

Constraints on the Age and Provenance of the Chugach Accretionary Complex from Detrital Zircons in the Sitka Graywacke near Sitka, Alaska

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Abstract

The Sitka Graywacke is the westernmost and youngest unit of the Chugach accretionary complex in southeastern Alaska. Using laser-ablation inductively coupled plasma mass spectroscopy, we obtained 492 detrital-zircon ages on seven typical samples of Sitka Graywacke turbidites, which were collected in a transect across much of the unit near Sitka, Alaska. Individual grains range in age from 66 to 1,802 m.y. The youngest peak ages on relative-probability plots of the western four samples (74, 72, 74, and 74 m.y., from west to east) are distinctly younger than the youngest peak ages of the eastern three samples (105, 103, and 97 m.y., from west to east). These youngest peak ages set maximum depositional ages for each sample. We suggest that these peak ages are not significantly older (within ~5 m.y.) than the depositional age of the Sitka Graywacke because the deposits accumulated in a trench along a convergent margin, where magmatic sources likely continuously introduced juvenile zircons. The differences in the youngest cluster of detrital-zircon ages between the eastern and western sample localities is likely due to both a change in provenance and a fault. The similarity of the youngest peak ages in the Sitka Graywacke to fossil ages in the Valdez Group, in Prince William Sound, implies that the western part of the Sitka Graywacke is correlative with the Valdez Group, as previously inferred. However, the eastern part of the Sitka Graywacke has youngest detrital-zircon ages older than fossil ages in the Valdez Group and younger than fossil ages in the McHugh Complex, which in south-central Alaska is the oldest part of the accretionary complex. The age distribution of zircons in the older, eastern sequence suggests sources along the British Columbia margin. The detrital-zircon ages in the younger, western sequence are similar to igneous ages from south-central Alaska to southern British Columbia. Right-lateral strike slip on various fault systems inboard of the Sitka Graywacke implies that it lay to the south when it was deposited and offscraped. Thus, although source areas as far north as the St. Elias Mountains and south-central Alaska are possible, they were most likely in coastal and interior British Columbia.

Introduction

Graywacke turbidites are notoriously poor in fossils, and so dating turbidite sequences can be a great challenge. The Sitka Graywacke, which lies along the west coast of southeastern Alaska, is one such turbidite sequence (Berg and Hinckley, 1963; Loney and others, 1963; Decker and others, 1979; Decker, 1980). The unit is part of a Mesozoic through early Tertiary accretionary complex rimming southern Alaska (fig. 1; Plafker and others, 1977) that is commonly referred to as the Chugach terrane (for example, Plafker and others, 1994). Various workers (Nilsen and Moore, 1979; Nilsen and Zuffa, 1982; Plafker and others, 1994) have proposed that these accretionary-complex turbidites were deposited by northward-flowing trench-parallel turbidity currents derived from erosion of the Coast Mountains of southeastern Alaska and British Columbia.

Regional correlations between the Sitka Graywacke and other units of the Chugach accretionary complex are hindered by the absence of fossils in the turbidites. Thus far, only two fossils have been identified from the Sitka Graywacke. Reifentstahl (1986) obtained a gastropod with a nondiagnostic age ranging from Silurian to Eocene. Also, a small limestone pod, with sharp contacts, in graywacke (at the location of our sample 65, fig. 2) contained a poorly preserved cone-shaped nassellarian radiolarian of possible Jurassic age (C.D. Blome, written commun., 1994).

Detrital-zircon ages can provide a maximum depositional age because the host sediment must have been laid down after the youngest detrital-zircon grain crystallized. In this study, we use the youngest ages from sandstone of the Sitka Graywacke to constrain depositional ages, subdivide the unit, and interpret the provenance of the detrital-zircon grains on the basis of a comparison with the ages of potential igneous source rocks along the northern Cordilleran margin.

Methods

We collected seven samples in an east-west-trending transect across much of the Sitka Graywacke near Sitka, Alaska (fig. 2; table 1). All samples were collected from normally

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graded sandy turbidite beds, 5 to 20 cm thick, with sharp basal contacts. The sampled beds were typical of the beds at each sample locality. Detrital zircons were extracted by using standard mineral-separation techniques.

Zircons were mounted in epoxy and polished to expose the interior of the grains. Isotopic analyses were performed by using a Micromass IsoProbe multicollector inductively coupled plasma mass spectrometer (MC-ICPMS) with a laser-ablation system at the geochronology laboratory of the University of Arizona, Tucson. Laser-beam diameter was 35 μm , yielding ablation pits $\sim 15 \mu\text{m}$ deep. Interelement fractionation was monitored by analyzing fragments of a large concordant zircon with a known (isotope-dilution thermal-ionization mass-spectrographic, or ID-TIMS) age of 564 ± 4 m.y. (2σ error). This standard was analyzed once for every five unknowns. Grains were selected randomly from all sizes and morphologies present, avoiding grains with fractures or inclusions. The ablated material was carried in Ar gas into the plasma source of the mass spectrometer, which was equipped with a flight tube of sufficient width that U, Th, and Pb isotopes were analyzed simultaneously. All measurements were made in static mode, using Faraday detectors for ^{238}U , ^{232}Th , and $^{208-206}\text{Pb}$ and an ion-counting channel for ^{204}Pb . Ion yields were 0.5 mV/ppm. Each analysis consisted of 1 20-s integration on peaks with the laser off (for backgrounds), 20 1-s

integrations with the laser firing, and a 30-s delay to purge the previous sample and prepare for the next analysis.

Common Pb correction was made by using the measured ^{204}Pb content and assuming an initial Pb composition from Stacey and Kramers (1975) (with uncertainties of ± 1.0 for $^{206}\text{Pb}/^{204}\text{Pb}$ ratio and ± 0.3 for $^{207}\text{Pb}/^{204}\text{Pb}$ ratio). Our ^{204}Pb measurement is unaffected by the presence of ^{204}Hg because backgrounds are measured on peaks (thereby subtracting any background ^{204}Hg and ^{204}Pb) and only trace Hg was present in the Ar gas.

Interelement fractionation of Pb/U is generally < 20 percent, whereas fractionation of Pb isotopes is generally < 5 percent. In-run analysis, generally every fifth measurement, of fragments of a large zircon crystal standard (Sri Lanka) with a known age of 564 ± 4 m.y. (2σ error; G.E. Gehrels, unpub. data, 2004) was used to correct for this fractionation. The uncertainty resulting from the calibration correction (together with the uncertainty from decay constants and common Pb composition) is generally ± 3 percent (2σ) for $^{206}\text{Pb}/^{238}\text{U}$ ages and for $^{206}\text{Pb}/^{207}\text{Pb}$ ages > 1.2 b.y. Fractionation also increases with depth into the laser pit. The accepted isotopic ratios were accordingly determined by least-squares projection through the measured values back to the initial determination. The complete measured isotopic ratios and ages are listed in table 2. Errors from measuring $^{206}\text{Pb}/^{238}\text{U}$, $^{206}\text{Pb}/^{207}\text{Pb}$, and $^{206}\text{Pb}/^{204}\text{Pb}$

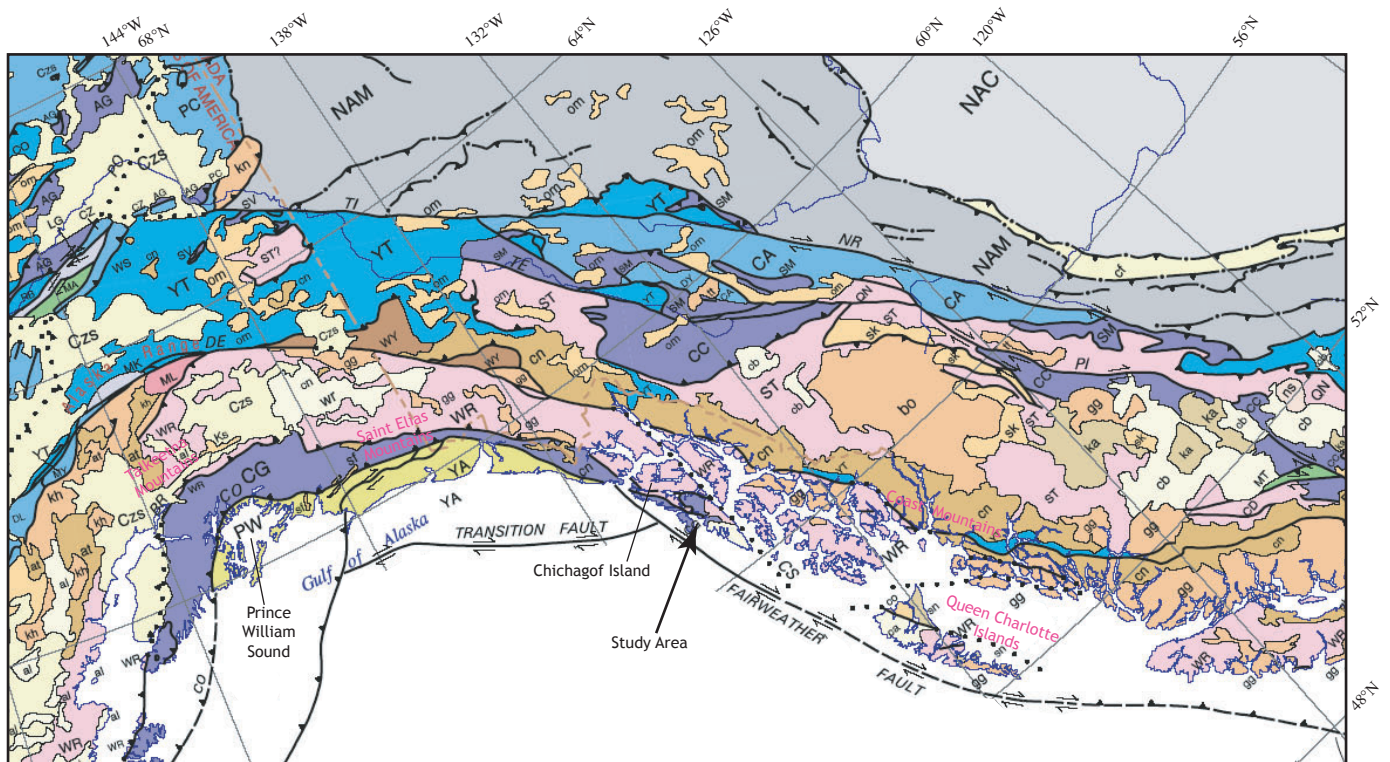


Figure 1. Gulf of Alaska area, showing location of study area and tectonostratigraphic terranes of Alaska, Yukon Territory, and British Columbia (from Nokleberg and others, 1998). The Sitka Graywacke, the Valdez Group, and the McHugh Complex are all parts of the Chugach accretionary complex. Terranes: CG, Chugach; ST, Stikine; YT, Yukon-Tanana. Faults: BR, Border Ranges; CS, Chatham Straight; DE, Denali; TI, Tintina.

Table 1. Summary data for samples of Sitka Graywacke sandstone analyzed for detrital zircons

[Latitude and longitude from NAD27 datum; youngest detrital-zircon age peaks from relative-probability plots in figure 3. USGS, U.S. Geological Survey]

Sample	Field No.	USGS map sheet	Latitude(° N.)	Longitude(° W.)	Youngest detrital-zircon age peak (Ma)
62	92PH062	Port Alexander D-5	56.93681	135.43308	74
63	92PH063	Port Alexander D-5	56.95279	135.39284	72
64	92PH064	Port Alexander D-5	56.97973	135.37554	74
65	92PH065	Port Alexander D-4	56.99984	135.32176	74
66	92PH066	Sitka A-4	57.01497	135.27063	105
67	92PH067	Sitka A-4	57.03072	135.23172	103
69	92PH069	Port Alexander D-4	56.99345	135.15739	97

ratios are reported at the 1σ level. Additional errors that affect all ages include uncertainties from (1) U decay constants, (2) the composition of common Pb (assumed to be ± 1.0 for the $^{206}\text{Pb}/^{204}\text{Pb}$ ratio and ± 0.3 for the $^{207}\text{Pb}/^{204}\text{Pb}$ ratio), and (3) calibration correction. These systematic errors add an additional 2-percent (1σ) uncertainty to $^{206}\text{Pb}/^{238}\text{U}$ ages and to $^{206}\text{Pb}/^{207}\text{Pb}$ ages > 1.2 b.y. A total of 69 to 75 ages were determined for each of the seven samples.

Age interpretations for the analyses are based largely on $^{206}\text{Pb}/^{238}\text{U}$ ages. $^{206}\text{Pb}/^{207}\text{Pb}$ ages for these relatively young zircons are much less reliable given the low ^{207}Pb content. As a result, discordance on an individual-grain basis cannot be assessed for grains younger than 1.2 b.y. Thus, $^{206}\text{Pb}/^{207}\text{Pb}$ ages were used for a total of four grains older than 1.2 b.y.

The U-Pb ages are plotted as histograms with a normalized relative-probability distribution in figure 3 (Ludwig, 2003), where the relative height of peaks corresponds to the significance of that age population.

Results and Discussion

We interpret the youngest cluster of detrital-zircon ages as setting a limiting depositional age for that sample. The youngest cluster is used, rather than the youngest individual grain, because isotopic disturbance by Pb loss can yield detrital-zircon ages that are significantly younger than the crystallization age but still analytically concordant. Because each zircon in a group of grains would almost certainly not lose exactly enough Pb to yield similar ages, we interpret the peak age of the youngest cluster (determined from the relative-probability plots in fig. 3) as a robust indicator of the age of the youngest igneous source, accordingly setting a maximum depositional age for the host graywacke.

The detrital-zircon ages indicate that the possible Jurassic nasselarian from the outcrop at sample locality 65, mentioned previously, is likely younger than Jurassic. The youngest peak

in ages at this locality is much younger—74 m.y. This detrital-zircon age helps constrain the loose fossil age.

The youngest peak ages of the western four samples (74, 72, 74, and 74 m.y., from west to east) are distinctly younger

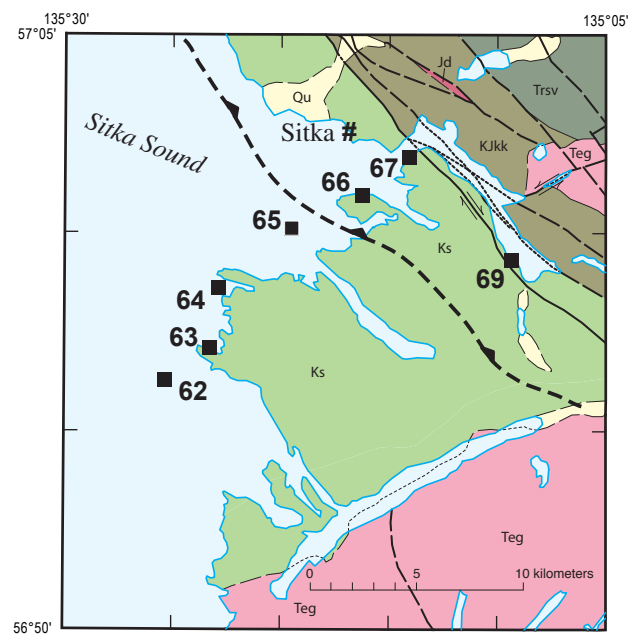


Figure 2. Study area near Sitka, Alaska, showing regional geology simplified from Karl and others (in press). Squares, location of detrital zircon samples. Samples 62 and 65 are on small islands hidden beneath symbol. Irregular dashed line, inferred thrust fault (teeth on upper plate) between samples 65 and 66 that separates older, eastern samples from younger, western samples. Units: Jd, Jurassic diorite; KJkk and Trsv, Jurassic through Cretaceous and Triassic parts of the Kelp Bay Group, respectively; Ks, Cretaceous Sitka Graywacke; Qu, Quaternary sedimentary rocks, undivided; Teg, Eocene granitic rocks. From U.S. Geological Survey Sitka A-4 and Port Alexander D-5 1:63,360-scale maps.

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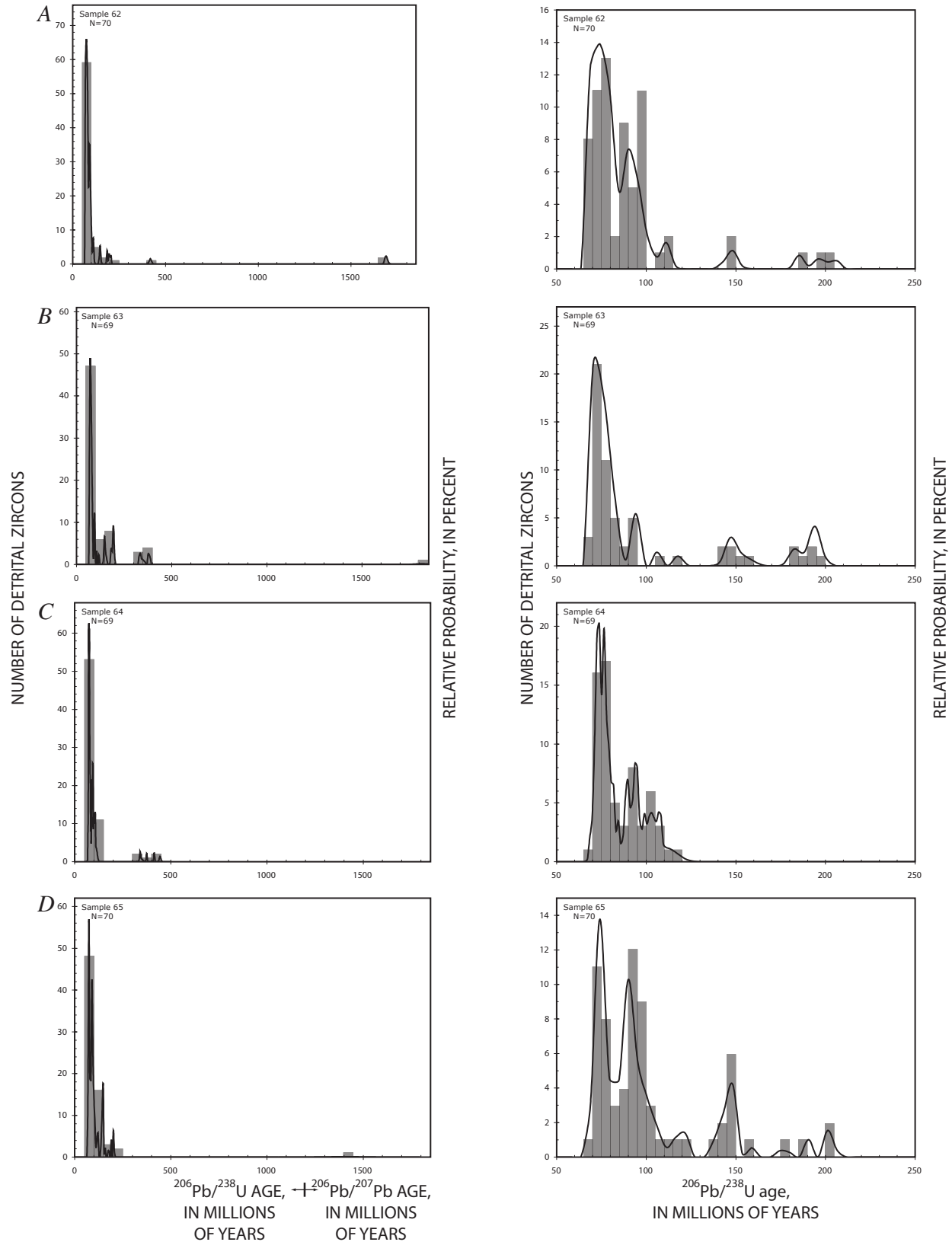


Figure 3. Detrital-zircon histograms (gray bars) and relative-probability plots (solid curve) for samples 62 (A), 63 (B), 64 (C), 65 (D), 66 (E), 67 (F), and 69 (G). Left plot shows complete age range from 0 to 1,850 m.y., with a 50-m.y. bar width; right plot shows detailed area from 50 to 250 m.y., with a 5-m.y. bar width. All plots calculated with Isoplot software (Ludwig, 2003).

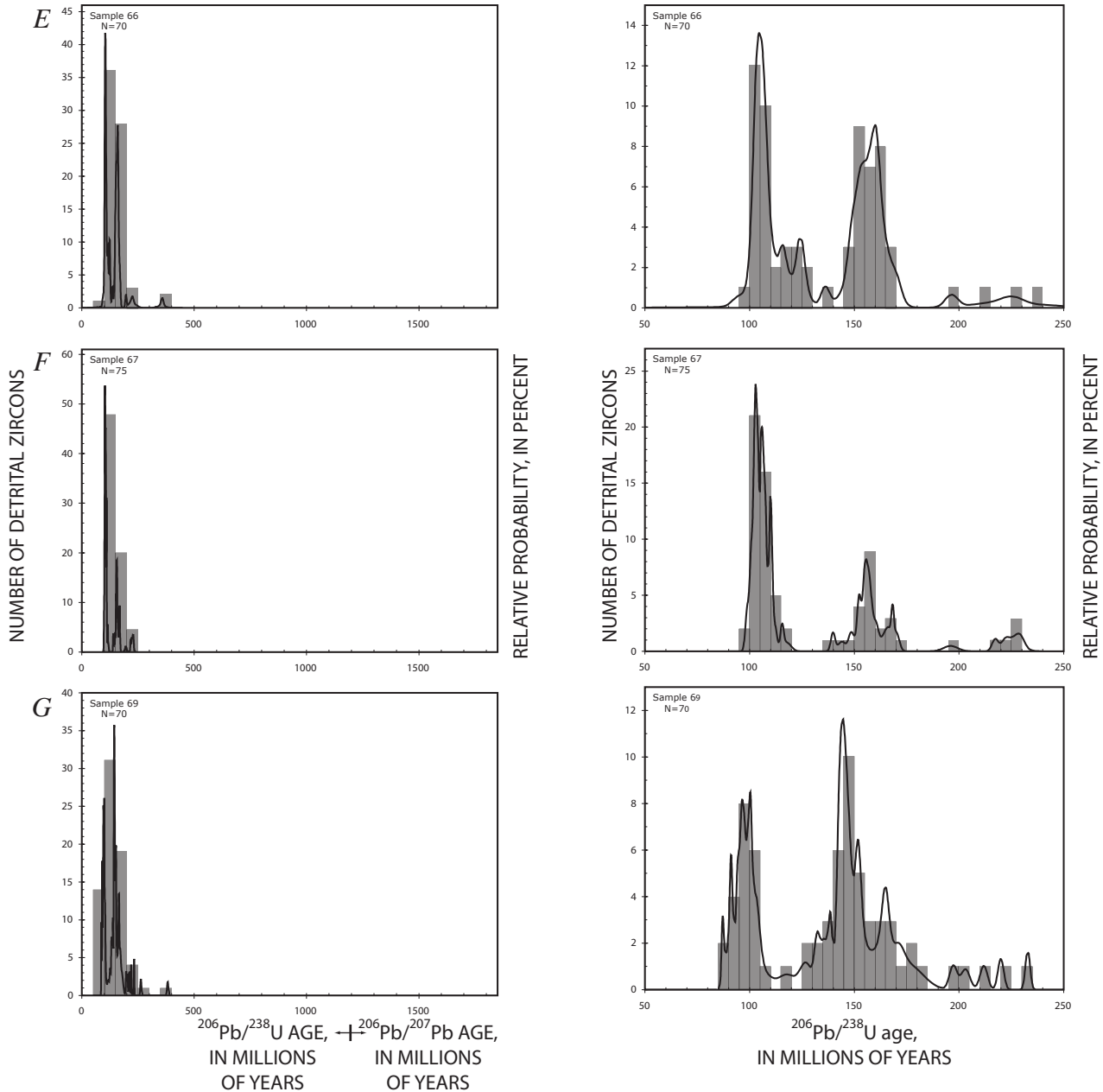


Figure 3.—Continued.

than the youngest peak ages of the eastern three samples (105, 103, and 97 m.y., from west to east; figs. 2, 3; tables 1, 2). Interpretation of these limiting ages depends on the time interval between crystallization and deposition of the zircons, which would largely be influenced by the presence or absence of igneous sources, the rate of exhumation of source terrains, and the geometry and efficiency of depositional systems.

We suggest that the youngest peak ages are close to the depositional age of the Sitka Graywacke. A subduction zone with associated arc magmatism was active along the western margin of North America during Late Cretaceous through early Tertiary time (for example, Miller, 1994; Moll-Stalcup and others, 1994; Plafker and others, 1994; Breitsprecher and Mortensen, 2004a, b), and so numerous sources of juvenile

zircons were present. Also occurring at that time was rapid uplift of the Coast Mountains of southeastern Alaska and British Columbia (for example, Hollister, 1982; Crawford and others, 2000; McClelland and Mattinson, 2000), which would have provided additional sources of young zircons. Moreover, because the Chugach terrane turbidites are considered to have been deposited by northward-flowing currents in the paleotrench (Nilsen and Moore, 1979; Nilsen and Zuffa, 1982; Plafker and others, 1994), little time intervened between deposition and accretion of the trench sediment. Taking these considerations together, we infer little time, possibly ≤ 5 m.y. between crystallization and deposition of the detrital zircons. Thus, these detrital-zircon limiting ages would indicate a range in the maximum age of these Sitka Graywacke samples from ~ 72 to ~ 105 m.y.

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The difference in the youngest peak ages between the eastern three and western four sample localities may be due to both a change in provenance and a fault. Histograms and relative-probability plots of detrital-zircon ages from the eastern and western samples differ (fig. 3), and so the two population groups are here considered separately. Because the two groups appear to be internally similar, we combined the data for the western four and eastern three samples (fig. 4). Because the patterns of the combined histograms and relative-probability plots differ for the two groups, the provenance must have changed. Is the difference between the two areas due only to a change in provenance? If so, then the entire sampling area would consist of a single stratigraphic section more than 15 km thick. Such a great thickness of a single panel in an accretionary complex would be extremely unusual. The overall westward-younging pattern of the youngest detrital-zircon ages is consistent with the standard model of accretionary prisms, which has trenchward-younging, fault-bounded panels of accreted sediment. Moreover, intraformational thrust faults are nearly ubiquitous in Chugach terrane turbidites (for example, Sample and Moore, 1987). Numerous faults occur within the

Sitka Graywacke, most of which bound thrust panels and others of which are later-stage right-lateral strike-slip faults (Davis and others, 1998; Karl and others, in press). No previous fault was mapped between the western and eastern localities, and no obvious sedimentologic differences exist. We infer a fault, probably a thrust fault, in an uncertain location between the eastern three and western four sample localities (fig. 2).

The detrital-zircon limiting ages for the Sitka Graywacke allow comparison and correlation to other parts of the Chugach accretionary complex in southern Alaska. The Valdez Group comprises the largest area of the Chugach terrane in south-central Alaska (fig. 1). The western four samples of Sitka Graywacke have youngest peak ages of late Campanian (fig. 6) on the time scale of Gradstein and others (2005). Assuming that the detrital zircons were deposited within several million years of crystallization, they indicate that the western part of the Sitka Graywacke is correlative with the Valdez Group (fig. 6; Grant and Higgins, 1910; Tysdal and Plafker, 1978). All reliable fossil ages from the Valdez Group are Maastrichtian (65.5–70.6 m.y.; Tysdal and Plafker, 1978; Nelson and others, 1985), although two samples have questionable age calls

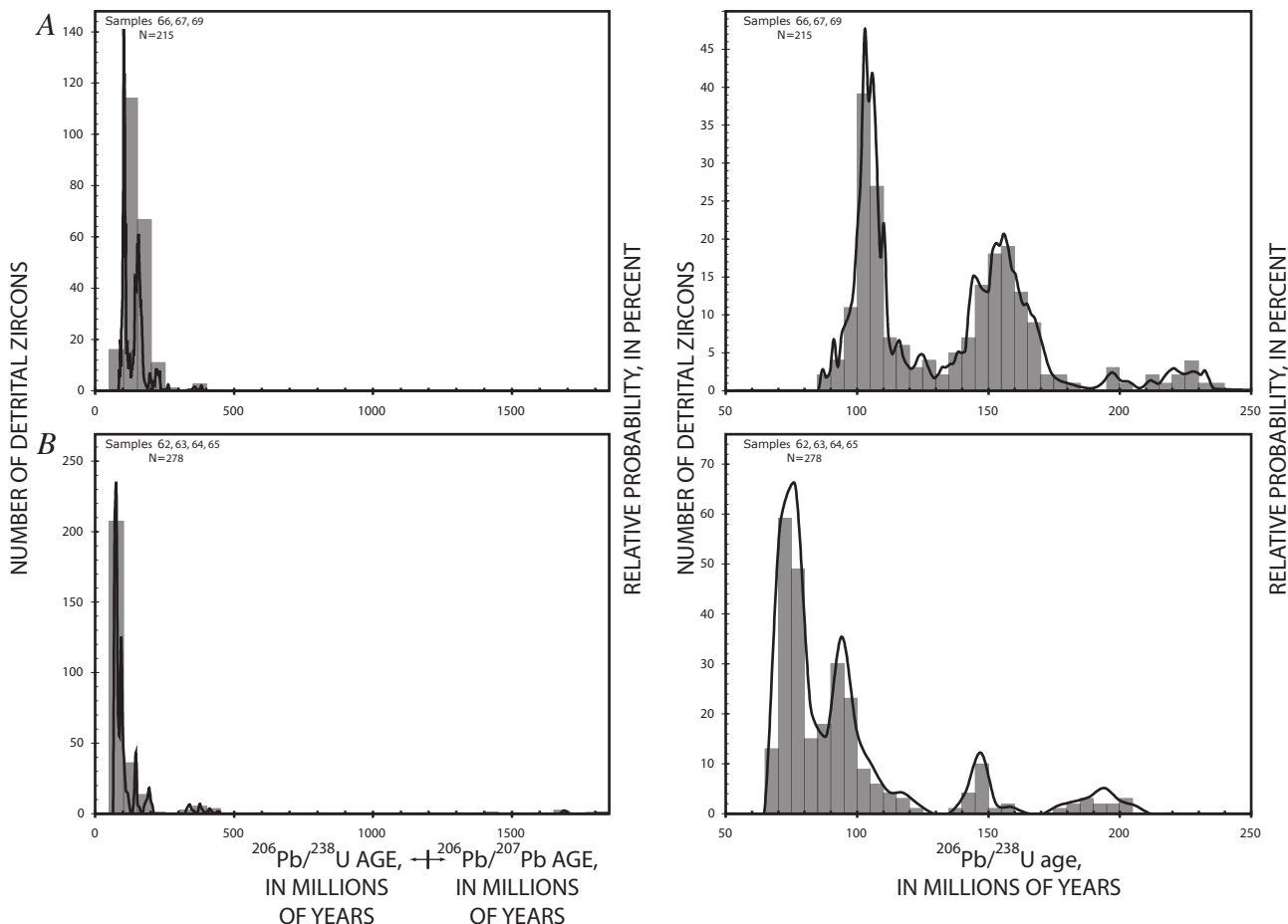


Figure 4. Detrital-zircon histograms and relative-probability plots for eastern three samples (66, 67, 69) (A) and western four samples (62, 63, 64, 65) (B). Same symbols as in figure 3.

that are Campanian-Maastrichtian (65.5–83.5 m.y.; Tysdal and Plafker, 1978; Winkler, 1992). If less than 5 m.y. intervened between crystallization and deposition of the western Sitka Graywacke samples, they would be Maastrichtian.

The younger part of the Sitka Graywacke also appears to be correlative with the upper part of the Yakutat Group of Tarr and Butler (1909). The Yakutat Group consists broadly of an eastern, or “lower,” melange facies and a western, or “upper,” flysch facies, both of which are considered correlative with the Chugach terrane (Plafker and others, 1977, 1994). The Yakutat Group is in the eastern part of the Yakutat terrane (fig. 1), which is presently accreting onto the southern Alaska margin (Plafker and others, 1978, 1994). The Yakutat terrane likely lay south of the Sitka Graywacke, off the coast of southern British Columbia, when it was offscraped (Plafker and others, 1994). Thus, on a large scale, the Yakutat Group was the now-dismembered part of the accretionary complex that lay to the south of the Sitka Graywacke. One Campanian fossil was obtained in outcrop from the upper part of the Yakutat Group (Jones and Clark, 1973), and Campanian and Maastrichtian foraminifers were recovered from an offshore borehole that penetrated the Yakutat Group (W.V. Sliter, in Rau and others, 1983). Flysch of the upper part of the Yakutat Group is more quartzofeldspathic and has a higher proportion of plutonic-rock fragments relative to the Sitka Graywacke and the Valdez Group (Nilsen, 1984). Nonetheless, the similarity of the youngest peak ages for the western four samples of Sitka Graywacke to the fossil ages for the Yakutat Group indicate that the rocks are approximately coeval, as previously considered (for example, Plafker and others, 1994), although they differ compositionally.

The Sitka Graywacke is at least ~30 m.y. older than the older part of the Chugach accretionary complex in southeastern Alaska (fig. 6). The Kelp Bay Group of Berg and Hinkley (1963), a melange containing fossils as young as Valanginian (Plafker and others, 1976; Decker, 1980), lies in fault contact with the Sitka Graywacke (see units KJkk and Trsv, fig. 2). The oldest detrital-zircon limiting age for the Sitka Graywacke is 105 m.y. (sample 66), and so a substantial age gap of ~30 m.y. exists on the time scale of Gradstein and others (2005). Of course, this age gap may be smaller if younger fossils are present in the Kelp Bay Group, or if older samples are present in the Sitka Graywacke.

The youngest peak ages for the eastern three samples of Sitka Graywacke (fig. 2) corresponds to the youngest fossil ages for melange of the McHugh Complex in south-central Alaska (fig. 6; see Plafker and others, 1994, for an overview of ages), which are Albian (99.6–112.0 m.y.) or Cenomanian (99.6–93.5 m.y.; Winkler and others, 1980). The McHugh Complex has been correlated with the Kelp Bay Group (Plafker and others, 1977, 1994; Decker, 1980). The oldest two samples of Sitka Graywacke are Albian (99.6–112.0 m.y.), and the easternmost sample (69) is Cenomanian (93.5–99.6 m.y.). Radiolarians from one small stratigraphically and structurally enigmatic area in Prince William Sound are dated at Coniacian through Turonian (89.3–83.5 m.y.; Haeussler and Nelson, 1993; Nelson and others, 1999), between the ages of the

Valdez Group and the McHugh Complex. Thus, assuming that the detrital-zircon limiting ages are close to the depositional age of the Sitka Graywacke, the age range of the unit is at least ~26–32 m.y.², significantly longer than the age range of reliably dated turbidites of the Valdez Group. Considering all these turbidite sequences together suggests relatively continuous turbidite sedimentation between Albian and Maastrichtian time; however, the volume of sediment deposited varies, and clearly a large pulse was deposited in Maastrichtian time (for example, Plafker and others, 1994).

The distribution of detrital-zircon ages for the Sitka Graywacke corresponds to that of magmatic ages from south-central Alaska to British Columbia. The eastern three samples contain the largest number of zircons dated at 100–110 m.y. (particularly samples 66 and 67) and at 145–165 m.y. (figs. 3, 4). Histograms derived from a compilation of igneous-rock ages from British Columbia (Breitsprecher and Mortenson, 2004a) and Yukon Territory (Breitsprecher and Mortenson, 2004b) are shown in figure 5. Numerous plutons in the Coast Mountains of British Columbia and southeastern Alaska dated at 90–110 m.y. may be the source of these zircons. In south-central Alaska, little magmatism has been reported. Numerous sources are possible, both near and far, for the older ages of 145–165 m.y. that may reflect (1) the Stikine terrane magmatic arc, which lies to the east and south of the study area (fig. 2; Van der Heyden, 1992); (2) plutons on the outer Queen Charlotte Islands in British Columbia (Butler and others, in press), (3) Middle to Late Jurassic plutons in southern British Columbia (for example, van der Heyden, 1992); (4) intrusions in the Saint Elias Mountains (Dodds and Campbell, 1988); or (5) younger phases of the Talkeetna and Chitina plutonic arcs in south-central Alaska (Plafker and others, 1989; Roeske and others, 2003; Trop and others, 2005). These detrital-zircon ages do not uniquely define the source terrain.

The western four samples have ages that are consistent with igneous sources from south-central Alaska to southern British Columbia. All four samples contain zircons dated at ~74 m.y. Plutons at the lower end of this age range occur in the western Alaska Range. Voluminous volcanism occurred in the Talkeetna Mountains of south-central Alaska between 80 and 70 m.y. Detrital zircons from the Cretaceous Matanuska Formation of south-central Alaska have a significant peak in ages of 75–80 m.y. (J.M. Trop, written commun., 2005). Numerous intrusions dated at 57–71 m.y. occur in the region between Juneau and Haines (Gehrels, 2000). A significant peak in detrital-zircon ages from 90–100 m.y. corresponds to a period of voluminous magmatism in southeastern Alaska and British Columbia (fig. 5), and so this peak age seems to reflect

²The younger age comes from 100 m.y. (the youngest peak age for the eastern three samples, at loc. 69) minus 74 m.y. (the oldest peak age for the western four samples, at locs. 62, 64, and 65) equals 26 m.y., whereas the older age comes from 104 m.y. (the oldest peak age for the eastern three samples, at loc. 66) minus 72 m.y. (the youngest peak age for the western four samples, at loc. 63) equals 32 m.y.

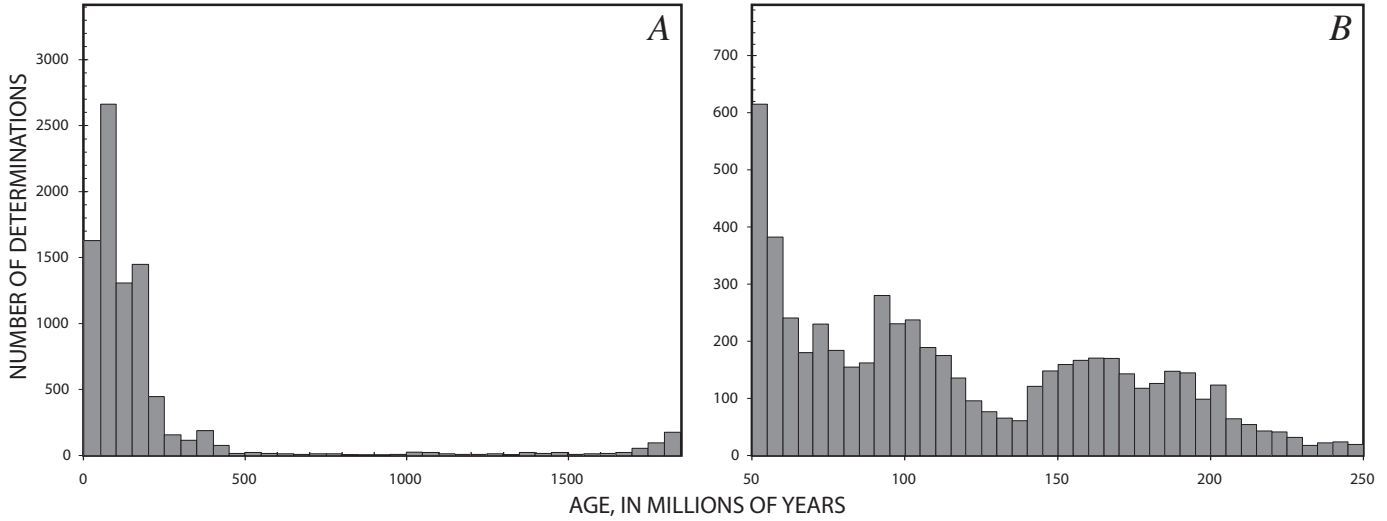


Figure 5. Histograms of 9,318 isotopic-age determinations from British Columbia and Yukon Territory, from compilations by Breitsprecher and Mortensen (2004a, b). *A*, All data. *B*, Data from 50 to 250 m.y., with a 5-m.y. bar width as in figures 3 and 4.

indeterminate sources from this region. Moreover, the relatively small number of zircons dated at 80–90 m.y. is also consistent with the absence of plutons in that interval in southern southeastern Alaska and British Columbia. Interestingly, the only >1-b.y. zircons were obtained from the western, young group of samples: 1,686±6, 1,693±8, 1,813±32, and 1,407±96 m.y. (²⁰⁶Pb/²⁰⁷Pb ages). These old ages may reflect sources in the northern Cordilleran miogeocline (Gehrels, 2000) or in the Yukon-Tanana terrane of southeastern Alaska and British Columbia (Kapp and Gehrels, 1998; Gehrels, 2001).

In conclusion, the detrital zircons in the samples of Sitka Graywacke do not uniquely identify a source region. Instead, possible source regions cover a vast area from south-central Alaska to southern British Columbia. However, all considerations of the regional tectonics of southeastern Alaska involve some amount of right-lateral strike slip eastward of the Sitka Graywacke. The Border Ranges Fault bounds the east side of the Chugach terrane, and speculations on the amount of right slip range from hundreds of kilometers to more than 1,000 km (Smart and others, 1996; Roeske and others, 2003). The slip must have occurred before 50 m.y., when a pluton intruded across the fault on Chichagof Island (Johnson and Karl, 1985, L.W. Snee, written commun., 1997). The Chatham Strait Fault has ~150 km of right slip (see discussion in Hudson and others, 1982). The Tintina Fault in British Columbia has ~420 km of right slip (Gabrielse, 1985) that likely occurred in early Tertiary time (Jackson and Mortensen, 2000). The eastern Denali Fault has ~370 km of right-lateral offset (Lowey, 1998). Thus, significant margin-parallel right slip clearly occurred in the interval between deposition of the Sitka Graywacke and ~50 m.y., after which major activity ceased on all these fault systems except the Chatham Strait Fault. Regardless of the uncertainties as to how much slip occurred on different faults at particular times, the Sitka Graywacke was situated south of its present position. Therefore, zircon sources in coastal and central to

southern interior British Columbia seem most likely. Although some ages are consistent with source regions in the St. Elias Mountains and in south-central Alaska, they are not demanded.

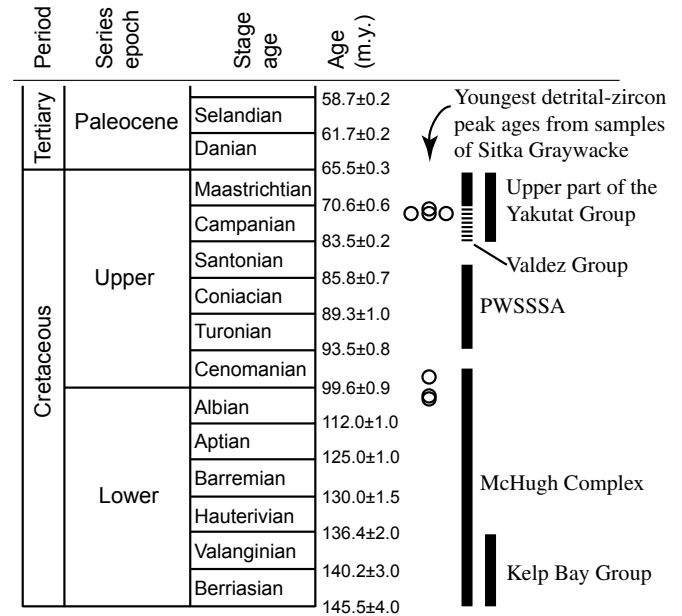


Figure 6. Age range for the Valdez Group (Tysdal and Plafker, 1978, Winkler, 1992; Plafker and others, 1994), younger part of the Kelp Bay Group (Plafker and others, 1976, Decker, 1980), and the McHugh Complex (Winkler, 1992; Plafker and others, 1994), fossiliferous rocks of central Prince William Sound (PWSSSA [Prince William Sound Special Study Area]; Nelson and others, 1999), “upper” flysch facies of the Yakutat Group (Jones and Clark, 1973, W.V. Sliter, in Rau and others, 1983), and detrital-zircon youngest peak ages for the Sitka Graywacke (this study). Time scale from Gradstein and others (2005).

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References Cited

- Berg, H.C., and Hinckley, D.W., 1963, Reconnaissance geology of northern Baranof Island, Alaska, *in* Contributions to general geology, 1961: U.S. Geological Survey Bulletin 1141-O, p. O1-O24 [includes geologic map, scale 1:125,000].
- Breitsprecher, Katrin, and Mortensen, J.K., 2004a, BCAGE 2004A-1—a database of isotopic age determinations for rock units from British Columbia: British Columbia Geological Survey Open File 2004-3 [release 3.0].
- Breitsprecher, Katrin, and Mortensen, J.K., 2004b, YukonAge 2004; a database of isotopic age determinations for rock units from Yukon Territory: Whitehorse, Yukon Territory, Canada, Yukon Geological Survey [CD-ROM].
- Butler, R.F., Gehrels, G.E., Hart, William, Davidson, Cameron, and Crawford, M.L., in press, Paleomagnetism of Late Jurassic to mid-Cretaceous plutons near Prince Rupert, British Columbia, *in* Haggart, J.W., Enkin, R.J., and Monger, J.W.H., eds., Paleogeography of western North America; constraints on large-scale latitudinal displacements: Geological Association of Canada Special Paper.
- Crawford, M.L., Crawford, W.A., and Gehrels, G.E., 2000, Terrane assembly and structural relationships in the eastern Prince Rupert quadrangle, British Columbia, *in* Stowell, H.H., and McClelland, W.C., eds., Tectonics of the Coast Mountains, southern Alaska and British Columbia: Geological Society of America Special Paper 343, p. 1-21.
- Davis, S.J., Roeske, S.M., and Karl, S.M., 1998, Late Cretaceous to early Tertiary transtension and strain partitioning in the Chugach accretionary complex, SE Alaska: *Journal of Structural Geology*, v. 20, no. 5, p. 639-654.
- Decker, J.E., 1980, Geology of a Cretaceous subduction complex, western Chichagof Island, southeastern Alaska: Stanford, Calif., Stanford University, Ph.D. thesis, 135 p.
- Decker, John, Nilson, T.H., and Karl, S.M., 1979, Turbidite facies of the Sitka Graywacke, southeastern Alaska, *in* Johnson, K.M., and Williams, J.R., eds., The United States Geological Survey in Alaska; accomplishments during 1978: U.S. Geological Survey Circular 804-B, p. B125-B129.
- Dodds, C.J., and Campbell, R.B., 1988, Potassium-argon ages of mainly intrusive rocks in the Saint Elias Mountains, Yukon and British Columbia: Geological Survey of Canada Paper 87-16, 43 p.
- Gabrielse, Hubert, 1985, Major dextral transcurrent displacements along the northern Rocky Mountain Trench and related lineaments in north-central British Columbia: *Geological Society of America Bulletin*, v. 96, no. 1, p. 1-14.
- Gehrels, G.E., 2000, Reconnaissance geology and U-Pb geochronology of the west flank of the Coast Mountains between Juneau and Skagway, southeastern Alaska, *in* Stowell, H.H., and McClelland, W.C., eds., Tectonics of the Coast Mountains, southern Alaska and British Columbia: Geological Society of America Special Paper 343, p. 213-233.
- Gehrels, G.E., 2001, Geology of the Chatham Sound region, Southeast Alaska and coastal British Columbia: *Canadian Journal of Earth Sciences*, v. 38, no. 11, p. 1579-1599.
- Gradstein, F.M., Ogg, J.G., and Smith, A.G., 2005, A geologic time scale 2004: London, Cambridge University Press, 610 p.
- Grant, U.S., and Higgins, D.F., 1910, Reconnaissance of the geology and mineral resources of Prince William Sound, Alaska: U.S. Geological Survey Bulletin 443, 89 p. [includes geologic map, scale 1:250,000].
- Haeussler, P.J., and Nelson, S.W., 1993, Structural evolution of the Chugach-Prince William Terrane at the hinge of the orocline in Prince William Sound, and implications for ore deposits, *in* Dusel-Bacon, Cynthia, and Till, A.B., eds., Geologic studies in Alaska by the U.S. Geological Survey in 1992: U.S. Geological Survey Bulletin 2068, p. 143-162.
- Hollister, L.S., 1982, Metamorphic evidence for rapid (2 mm/yr) uplift of a portion of the Central Gneiss Complex, Coast Mountains, B.C.: *Canadian Mineralogist*, v. 20, no. 3, p. 319-332.
- Hudson, Travis, Plafker, George, and Dixon, Kirk, 1982, Horizontal offset history of the Chatham Strait fault, *in* Coonrad, W.L., ed., The United States Geological Survey in Alaska; accomplishments during 1980: U.S. Geological Survey Circular 844, p. 128-132.
- Jackson, L.E., Jr., and Mortensen, J.K., 2000, New constraints indicate mainly early Paleogene displacement on the Tintina fault zone in the northern Cordillera [abs.]: *Geological Society of America Abstracts with Programs*, v. 32, no. 6, p. 21.

10 Constraints on the Age and Provenance of the Chugach Accretionary Complex, Alaska

- Johnson, B.R., and Karl, S.M., 1985, Geologic map of western Chichagof and Yakobi Islands, southeastern Alaska: U.S. Geological Survey Miscellaneous Investigations Map I-1506, 15 p., scale 1:125,000.
- Jones, D.L., and Clark, S.H.B., 1973, Upper Cretaceous (Maestrichtian) fossils from the Kenai-Chugach mountains, Kodiak and Shumagin Islands, southern Alaska: U.S. Geological Survey Journal of Research, v. 1, no. 2, p. 125-136.
- Kapp, P.A., and Gehrels, G.E., 1998, Detrital zircon constraints on the tectonic evolution of the Gravina Belt, Southeastern Alaska: Canadian Journal of Earth Sciences, v. 35, no. 3, p. 253-268.
- Karl, S.K., Haeussler, P.J., Himmelberg, Glenn, and Zumsteg, C.Z., in press, Geologic map of Baranof Island, Alaska: U.S. Geological Survey Scientific Investigations Map, scale 1:250,000.
- Loney, R.A., Berg, H.C., Pomeroy, J.S., and Brew, D.A., 1963, Reconnaissance geologic map of Chichagof Island and northwestern Baranof Island, Alaska: U.S. Geological Survey Miscellaneous Geologic Investigations Map I-388, scale 1:250,000.
- Lowey, G.W., 1998, A new estimate of the amount of displacement on the Denali Fault system based on the occurrence of carbonate megaboulders in the Dezadeash Formation (Jura-Cretaceous), Yukon, and the Nutzotin Mountains Sequence (Jura-Cretaceous), Alaska: Bulletin of Canadian Petroleum Geology, v. 46, no. 3, p. 379-386.
- Ludwig, K.R., 2003, User's manual for Isoplot 3.0: a geochronological toolkit for Microsoft Excel: Berkeley, Calif., Berkeley Geochronology Center Special Publication 4, 71 p.
- McClelland, W.C., and Mattinson, J.M., 2000, Cretaceous-Tertiary evolution of the western Coast Mountains, central southeastern Alaska, in Stowell, H.H., and McClelland, W.C., eds., Tectonics of the Coast Mountains, southern Alaska and British Columbia: Geological Society of America Special Paper 343, p. 159-182.
- Miller, T.P., 1994, Pre-Cenozoic plutonic rocks in mainland Alaska, in Plafker, George, and Berg, H.C., eds., The geology of Alaska, v. G-1 of The geology of North America: Boulder, Colo., Geological Society of America, p. 535-554.
- Moll-Stalcup, E.J., Brew, D.A., and Vallier, T.L., 1994, Latest Cretaceous and Cenozoic magmatic rocks of Alaska, in Plafker, George, and Berg, H.C., eds., The geology of Alaska, v. G-1 of The geology of North America: Boulder, Colo., Geological Society of America, scale 1:2,500,000.
- Nelson, S.W., Dumoulin, J.A., and Miller, M.L., 1985, Geologic map of the Chugach National Forest, Alaska: U.S. Geological Survey Miscellaneous Field Studies Map MF-1645-B, 16 p., scale 1:250,000.
- Nelson, S.W., Miller, M.L., Haeussler, P.J., Snee, L.W., Phillips, P.J., and Huber, Carol, 1999, Preliminary geologic map of the Chugach National Forest special study area, Alaska: U.S. Geological Survey Open-File Report 99-362, scale 1:63,360 [<http://wrgis.wr.usgs.gov/open-file/of99-362/>].
- Nilsen, T.H., 1984, Trench-fill submarine-fan associations of the Upper Cretaceous Chugach terrane, southern Alaska: Geo-Marine Letters, v. 3, no. 2-4, p. 179-185.
- Nilsen, T.H., and Moore, G.W., 1979, Reconnaissance study of Upper Cretaceous to Miocene stratigraphic units and sedimentary facies, Kodiak and adjacent islands, Alaska: U.S. Geological Survey Professional Paper 1093, 34 p.
- Nilson, T.H., and Zuffa, G.G., 1982, The Chugach terrane, a Cretaceous trench-fill deposit southern Alaska, in Leggett, J.K., ed., Trench-forearc geology—sedimentation and tectonics on modern and ancient active plate margins: Geological Society of London, Special Publication 10, p. 213-227.
- Nokleberg, W.J., West, T.D., Dawson, K.M., Shpikerman, V.I., Bundtzen, T.K., Parfenov, L.M., Monger, J.W.H., Ratkin, V.V., Baranov, B.V., Byalobzhesky, S.G., Diggles, M.F., Eremin, R.A., Fugita, Kazuga, Gordey, S.P., Gorodinskiy, M.E., Goryachev, N.A., Feeney, T.D., Frolov, T.F., Grantz, Arthur, Khanchuk, A.I., Koch, R.D., Natalin, B.A., Natapov, L.M., Norton, I.O., Patton, W.W., Jr., Plafker, George, Pozdeev, A.I., Rozenblum, I.S., Scholl, D.W., Sokolov, S.D., Sosunov, G.M., Stone, D.B., Tabor, R.W., Tsukanov, N.V., and Vallier, T.L., 1998, Summary terrane, mineral deposit, and metallogenic belt maps of the Russian Far East, Alaska, and the Canadian Cordillera: U.S. Geological Survey Open-File Report 98-136, CD-ROM [<http://wrgis.wr.usgs.gov/open-file/of98-136/>].
- Plafker, George, Hudson, Travis, Bruns, T.R., and Rubin, Meyer, 1978, Late Quaternary offsets along the Fairweather fault and crustal plate interactions in southern Alaska: Canadian Journal of Earth Sciences, v. 15, no. 5, p. 805-816.
- Plafker, George, Jones, D.L., Hudson, Travis, and Berg, H.C., 1976, The Border Ranges Fault system in the Saint Elias Mountains and Alexander Archipelago, in Cobb, E.H., ed., The United States Geological Survey in Alaska; accomplishments during 1975: U.S. Geological Survey Circular 733, p. 14-16.

- Plafker, George, Nokleberg, W.J., and Lull, J.S., 1989, Bedrock geology and tectonic evolution of the Wrangellia, Peninsular, and Chugach terranes along the Trans-Alaska Crustal Transect in the Chugach Mountains and southern Copper River basin, Alaska: *Journal of Geophysical Research*, v. 94, no. B4, p. 4255–4295.
- Plafker, George, Jones, D.L., and Pessagno, E.A., Jr., 1977, A Cretaceous accretionary flysch and mélangé terrane along the Gulf of Alaska margin, in Blean, K.M., ed., *The United States Geological Survey in Alaska—accomplishments during 1976*: U.S. Geological Survey Circular 751–B, p. B41–B43.
- Plafker, George, Moore, J.C., and Winkler, G.R., 1994, Geology of the southern Alaska margin, in Plafker, George, and Berg, H.C., eds., *The geology of Alaska*, v. G–1 of *The geology of North America*: Boulder, Colo., Geological Society of America, p. 389–449.
- Rau, W.W., Plafker, George, and Winkler, G.R., 1983, Foraminiferal biostratigraphy and correlations in the Gulf of Alaska Tertiary province: U.S. Geological Survey Oil and Gas Investigations Chart 120, 11 p., 3 sheets.
- Reifenstuhel, R.R., 1986, Geology of the Goddard Hot Springs area, Baranof Island, southeastern Alaska: Alaska Division of Geological and Geophysical Surveys Public-Data File 86–2, 82 p.
- Roeske, S.M., Snee, L.W., and Pavlis, T.L., 2003, Dextral-slip reactivation of an arc-forearc boundary during Late Cretaceous–Early Eocene oblique convergence in the northern Cordillera, in Sisson, V.B., Roeske, S.M., and Pavlis, T.L., eds., *Geology of a transpressional orogen developed during ridge-trench interaction along the North Pacific margin*: Geological Society of America Special Paper 371, p. 141–169.
- Sample, J.C., and Moore, J.C., 1987, Structural style and kinematics of an underplated slate belt, Kodiak and adjacent islands, Alaska: *Geological Society of America Bulletin*, v. 19, no. 1, p. 7–20.
- Smart, K.J., Pavlis, T.L., Sisson, V.B., Roeske, S.M., and Snee, L.W., 1996, The Border Ranges fault system in Glacier Bay National Park, Alaska; evidence for major early Cenozoic dextral strike-slip motion: *Canadian Journal of Earth Sciences*, v. 33, no. 9, p. 1268–1282.
- Stacey, J.S., and Kramers, J.D., 1975, Approximation of terrestrial lead isotope evolution by a two-stage model: *Earth and Planetary Science Letters*, v. 26, no. 2, p. 207–221.
- Tarr, R.S., and Butler, B.S., 1909, *The Yakutat Bay region, Alaska*: U.S. Geological Survey Professional Paper 64, 183 p.
- Trop, J.M., Szuch, D.A., Rioux, Matthew, and Blodgett, R.B., 2005, Sedimentology and provenance of the Upper Jurassic Naknek Formation, Talkeetna Mountains, Alaska; bearings on the accretionary tectonic history of the Wrangellia composite terrane: *Geological Society of America Bulletin*, v. 117, no. 5–6, p. 570–588.
- Tysdal, R.G., and Plafker, George, 1978, Age and continuity of the Valdez Group, southern Alaska, in Sohl, N.F., and Wright, W.B., eds., *Changes in stratigraphic nomenclature by the U.S. Geological Survey, 1977*: U.S. Geological Survey Bulletin 1457–A, p. A120–A124.
- van der Heyden, Peter, 1992, A Middle Jurassic to early Tertiary Andean-Sierran arc model for the Coast Belt of British Columbia: *Tectonics*, v. 11, no. 1, p. 82–97.
- Winkler, G.R., 1992, Geologic map and summary geochronology of the Anchorage 1°×3° quadrangle, southern Alaska: U.S. Geological Survey Miscellaneous Investigations Map I–2283, scale 1:250,000.
- Winkler, G.R., Silberman, M.L., Grantz, Arthur, Miller, R.J., and MacKevett, E.M., Jr., 1980, Geologic map and summary geochronology of the Valdez quadrangle, southern Alaska: U. S. Geological Survey Open-File Report 80–892–A, 2 sheets, scale 1:250,000.

Table 2

Table 2. Analytical data on detrital zircons.

[All analyses by laser-ablation inductively coupled plasma mass spectroscopy (ICPMS)]

Analysis	U content (ppm)	²⁰⁶ Pb/ ²⁰⁴ Pb ratio	U/Th ratio	²⁰⁶ Pb/ ²⁰⁸ Pb ratio	± (pct)	²⁰⁵ Pb/ ²³⁵ U ratio	± (pct)	²⁰⁶ Pb/ ²⁰⁷ Pb ratio	± (pct)	Error correla- tion	²⁰⁶ Pb/ ²³⁸ U age (m.y.)	± (m.y.)	²⁰⁷ Pb/ ²³⁵ U age (m.y.)	± (m.y.)	²⁰⁶ Pb/ ²⁰⁷ Pb age (m.y.)	± (m.y.)
43	68	2,025	28	0.02839	3.2	0.16739	8.3	23.39	7.7	0.38	180.5	5.8	157.1	14.0	-182	96
44	137	3,031	22	0.02255	0.9	0.14658	3.6	21.22	3.4	0.25	143.8	1.3	138.9	5.3	56	41
45	79	2,484	54	0.03677	0.5	0.24964	6.3	20.31	6.2	0.08	232.8	1.1	226.3	15.7	159	73
46	112	2,152	24	0.02173	0.6	0.13592	7.4	22.04	7.4	0.08	138.6	0.9	129.4	10.2	-36	90
47	83	2,526	44	0.02592	1.0	0.17981	3.5	19.88	3.4	0.28	165.0	1.7	167.9	6.4	209	39
48	305	6,703	14	0.02118	1.3	0.14224	2.7	20.53	2.3	0.49	135.1	1.8	135.0	3.9	134	28
49	229	5,653	40	0.03472	0.6	0.25251	1.8	18.96	1.7	0.35	220.0	1.4	228.6	4.6	318	19
50	133	3,874	73	0.02510	2.6	0.17077	5.2	20.27	4.5	0.50	159.8	4.2	160.1	9.0	164	53
51	267	3,995	52	0.01966	3.7	0.13103	5.0	20.69	3.4	0.74	125.5	4.7	125.0	6.6	115	40
52	68	1,714	79	0.02576	1.5	0.19141	8.3	18.56	8.2	0.18	164.0	2.5	177.8	16.0	367	92
53	471	5,846	47	0.01612	1.5	0.10415	1.9	21.34	1.3	0.76	103.1	1.5	100.6	2.0	43	15
54	137	2,823	71	0.01984	2.0	0.13186	5.2	20.75	4.8	0.38	126.7	2.5	125.8	6.9	109	57
55	179	4,260	38	0.02394	1.4	0.15326	4.5	21.53	4.3	0.30	152.5	2.1	144.8	7.0	21	52
56	295	6,890	22	0.02318	0.8	0.15826	2.0	20.20	1.9	0.39	147.7	1.2	149.2	3.2	172	22
57	69	2,487	25	0.02776	2.6	0.19558	8.2	19.57	7.8	0.32	176.5	4.6	181.4	16.2	245	90
58	69	1,265	30	0.02553	2.2	0.13430	4.6	26.21	4.1	0.46	162.5	3.5	128.0	6.3	-474	54
59	105	3,750	53	0.02343	1.9	0.16386	4.8	19.71	4.4	0.39	149.3	2.8	154.1	7.9	228	51
60	194	5,028	14	0.03107	0.9	0.21431	2.0	19.99	1.8	0.46	197.3	1.8	197.2	4.3	196	21
61	77	1,922	31	0.02295	1.6	0.14263	5.8	22.19	5.6	0.28	146.3	2.4	135.4	8.4	-52	68
62	745	5,856	43	0.02696	1.5	0.18534	2.2	20.06	1.6	0.68	171.5	2.6	172.6	4.1	188	19
63	106	1,852	101	0.02113	2.6	0.13920	14.1	20.93	13.8	0.19	134.8	3.6	132.3	19.7	88	164
64	90	955	70	0.02381	1.3	0.14263	5.6	23.01	5.5	0.22	151.7	1.9	135.4	8.1	-142	68
65	198	2,938	55	0.01595	1.5	0.10277	4.2	21.40	3.9	0.36	102.0	1.6	99.3	4.4	35	47
66	77	1,443	61	0.02278	1.0	0.11866	10.6	26.46	10.5	0.09	145.2	1.5	113.9	12.7	-500	140
67	67	1,261	56	0.02467	4.5	0.15255	9.5	22.30	8.4	0.47	157.1	7.2	144.2	14.6	-64	102
68	86	1,120	28	0.01566	1.4	0.08348	9.0	25.86	8.9	0.15	100.1	1.4	81.4	7.6	-439	117
69	163	2,549	21	0.02237	0.9	0.16881	6.4	18.28	6.3	0.14	142.6	1.3	158.4	10.9	401	71
70	135	2,362	28	0.01571	0.7	0.11183	7.7	19.37	7.7	0.09	100.5	0.7	107.6	8.7	269	88