenhouse gas concentrations and the observed depletion of stratospheric ozone⁶.

Precipitation patterns are changing

Precipitation is not distributed evenly over the globe. Its average distribution is governed primarily by atmospheric circulation patterns and the availability of moisture, which in turn are influenced by temperature. Because of human-caused changes in atmospheric temperature, changes are expected in atmospheric circulation, and therefore in precipitation patterns.

Observations show that such shifts are occurring. Changes have been observed in the amount, intensity, frequency, and type of precipitation. Pronounced increases in precipitation over the past 100 years have been observed in eastern North America, southern South America, and northern Europe. Decreases have been seen in the Mediterranean, most of Africa, and southern Asia. The geographical distribution of droughts and flooding has been complex. In some regions, there have been increases in the occurrences of both droughts and floods¹⁴. As the world warms, northern regions and mountainous areas are experiencing more precipitation falling as rain rather than snow¹⁸. Widespread

increases in heavy precipitation events have occurred, even in places where total amounts have decreased. These changes are associated with the fact that warmer air holds more water vapor evaporating from the world's oceans and land surface¹⁷. This increase in atmospheric water vapor has been observed from satellites, and is primarily due to human influences19,20.

Sea level is rising

I.45 After at least 2000 years of little change, sea level rose by roughly L46 8 inches over the past 100 years. L47 L48 Satellite data available over the past I 49 15 years shows sea-level rising at a L50

rate roughly double the rate observed over the past century²¹.

Global warming causes sea level to rise in two ways. First, ocean water expands as it warms, and therefore takes up more space. Warming has been observed in each of the world's major ocean basins, and has been directly linked to human influences^{22,23}.

Second, warming leads to the melting of glaciers and ice sheets, which raises sea level by adding water to the oceans. Glaciers have been retreating worldwide, and the rate of retreat has increased in the past decade²⁴. Only a few glaciers are actually advancing (in locations that were well below freezing, and where increased precipitation has outpaced melting). The total volume of glaciers on Earth is declining sharply. The progressive disappearance of glaciers has implications not only for the rise in global sea level, but also for water supplies in certain densely-populated regions of Asia and South America.

The Earth has two major ice sheets. The Greenland Ice Sheet contains enough water to raise sea level by about 20 feet. Melting of the entire Antarctic Ice Sheet would raise sea levels by over 200 feet. Both of these ice sheets are currently melting around parts of their edges. Complete melting of either of these ice sheets over this century or the next is

Cumulative Decrease in Global Glacier Ice



virtually impossible. The Greenland Ice Sheet has
also been experiencing record amounts of surface
melting, and a large increase in the rate of mass loss
in the past decade²⁵.

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L12 In 1996, the IPCC Second Assessment Report²⁶ L13 cautiously concluded that "the balance of evi-L14 dence suggests a discernible human influence on global climate". Since then, a number of national L15 and international assessments have come to much L16 L17 stronger conclusions about the reality of human effects on climate. Recent scientific assessments L18 find that most of the warming of the Earth's surface L19 over the past 50 years has been caused by human L20 activities^{27,28}. What evidence allowed scientists to L21 L22 identify human influences as the major cause of the observed warming? How can we be sure that "it's L23 L24 mostly us"? L25

L26 This conclusion rests on multiple lines of evidence. L27 Like the warming "signal" that has gradually emerged from the "noise" of natural climate vari-L28 ability, the scientific evidence for a human influ-1.29 L30 ence on global climate has accumulated slowly over L31 the past several decades, from many hundreds of L32 studies. No single study is a "smoking gun". Nor L33 has any single study undermined the large body of evidence supporting the conclusion that human L34 activity is the primary driver of recent warming. L35 L36 L37 The first line of evidence is our basic physical un-L38 derstanding of how greenhouse gases trap heat, how the climate system responds to increases in green-L39 house gases, and how other human and natural L40 L41 factors influence climate. The second line of evi-L42 dence is from indirect estimates of climate changes I.43 over the last 1,000 to 2,000 years. These so-called L44 "paleodata" are obtained from living things (like I.45 tree rings and corals) and from physical quantities L46 (like the ratio between lighter and heavier isotopes L47 of oxygen in ice cores) which change in measurable ways as climate changes. The lesson from paleo-L48 data is that global surface temperatures over the I.49 L50 last several decades are clearly unusual, in that they

were higher than at any time during at least the past 400 years²⁹. For the Northern Hemisphere, recent temperature rises are clearly unusual in at least the last 1,000 years^{29,30}.

The third line of evidence is based on the broad, qualitative consistency between observed changes in climate and the computer model predictions of how climate would be expected to change in response to human activities. For example, when climate models are run with historical increases in greenhouse gases, they show gradual warming of the Earth and ocean surface, increases in ocean heat content and the temperature of the lower atmosphere, a rise in global sea level, retreat of sea-ice and snow cover, cooling of the stratosphere, an increase in the amount of atmospheric water vapor, and changes in large-scale precipitation and pressure patterns. These and other aspects of modeled climate change are in agreement with observations^{6,31}.

Finally, there is statistical evidence from so-called "fingerprint" studies. Each factor that affects climate produces a unique pattern of climate response, much as each person has a unique fingerprint. Fingerprint studies exploit these unique signatures, and make detailed comparisons of modeled and observed climate change patterns³¹. Scientists rely on such studies to attribute observed changes in climate to a particular cause or set of causes. In the real world, the climate changes that have occurred since the Industrial Revolution are due to a complex mixture of human and natural causes. The importance of each individual influence in this mixture changes over time. Of course, there are not multiple Earths, which would allow an experimenter to change one factor at a time on each Earth, thus helping to isolate different fingerprints. Climate models can be used to perform the systematic experiments that are not possible in the real world: a single factor (like greenhouse gases) or a set of factors can be varied, and the response of the climate system to these individual or combined changes can thus be studied³².

For example, when climate model simulations of the last century include all of the major influences on climate, both human-induced and natural, they can reproduce many important features of observed R5 R6 R7 R8

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Separating Human and Natural Influences on Climate with human effects 58 Temperature (°F) observed 57 56 natural forces only 1900 1950 2000 Year Observations Models using only natural forces Models using both natural and human forces Hegerl et al.³¹

The blue band shows how global average temperatures would have changed due to natural forces only, as simulated by climate models. The red band shows model projections of the effects of human and natural forces combined. The black line shows actual observed global average temperatures. As the blue line indicates, without human influences, temperature over the past century would actually have first warmed and then cooled slightly over recent decades.

climate change patterns. When human influences are removed from the model experiments, results suggest that the surface of the Earth would actually have cooled slightly over the last 50 years. The clear message from fingerprint studies is that the observed warming over the last half-century cannot be explained by natural factors alone^{6,33}.

L31 Another fingerprint of human effects on climate L32 has been identified when one looks at a slice through the layers of the atmosphere, and studies L33 L34 the pattern of temperature changes from the surface L35 up through the stratosphere. In all climate models, L36 increases in carbon dioxide cause warming at the surface and in the troposphere, but lead to cool-L37 L38 ing of the stratosphere. Models also show that the L39 human-caused depletion of stratospheric ozone has L40 a strong cooling effect in the stratosphere. There L41 is a good match between the model fingerprint in L42 response to combined carbon dioxide and ozone changes and the observed pattern of tropospheric L43 L44 warming and stratospheric cooling⁶. L45

L46 In contrast, if most of the observed temperature
L47 change had been due to an increase in solar outL48 put rather than an increase in greenhouse gases,
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L50 its full vertical extent, including the stratosphere⁶.

Global Climate Change Impacts in the United States

The observed pattern of atmospheric temperatureR1changes, with its pronounced cooling in the strato-R2sphere, is therefore inconsistent with the hypothesisR3that changes in the Sun can explain the warming ofR4recent decades. Moreover, direct satellite measure-R5ments of solar output show slight decreases duringR6the recent period of warming.R7

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The earliest fingerprint work³⁴ focused on changes R9 in surface and atmospheric temperature. Scientists R10 then applied fingerprint methods to a whole range R11 of climate variables^{31,35}, identifying human-caused R12 climate signals in the heat content of the oceans^{22,23}, R13 the height of the tropopause³⁶ (the boundary be-R14 tween the troposphere and stratosphere, which has R15 shifted upward by hundreds of feet in recent de-R16 cades), the geographical patterns of precipitation³⁷, R17 drought³⁸, surface pressure³⁹, and the runoff from R18 major river basins⁴⁰. R19

Studies published after the appearance of the IPCC R21 Fourth Assessment Report in 2007 have found hu-R22 man fingerprints in the increased levels of atmo-R23 spheric moisture^{19,20} (both close to the surface and R24 over the full extent of the atmosphere), in the de-R25 cline of Arctic sea ice extent⁴¹, and in the patterns R26 of changes in Arctic and Antarctic surface tempera-R27 tures⁴². The message from this entire body of work R28 is that the climate system is telling a consistent R29 story of increasingly dominant human influence-R30 the changes in temperature, ice extent, moisture, R31 and circulation patterns fit together in a physically R32 consistent way, like pieces in a complex puzzle. R33

Increasingly, this type of fingerprint work is R35 shifting its emphasis. As noted, clear and compel-R36 ling scientific evidence supports the case for a R37 pronounced human influence on global climate. R38 Much of the recent attention is now on climate R39 changes at continental and regional scales^{43,44}, and R40 on variables that can have large impacts on societ-R41 ies. For example, scientists have established causal R42 links between human activities and the changes in R43 snowpack, maximum and minimum temperature, R44 and the seasonal timing of runoff over mountain-R45 ous regions of the western United States¹⁸. A large R46 human component has been identified in the ocean R47 surface temperature changes in hurricane formation R48 regions^{45,46}. Researchers are also looking beyond R49 the physical climate system, and are beginning to R50

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L1tie changes in the distribution and seasonal behav-12 ior of plant and animal species to human-caused L3 changes in temperature and precipitation^{47,48}. IA 1.5 For over a decade, one aspect of the climate change L6 story seemed to show a significant difference L7 between models and observations⁶. In the tropics, 1.8 all models predicted that with a rise in greenhouse 19 gases, the troposphere would be expected to warm L10 more rapidly than the surface. Observations from L11 weather balloons, satellites, and surface thermom-L12 eters seemed to show exactly the opposite behav-L13 ior (more rapid warming of the surface than the L14 troposphere). This issue was a stumbling block in L15 our understanding of the causes of climate change. It is now largely resolved⁴⁹. Research showed that L16 L17 there were large uncertainties in the satellite and weather balloon data. When uncertainties in mod-L18 L19 els and observations are properly accounted for,

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newer observational datasets (with better treatment of known problems) are in agreement with climate model results^{17,50-53}.

This does not mean, however, that all remaining differences between models and observations have been resolved. The observed changes in some climate variables, such as Arctic sea ice⁴¹, some aspects of precipitation^{37,55}, and patterns of surface pressure, appear to be proceeding much more rapidly than models have projected. The reasons for these differences are not well understood. Nevertheless, the bottom-line conclusion from climate fingerprinting is that most of the observed changes studied to date are consistent with each other, and are also consistent with our scientific understanding of how the climate system would be expected to respond to the increase in heat-trapping gases resulting from human activities^{6,31}.



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L1 Scientists are sometimes asked whether extreme 1.2 weather events can be linked to human activities⁵⁶. L3 Scientific research has concluded that human influ-IA ences on climate are indeed changing the likelihood 1.5 of certain types of extreme events. For example, an analysis of the European summer heat wave of L6 L7 2003 found that the risk of such a heat wave is now 1.8 roughly four times as great due to human influences on climate^{57,58}.

> Like fingerprint work, such analyses of humancaused changes in the risks of extreme events rely on information from climate models, and on our understanding of the physics of the climate system. All of the models used in this work have imperfections in their representation of the complexities of the "real world" climate system^{59,60}. These are due to both limits in our understanding of the climate system, and in our ability to represent its complex behavior with available computer resources. Despite this, models are extremely useful, for a number of reasons.

L24 First, despite the existence of systematic errors, the L25 current generation of climate models accurately L26 portrays many important aspects of today's weather L27 patterns and climate^{59,60}. Models are constantly being improved, and are routinely tested against L28 1.29 many observations of Earth's climate system. L30 Second, the fingerprint work shows that models L31 capture not only our present-day climate, but also L32 key features of the observed climate changes over L33 the past century²⁹. Third, many of the large-scale L34 observed climate changes (such as the warming of L35 the surface and troposphere, and the increase in the L36 amount of moisture in the atmosphere) are driven L37 by very basic physics, which is well-represented L38 in models¹⁹. Fourth, climate models can be used to L39 predict changes in climate that can be verified in L40 the real world. Examples include the global cooling L41 subsequent to the eruption of Mount Pinatubo and L42 the stratospheric cooling with increasing carbon L43 dioxide. Finally, models are the only tools that exist L44 for trying to understand the climate changes likely L45 to be experienced over the course of this century. L46 No period in Earth's geological history provides an L47 exact analogue for the climatic conditions that will L48 unfold in the coming decades.

Global temperatures will continue to rise over this century; by how much and for how long depends on a number of factors, including the amount of heattrapping emissions and how sensitive the climate is to those emissions. R1

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Some continued warming of the planet is inevitable over the next few decades. The amount of future warming will be determined largely by choices made now and over the next few decades. Lower levels of heat-trapping emissions will yield less future warming, while higher levels will result in more warming, and more severe impacts on society and the natural world.

Rising global temperature

All climate models project that human-caused emissions of heat-trapping gases will cause further warming in the future. Based on scenarios that do not assume explicit climate policies to reduce greenhouse gas emissions, global average temperature is projected to rise by 2 to 11.5°F by the end of this century⁶¹ (relative to the 1980-1999 time period). Whether the actual warming in 2100 will be closer to the low or the high end of this range depends primarily on two factors: first, the future level of emissions of heat-trapping gases. and second, how sensitive climate will be, that is, how much climate will change in response to those emissions. The range of possible outcomes has been explored using a range of different emissions scenarios, and a variety of climate models that encompass the known range of climate sensitivity.

The IPCC developed a set of scenarios in a Special R36 Report on Emissions Scenarios (SRES)⁶². These R37 have been extensively used to explore the potential R38 for future climate change. None of these scenarios R39 assumes explicit policies to limit climate change. R40 Rather, emissions in these scenarios vary based on R41 different assumptions about changes in population, R42 adoption of new technologies, economic growth, R43 and other factors. None of them involve stabilizing R44 atmospheric concentrations of heat-trapping gases R45 at a level that would avoid dangerous human inter-R46 ference with the climate system as required by the R47 United Nations' Framework Convention on Climate R48 Change, which was signed in 1992 by the United R49 States and most other countries. R50

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L1 Changing precipitation patterns

1.2 Projections of changes in precipitation largely follow recently observed patterns of change, with L3 IA overall increases in the global average but substan-L5 tial shifts in where and how precipitation falls⁶¹. Generally, higher latitudes are projected to receive L6 L7 more precipitation, while the sub-tropics expand further poleward⁶³ and also receive less rain. 1.8 19 Increases in tropical precipitation are projected during rainy seasons (such as monsoons), and es-L10 L11 pecially over the tropical Pacific. Certain regions, L12 including the U.S. West (especially the Southwest) L13 and the Mediterranean, are expected to become L14 drier. The trend towards more heavy downpours is expected to continue, with precipitation becoming L15 less frequent but more intense⁶¹. More precipitation L16 L17 is expected to fall as rain rather than snow.

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L19 Currently rare extreme events are becoming L20 more common

In a warmer future climate, models project there L21 L22 will be an increased risk of more intense, more L23 frequent and longer-lasting heat waves⁶¹. The Eu-L24 ropean heat wave of 2003 is an example of the type of extreme heat event that is likely to become more L25 L26 common⁶¹, with the likelihood of such a heat wave L27 projected to increase 100-fold in the next 40 years. If greenhouse gas emissions continue to increase, L28 1.29 by the 2040s more than half of European summers L30 will be hotter than the summer of 2003, and by the L31 end of this century, a summer as hot as that of 2003 L32 will be considered unusually cool⁵⁷. L33 L34 Increased extremes of summer dryness and winter wetness are projected for much of the globe, mean-L35 L36 ing a generally greater risk of droughts and floods. This has already been observed³⁸, and is projected L37

L38 to continue, because in a warmer world, precipitaL39 tion tends to be concentrated into more intense
L40 events, with longer periods of little precipitation in
L41 between⁶¹.

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L43 Models project a general tendency for more intense
L44 but fewer storms overall outside the tropics, with
L45 more extreme wind events and higher ocean waves
L46 in a number of regions in association with those
L47 storms. Models also project a shift of storm tracks
L48 toward the poles in both hemispheres⁶¹.

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Changes in hurricanes are difficult to project because there are countervailing forces. Higher ocean temperatures lead to stronger storms with higher wind speeds and more rainfall⁶⁴. But changes in wind speed and direction with height are also projected to increase in some regions, and this tends to work against storm formation and growth⁶⁵. It currently appears that stronger, more rain-producing tropical storms and hurricanes are generally more likely, though more research is required on these issues.

Sea level will continue to rise

Projecting future sea-level rise presents special challenges. Scientists have a well-developed understanding of the contributions of thermal expansion and melting glaciers to sea-level rise, so the models used to project sea-level rise include these processes. However, recent observations of the polar ice sheets show that additional processes are operating that affect the responses of ice sheets to warming. Although these processes are not well understood, they are already producing substantial additional loss of ice mass, but it is difficult to predict their future contributions to sea-level rise.

Thus, most current estimates offer only a likely lower bound for future sea-level rise projections, with a highly uncertain upper bound. The 2007 assessment by the IPCC, for example, which did not attempt to include the highly uncertain contributions to sea-level rise due to changes in ice sheet dynamics, projected a rise of the world's oceans from 8 inches to 2 feet by the end of this century⁶¹.

Recent research has led to more comprehensive estimates of the accelerated flow to the sea of ice sheets in a warmer climate and how this contributes to sea-level rise. This work suggests that the upper and lower limits on sea-level rise over this century are substantially greater than previously projected^{13,66-68}.

The changes in sea level experienced at any particular location along the coast depend not only on the increase in the global average sea level, but also on changes in regional currents and winds and, particularly, on the vertical movements of the land due to geological forces. The consequences of sea-level rise at any particular location depend on L1 L2 L3 L4 L5

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Global Increase in Heavy Precipitation



Observed and projected changes in the heaviest 5 percent of precipitation events. The shaded areas show the possible ranges while the lines show the central projections from a set of climate models.

Observed and Projected Global Average Temperature





Emissions scenarios

The IPCC emission scenarios do not encompass the
full range of possible futures: climate can change
less than those scenarios imply, or it can change
more. Current carbon dioxide emissions are, inR2
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fact, above the highest emissions scenario[†] developed by the IPCC⁷⁰ (see figure on page 25). Whether this will continue is uncertain. R1

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There are also lower possible emis-R11 sions paths than those put forth by R12 the IPCC. The Framework Conven-R13 tion on Climate Change, to which R14 the United States and most other R15 countries are signatories, calls for R16 stabilizing concentrations of green-R17 house gases in the atmosphere at a R18 level that would avoid dangerous R19 human interference with the cli-R20 mate system. What exactly consti-R21 tutes such interference is subject to R22 interpretation. R23

A variety of research studies sug-R25 gest that a further 2°F increase R26 (relative to the 1980-1999 period) R27 would lead to severe, widespread, R28 and irreversible impacts⁷¹⁻⁷³. To R29 have a good chance (but not a R30 guarantee) of avoiding tempera-R31 R32 tures above those levels, it has been estimated that atmospheric concen-R33 trations of carbon dioxide would R34 need to stabilize in the long term at R35 around today's levels74-77. R36

The graphs above show emis-R38 sions scenarios and resulting CO₂ R39 concentrations for three IPCC R40 scenarios^{†,61} and two stabilization R41 scenarios78. The stabilization sce-R42 narios are aimed at stabilizing at-R43 mospheric CO₂ at roughly 450 and R44 550 parts per million (ppm); this R45 is 70 to 170 ppm above the current R46 concentration of about 380 ppm. R47 Resulting temperature changes R48 depend on the level of CO_2 , how R49 sensitive the climate system is, and R50

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The graphs show recent and projected global emissions of carbon dioxide in gigatons of carbon, on the left, and atmospheric concentrations on the right under five emissions scenarios. The top three in the key are IPCC scenarios that assume no explicit climate policies (these are used in model projections that appear throughout this report). The bottom two are "stabilization scenarios," designed to stabilize atmospheric carbon dioxide concentrations at 450 or 550 parts per million. The inset expanded below these charts shows emissions for the current two decades under these five scenarios along with actual emissions (in black).

the amount of particles in the atmosphere⁷⁵. Only
the 450 ppm stabilization target has the potential to
keep the global temperature rise at or below about
3.5°F from pre-industrial and 2°F above current,
a level beyond which many concerns have been
raised about dangerous human interference with the
climate system^{76,77}.

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Carbon (Gigatons)

L41 A further complication is that carbon dioxide is not L42 the only greenhouse gas of concern. Concentrations of other heat-trapping gases like methane and I.43 L44 nitrous oxide and particles like soot will also have I.45 to be stabilized at low enough levels to prevent global temperatures from rising higher than the L46 level mentioned above. When these other gases L47 L48 are added, including the offsetting cooling effects I.49 of sulfate aerosol particles, analyses suggest that L50 stabilizing concentrations around 400 parts per

million of equivalent CO_2 would yield about an 80 percent chance of avoiding exceeding the 2°F above present temperature threshold. This would be true even if concentrations temporarily peaked as high as 475 parts per million and then stabilized at 400 parts per million roughly a century later^{50,69,76,77,80.81}.

Rapid climate change

There is also the possibility of even larger climate change than current scenarios and models project. Not all changes in the climate are gradual. The long record of climate found in ice cores, tree rings, and other natural records show that Earth's climate patterns have undergone rapid shifts from one stable state to another within as short a period as a decade. The occurrence of rapid climate changes becomes increasingly more likely as the human disturbance of the climate system grows⁶¹. Such R28

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changes can occur so rapidly that they would challenge the ability of human and natural systems to adapt⁸². Examples of such changes are rapid shifts in drought frequency and duration. Ancient climate records suggest that in the United States, the Southwest may be at greatest risk for this kind of change, but that other regions including the Midwest and Great Plains have also had these kinds of rapid shifts in the past and could experience them again in the future.

Rapid ice sheet collapse with related sea-level rise is another type of rapid change that is not well understood or modeled that poses a risk for the future. Recent observations show that melting on the surface of an ice sheet produces water that flows down through large cracks that create conduits through the ice to the base of the ice sheet where it lubricates ice previously frozen to the rock below⁸². Further, the interaction with warm ocean water, where ice meets the sea, can lead to sudden losses in ice mass and accompanying rapid global sealevel rise. Observations indicate that ice loss has increased dramatically over the last decade, though scientists are not yet confident that they can project how the ice sheets will respond in the future.

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