The Paleoenvironment of Humans and Marine Birds of the Aleutian

Islands: Three Millennia of Change.

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ABSTRACT

A unique window into the biological history of the Aleutian Islands is provided by the zooarcheology of early human sites. We focus on the paleoavifauna hunted by early Aleuts who inhabited Amchitka and Buldir islands (central Aleutians), and Shemya Island (western Aleutians) from about 3,500 years ago to the present. Most of the seabird species recovered from these early sites varied widely in distribution and abundance through time and space. Pelagic procellariids such as Short-tailed Albatrosses and Short-tailed Shearwaters were present and abundant at most sites and at most times. During periods of increased temperatures and precipitation (e.g., 650-1100 ybp), nearshore foragers such as cormorants and parakeet auklets increased in abundance, but during periods of cooling (e.g., 1800-2100 ybp), piscivorous birds feeding offshore such as murres and kittiwakes predominated. Over three millennia, we found that marine bird populations were negatively correlated with temperature and positively correlated with precipitation. We detected hunter-related depletions of populations breeding in accessible colonies at small scales of space and time, but we did not observe widespread or long-term effects. We conclude that local oceanography and regional changes in prey bases caused by environmental and climate change in the past had a significant impact on the distribution and abundance of Aleutian marine birds.

Key Words: Paleobiology, zooarchaeology, environmental change, population dynamics, marine birds, Aleutian Islands.

INTRODUCTION

"The natives of Attu inform me that the Cod has not long been an inhabitant of the waters around that island. Its advent was near 1873. Previous to that time individuals had been obtained but rarely and many of the men had not seen a Cod [before]...The Codfish are very numerous at the present day. They attain immense size there. I saw one individual in February, 1881, that weighed just out of the water an even thirty pounds."

(Turner, 1886)

The first Europeans to the Aleutian Islands (Fig. 1) viewed an environment that differed in many ways from the present. Davydov (1810) and Berkh (1823) noted abundances of marine vertebrate animals in the Aleutians far higher than exist today—even after more than a half-century of exploitation by Russian fur traders. The first scientifically detailed reports were by William Dall, a naturalist associated with the California Academy of Sciences. He observed that Black Brants (*Branta nigricans*) were breeding in the Near Islands, that Ancient Murrelets (*Synthliboramphus antiquus*) occurred in high abundances throughout the Aleutians, and that Parakeet Auklets (*Aethia psittacula*) were restricted to only Amchitka Island in the central Aleutians (Dall, 1874). None of these observations are the case today. The most quantitative reports, however, were by Lucien Turner, who worked as a weather observer and naturalist on Attu Island in 1878 - 1881. He reported that Pacific Cod (*Gadus macrocephalus*) and Atka Mackerel (*Pleurogrammus monopterygius*) were unknown in the Near Islands five years previous to his arrival, but during his time on Attu Island they occurred in prodigious numbers. For example, from June 1 to July 31, 1879, five families

on Attu Island collected 10,000 pounds of Atka Mackerel by hand. Capelin (*Mallotus villosus*) were perhaps even more abundant in the waters around Atka Island: "...dead fish were so thick on the beach that it was impossible to walk without stepping on hundreds of them. They could be gathered with a shovel, they laid so thickly" (Turner 1886: 102). [Capelin are now nearly absent from the region for reasons unknown (Piatt and Anderson, 1995)]. He described birds on Attu Island at that time as common species that are rarely or never seen there today (e.g., Red-legged Kittiwake *Rissa brevirostris*) and, based on eyewitness accounts by native Attuans, reported that Pallas's Cormorants (*Phalacrocorax perspicillatus*)—by then extinct—abounded there only 15 years earlier.

Many of the subsequent changes in the natural environment of the Aleutians and Bering Sea clearly were the direct or indirect effect of European activities. These include the near extirpation of sea otters, *Enhydra lutris* (Doroff *et al.*, 2003); declining fisheries in cod, pollock (*Theragra chalcogramma*), and decapoda (Jackson *et al.*, 2001); and disappearance of Short-tailed Albatross (*Phoebastria* [*Diomedea*] *albatrus*) and Aleutian Canada Geese (*Branta canadensis leucopareia*) from most breeding islands (Jones, 1963). Other changes, including the abrupt appearance in the 1880s of Arctic Cod and Atka Mackerel in the Near Islands, and the recent population crashes of sea otters, Steller sea lions (*Eumetopias jubatus*) and Northern fur seals (*Callorhinus ursinus*) (Braham *et al.*, 1980; Trites, 1992; Trites and Donnelly, 2003), are less attributable to anthropogenic effects and may be associated with environmental variability (Francis, 1996) or other yet unknown factors.

Recent evidence indicates that the Beringian and North Pacific regions are undergoing rapid environmental change (Trenbeth and Hurrell, 1994; Mantua *et al.*, 1997), with consequent cascading effects on the biota (Oedekoven *et al.*, 2001; Jones *et al.*, 2002; Steneck *et al.*, 2002). At the same time, the collapse of virtually every major fishery around the world in recent decades (Amorosi *et al.*, 1996; Lyman, 1996; Jackson *et al.*, 2001; Reitz, 2004) has made it urgent to understand the underlying processes driving change in marine ecosystems. As our knowledge of ecosystem processes increases, it is clear that present-day conditions and records are inadequate to characterize long-term, large-scale ecosystem variability and function. It is valuable, therefore, to place the present conditions in Beringian marine ecosystems into a deep temporal context of historical patterns and to provide a baseline for interpreting local and regional change.

A unique window into the biological history of the Beringian environment is provided by the zooarcheology of early human sites. When Europeans entered Beringia in the mid-18th century, the region was already populated by indigenous peoples for many millennia. At the time of first contact, the early Aleuts had been living in the Aleutian Islands for about 5,000 years (Dumond, 2001). They were superb hunters of the marine environment and had developed a sophisticated material culture capable of exploiting marine vertebrates and invertebrates in pelagic and nearshore environments. The remains of food—bones, invertebrate remains, and other domestic waste—and cultural artifacts formed "kitchen middens" near the dwellings. Lengthy residence led to the compaction and discontinuous development of plant cover and soil, and the gradual accumulation of sand, sandy loam, or loam deposits. Taphonomic processes, such as periodic soaking of the developing layer by atmospheric precipitation, led to its gradual transformation into a "cultural layer"—a stratum of associated horizons of soil enriched with bones, artifacts, and humus. Radiocarbon dating shows that cultural layers can be formed over hundreds and thousands of years (Dinesman *et al.*, 1999). Careful analysis of these materials, coupled with precision dating of depositional strata, can provide a detailed proxy of the dynamics of the vertebrate community thousands of years ago and up to the present (Klein and Cruz-Uribe, 1984; McCartney, 1984; Lefèvre *et al.*, 1997).

We will use three detailed studies of the early Aleutian environment as means to explore the historical patterns of change in the marine vertebrate community, and to serve as a model how zooarcheology and anthropology may help elucidate complex issues related to environmental correlates of climate change and perturbation. We will explore what relative roles human occupation and environmental change have had on the coastal biota of the central and western Aleutians. We will also indicate what insights paleo-environmental history can provide in interpreting the magnitude and direction of biotic change currently under study.

METHODS

Source Materials

Paleobiotic samples with radiocarbon dates from Beringian regions are increasingly available; most of the material used here was excavated in connection with archeological studies on early human occupation of islands and shores of the Gulf of Alaska. Alaska Peninsula, and Aleutian Islands (Desautels et al., 1971; Corbett et al., 1997; Corbett et al., In Press). The number of excavated archeological sites in the study region is highest in the Gulf of Alaska (e.g., Kachemak Bay, Kodiak I), and decreases with distance west of the Alaska Peninsula (e.g., Rat Islands, Near Islands). Although there has been a long history of archeological excavations in the Aleutian Islands (Dall, 1877; Jochelson, 1925; Everdam, 1934; Friedmann, 1934, 1935b, a; Everdam, 1936; Friedmann, 1937; Hrdlicka, 1945; Spaulding, 1962; Lippold, 1966, 1972), there are only a few that used quantified sampling methodologies, that have zooarcheological material identified to species, and that have the material directly associated with radiocarbon dates (Table 1). Only three of them (Amchitka Island; Buldir Island; Shemya Island) have sufficient radiocarbon-dated time depth in identified material for comparisons. We have focused on birds in this study because of the high diversity of bony material and large sample sizes of identified material.

Zooarcheological specimens are dated by their contextual matrix rather than by radiocarbon dating of individual objects, and we use published dates as ¹⁴C years before present (¹⁴C y BP) \pm error (Table 2). All dates are uncalibrated, since there have been few calibration studies from the Aleutian Island, and there seems to be great variability in the marine reservoir effect throughout the Bering Sea basin (Dumond and Griffin, 2002; Savinetsky and Khassanov, 2004). Identifications were made by comparison with reference

material of recent species of birds known from the Beringian region, including the Gulf of Alaska, Bering Sea, Alaska, Siberia and the Russian Far East, and regional seas of the Arctic Ocean.

Climatic reconstruction

To estimate the change in climatic conditions of the Aleutian Islands, we used evidence from published studies on Bering Island and Kodiak Island, located on the extreme western and eastern boundaries of the archipelago (Finney et al. 2000, Razjigaeva et al. 2004). These studies and several others are based on standard analysis of diatom presence and abundance in buried soils (Mann et al., 1998; Sawada et al., 1999; Finney et al., 2000; Calkin et al., 2001; Tunicliffe et al., 2001; Gorbarenko et al., 2002; Gorbarenko et al., 2004; Razjigaeva et al., 2004). The reconstruction of past temperatures and precipitation on Shemva Island, however, derived from analysis of layer-by-layer changes in ash content of peat and the degree of its decomposition sampled from deposits (Savinetsky et al., 2004) adjacent to the midden sites. In cases of good drainage, these changes in peat are determined by the temporal dynamics of warming and the saturation of the deposits (Dinesman et al., 1999). To determine the degree of change in precipitation over the time of formation of the whole layer, we used the allogenic ash content of the layers. The ash content of pure peat cannot exceed 15% (P'yavchenko, 1963), and all proportions that exceed this figure are caused by the introduction of mineral particles into the accumulating peat through the surface drainage of atmospheric precipitation from adjoining slopes, or during periods of flooding. To estimate the change in the summer temperature regime, we used indicators of layer-by-layer changes in the degree of peat decomposition. The degree of decomposition is determined by the percentage content of unstructured matrix, containing

small particles of non-humified remains along with humic matter. Aerobic microorganisms, which actively function only in the upper, peat-forming layer, play the primary role in the decomposition of organic remains. After being covered with a developing layer of peat, the degree of decomposition attained remains practically unchanged (Tyuremnov, 1976). Microorganism activity in the peat layer is only possible with suitable warmth and sufficient moisture, and is suppressed by low temperatures, drying, or waterlogging. If the humidity of the peat does not exceed its moisture capacity, the deciding factor influencing decomposition is temperature (Prozorova, 1988). Protracted waterlogging of a well-drained peat bog is unlikely, so that changes in the degree of peat decomposition reflect changes in the temperature regime during the warm half of the year (Savinetsky *et al.*, 2004).

Analytical methodology

The material analyzed here was originally excavated in 1 x 1 m quadrats by natural layers; that is, each layer of homogeneous depositional matrix was considered a single unit regardless of thickness. (However, the cultural deposit Pit 1 on Buldir Island was approximately 0.6 m deep but appeared to be a single unit; due to the high density of material, excavations were made from 10 cm horizons). All excavated material was screened by 6 mm (1/4 inch) and 3 mm (1/8 inch) mesh, and all retained material was reserved for identification using standard quantitative methodology (Klein and Cruz-Uribe, 1984; Lyman, 1994). Only bones identified to species were used in this analysis; however, for analytical purposes, some groups of closely related species were lumped together (e.g., Thick-billed Murres [*Uria lomvia*] and Common Murres [*U. aalge*], are reported as *Uria* spp. in Tables 4 - 6). Absolute numbers of bones for each species were grouped by anatomical element and position. The total number of identified specimens (NISP) for each

species from each discrete level and the Minimum Number of Individuals (MNI) index which estimates the least number of birds responsible for the observed material—were calculated using the total assemblage of elements (Grayson, 1984; Klein and Cruz-Uribe, 1984). The NISP and MNI are directly related (Grayson, 1984), but the former relies on fewer assumptions and is less sensitive to errors associated with artificial aggregation of elements and depositional histories. We report both indices, except for the material excavated from Amchitka Island, where the original data were given primarily in MNI values (Harrington, 1987).

The use of archeological faunal assemblages as proxy indicators of historical biological patterns is becoming increasingly common (Eyerdam, 1936; Lefèvre *et al.*, 1997; Bouchet *et al.*, 1999; Ervynck, 1999; Savinetsky *et al.*, 2004). There are, however, several intrinsic biases associated with these data, such as site-specific variability in preservation; misidentification of material; preference by the early hunters for particular species or cultural prohibitions for others (Yesner, 1981); and local effects by hunters on abundances of available prey (Broughton, 1994; Grayson, 2001). Preservation biases are minimized with the material used here. Due to the high volume of calcareous material in archeological midden sites in the Aleutians, it has long been recognized that bones and artifacts preserve extremely well (Dall, 1877; Jochelson, 1925; Corbett *et al.*, 1997; McCartney and Veltre, 1999). We minimized misidentification errors by using extensive reference collections and independent determinations by several of the authors. Only material identified to species were employed in these analyses.

Birds were classed by nesting, diet, and foraging strategies (Table 3) based on data in Byrd et al. (in press). To test whether there were hunter-associated effects in the avifaunal remains, we used as a control the temporal frequency of non-breeding pelagic species (i.e., albatrosses and shearwaters) in midden levels. These birds rarely are found on land in the Aleutian Islands and are minimally affected by human depredation or contact as are those species breeding in accessible surface- or ledge-based colonies (i.e., cormorants, murres, gulls).

Statistical Analyses

Several of the statistical tests we use are uncommon, so details are provided here. Residual Analysis [RA] (Haberman, 1972; Everitt, 1977) makes it possible to isolate sources of deviations from row-column independence in a comparison table. The standardized residual of each cell can be used to assess the particular nature of association between row and column profiles (i.e., individual cells in the comparison table). For cell (i,j), the standardized residual is defined as:

$$e_{ij} = \frac{(n_{ij} - E_{ij})}{\sqrt{E_{ij}}} \tag{1}$$

In simple terms, the standardized residual is the square root of a chi-square term with the sign of the deviation intact. An estimate of the variance of e_{ij} can be calculated as follows:

$$v_{ij} = \left(1 - \frac{n_{i+1}}{n_{i+1}}\right) \left(1 - \frac{n_{i+j}}{n_{i+1}}\right)$$
(2)

where "+" indicates row or column totals. The residual can be transformed to a standard normal deviate with mean of zero and unit variance with an estimate of term variance calculated from marginal totals. The transformation yields an adjusted standardized residual

$$d_{ij} = \frac{e_{ij}}{\sqrt{v_{ij}}} \tag{3}$$

that indicates the contribution of a cell to the total deviation of the table from independence. The significance of the adjusted standardized residual is equivalent to the one-tailed probability of Z in a Normal distribution (Zar, 1999).

Goodness-of-Fit Analysis [GFA] is useful to determine whether two temporal profiles of midden remains have the same underlying pattern or whether they differ significantly in the proportion of bony elements recovered through time. The Kolmogorov-Smirnov Two-Sample test compares paired cumulative frequencies of the element proportions for each time strata and calculates the test statistic D by the following (Hollander and Wolfe, 1973):

$$D = \max \left| F_1(i) - F_2(i) \right| \tag{4}$$

where $F_1(i)$ and $F_2(i)$ are the cumulative frequencies of category *i* for profiles F_1 and F_2 . For large samples, which are met with these data, the statistic *D* has a sampling distribution that is approximated by the chi-square distribution with df = 2 (Goodman, 1954):

$$\chi^{2} = 4D^{2} \left(\frac{n_{1}n_{2}}{n_{1} + n_{2}} \right)$$
(5)

where n_1 and n_2 are the sample sizes for each profile.

Significance levels were assessed using the Bonferroni inequality (Holm, 1979), which partitions the critical value α by the number of multiple tests performed. Specifically, the critical level $\hat{\alpha}_k(y)$ for rejection of a null hypothesis for the outcome *y* of the test statistic Y_k for test *k* of *n* total tests is:

$$\hat{\alpha}_k(Y_k) = \frac{\alpha}{n} \tag{6}$$

TEMPORAL PATTERNS OF BIOTIC CHANGE

Amchitka Island

Amchitka Island (51°23′ N 179°16′ E) is located approximately midway between the Alaska Peninsula and Kamchatka (Fig. 1), and is about 65 km long and 2 to 7 km wide with large expanses of cliffs and shoreline habitats suitable for seabird colonies (Krog, 1953; Emison *et al.*, 1971; White *et al.*, 1977). Numerous archeological sites have been identified, but only a few have been investigated. The most detailed study to date was made by Desautels et al. (1971) at four sites on the southwestern (Pacific) coast of the island. Sites were excavated in 1 x 1 m quadrats using standard procedures, and material was recovered from natural layers. Site 31 was the largest midden found on Amchitka Island (53 m x 29 m) and comprised 11 discrete strata; levels D, F, L, and M were not used in this analysis due to their small number of bony elements (< 5). Radiocarbon dates are available for three levels (Table 2): B (890 \pm 90 ¹⁴C y BP), H (1980 \pm 95 ¹⁴C y BP), and J (2550 \pm 95 ¹⁴C y BP). Details on analysis methodology are given in Siegel-Causey et al. (1991).

The bird bones were studied subsequently by Harrington (1987) and Siegel-Causey et al. (1991); results are summarized in Table 4. Harrington (1987) only reported MNIs for individual species; Siegel-Causey et al. (1991) reidentified all cormorants bones and report both NISP and MNI indices for all strata. In total, 4,723 bones representing at least 909 individuals of 38 species were recovered and identified. Most of the taxa were found in all levels, and cormorants (*Phalacrocorax* spp.) were the most abundant. Of the 1,567 cormorant bones excavated from site 31, those of Pelagic Shags (*P. pelagicus*) predominated (83%). Cormorants taken together were found in expected frequencies, except that they were more common than expected in level E (RA: $z_E = +4.74$, P << 0.001). Short-tailed

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Albatrosses were evident in nearly every level excavated, absent in only the earliest (deepest) levels, and occurred in proportions expected from the overall distribution of taxa through time. Northern Fulmars (*Fulmarus glacialis*) and Short-tailed Shearwaters (*Puffinus tenuirostris*) were more abundant than expected early on, in levels I and J (RA: fulmar, z_1 = +2.77; shearwater, z_1 = +2.61, z_3 = +3.25; P < 0.01). Canada Geese were present in most levels, but abundances were lower than expected in levels B and I, and greater than expected in levels E and K (RA: z_B = -2.21, z_E = +4,74, z_1 = -3.12, z_K = +3.03; *P* < 0.01). The small auklets (*Aethia* spp.) were significantly less frequent in level E and more frequent in level I (RA: z_E = -3.07, z_1 = +2.87; P < 0.01).

Overall, levels E, I, and K differed markedly from the overall pattern of bird bones excavated from site 31. Multivariate analysis (Correspondence Analysis, $\chi_c^2 = 29.5$, df = 66, P = 0.99) reported in Siegel-Causey et al. (1991) revealed that the greatest sources of divergences were the high proportion of cormorants in level E, high proportions of fulmars, shearwaters, and small auklets in level I, and greater than expected numbers of geese in level K.

When grouped into classes representing diet and foraging preferences (see Table 3), the temporal profile (Fig. 2a)—that is, the pattern of bone abundances in levels through time—of planktivorous birds feeding offshore differed significantly from birds feeding in nearshore waters (GFA: $D_{OP/N} = 0.149$, $\chi^2 = 10.6$, P = 0.005). The sources of the divergences were an overabundance of nearshore birds in level E (RA: $z_{E(N)} = +6.32$, P << 0.001) and of offshore planktivorous birds in level I (RA: $z_{I(OP)} = +3.13$, P < 0.01). The profiles of piscivorous offshore birds and nearshore birds were similar. The temporal profiles of birds classed within each foraging group were alike and indicates that the differences (and similarities) of profiles were not strongly driven by individual species.

The temporal profile of midden bones from ledge-nesting birds differed from that seen of non-breeding pelagic birds (GFA: $D_{S/X} = 0.134$, $\chi^2 = 9.07$, P = 0.011), while all other species had similar distributions to the controls through time. The sources of difference were that ledge-nesting birds were more frequent than non-breeding birds in levels E and H (RA: $z_E = +4.72$, $z_H = +3.88$; P << 0.001). Observed abundances were lower in adjacent levels later in time, i.e. levels D and G, but not significantly.

Buldir Island

Buldir Island (51°22'N 176°55'E) is the most isolated island in the Aleutian archipelago and is the westernmost island of the Rat Islands group (Fig. 1). It is the only landfall between Shemya Island, 104 km to the west, and Kiska Island, 117 km to the east (Byrd and Day, 1986; Lefèvre *et al.*, 1997). The only known archaeological site on the island was first located in 1962 (Jones, 1963), and was excavated over several years by the authors (Lefèvre and Siegel-Causey, 1993; Corbett *et al.*, 1997; Lefèvre *et al.*, 1997). The preservation of artifacts and bones is exceptional and the numerous seabird colonies were not disturbed by the introduction of foxes as was the case for many other islands (Bailey, 1993).

The first unit, Pit 1, a 1 x 1 m test pit, was located at the west end of the midden. Under thin vegetation cover, 80 cm of sterile sand was observed over a single 60 cm thick black cultural deposit, which was excavated in 10 cm horizons. Two radiocarbon dates indicate a range of occupation between the late 15^{th} and the late 17^{th} centuries (Table 2), and indicate mixing and continuous occupation rather than separate depositional events (Lefèvre *et al.*, 1997). The latest date (280¹⁴C y BP) was used as representative of the entire stratum.

The second unit, Pit 2, was located near the east end of the midden, 143 m east of Pit 1. Seven alternating cultural and sterile sandy layers were observed, each with distinct characteristics and artifact densities. Level C dated from the early 15th century and level G from the late 8th century (Table 2). Profiles of both pits and other supporting details are given in Corbett et al. 1997. In total, 6,087 bones representing at least 541 individuals of 45 species were recovered.

Analysis of the animal taxa from the Buldir Island excavations indicated a nearly equal representation of mammal and bird remains by volume, but birds predominated in abundance and diversity (Lefèvre *et al.*, 1997). Fish were less numerous and represented primarily by Scorpaeniformes, with rock greenling (*Hexagrammos lagocephalus*) as the most abundant. Pacific cod and Pacific halibut (*Hippoglossus stenolepis*) were represented by only a few vertebrae. The only mammal bones recovered were from Steller sea lions, *Eumetopias jubatus*. Table 5 gives the distribution of avian bones by stratigraphic level for both excavation sites.

At present, the breeding avifauna of Buldir comprises 32 species, 21 of them seabirds (Byrd and Day, 1986); in terms of numbers or biomass, 99.9% of the breeding birds of the island are seabirds (Lefèvre and Siegel-Causey, 1993; Lefèvre *et al.*, 1997). The results given in Table 5 reveal that there have been dramatic changes in the abundance of bones through time. Cassin's Auklet (*Ptychorhamphus aleuticus*) was found in every layer in similar abundances, except for level E of site 2 where an order of magnitude greater were recovered. This species is uncommonly encountered at present and are rarely noted to breed (Byrd and Day, 1986). Similarly, the Rhinoceros Auklet (*Cerorhinca monocerata*) was found throughout the midden levels in moderate abundances, but today is represented by

less than 20 breeding pairs. By contrast, kittiwakes (*Rissa* spp.) were found only in one level and Leach's Storm Petrels (*Oceanodroma leucorhoa*) were found in only three levels, but they are both common breeding species today on Buldir Island. Short-tailed Albatrosses were found only in two levels as well, but they are extraordinarily common in archeological contexts elsewhere in the Aleutians (Yesner, 1976; Causey, In Press). Unlike Amchitka, where cormorants were the most common species in the midden deposits, on Buldir they were rarely found in zooarcheological contexts. Cormorants breed in small numbers on Buldir and associated islets, but large numbers of non-breeding birds are commonly encountered in nearby waters (Byrd and Day, 1986).

The temporal profiles of bones deposited in the midden levels of groups based on diet and foraging preferences (Fig. 2) were all different (GFA: $D_{OF/OP} = 0.206$, $\chi^2 = 15.5$, P = 0.004; $D_{OF/N} = 0.212$, $\chi^2 = 14.6$, P = 0.007; $D_{OP/N} = 0.232$, $\chi^2 = 75.2$, P << 0.001). The sources of the differences primarily were due to the high abundance of nearshore birds in level C, offshore planktivorous birds in level D, and offshore piscivorous birds in level E. Temporal profiles of birds classed within each foraging group were alike, suggesting that the differences were associated with hunting and subsequent deposition within the midden rather than anomalies associated with particular species.

The temporal profile of midden bones from ledge-nesting and surface-nesting birds differed from that seen of non-breeding and other inaccessible birds (GFA: $D_{S-L/B-X} = 0.339$, $\chi^2 = 201.82$, $P \ll 0.001$). The sources of difference were that birds breeding in accessible colonies were less frequent in level C and more frequent in level E than non-breeding birds (RA: $z_{\rm C} = -16.96$, $z_{\rm E} = +11.63$; P $\ll 0.001$).

Shemya Island

Shemya Island (52°43'N 174°07'E) lies in the Semichi Islands group of the Near Islands (Fig. 1)-the westernmost group of Alaska's Aleutian Islands-and is a flat-topped seamount approximately 2.4 km wide and 5.7 km long on a west-east axis, with an area of 1,425 ha (3,520 acres) and a shoreline of 22 km (14 miles) on the North Pacific Ocean and Bering Sea. Among the ten archaeological sites located on Shemya Island, two (SH-1 and SH-5) are completely destroyed today, and a third, ATU-212, is represented by only a tiny remnant (Corbett *et al.*, In Press). The seven remaining sites have extensively suffered from WWII and subsequent military impact over the last 50 years, and all the upper levels of these sites have been either totally destroyed or severely disturbed. Four sites (ATU-003, ATU-021, ATU-022, ATU-061) yielded sufficient material from dated levels (Corbett et al., In Press) to allow comparison with the excavations on Amchitka and Buldir Islands (Table 5). In all, 5,482 bones were recovered representing at least 445 individuals and 37 species. Radiocarbon dates are based on the earliest levels obtained for each site (Table 3). The estimation errors associated with the radiocarbon dates obtained for sites ATU-021 and ATU-022 suggest that these may be contemporaneous; however, for analytical purposes the data from each site are reported separately.

Short-tailed Albatrosses, fulmars, and shearwaters were the most numerous birds found in midden deposits on Shemya and represented 69% of the major species identified. In the earliest excavation site ATU-061, however, they are significantly rarer than expected (RA: $z_{alb} = -6.60$; $z_{fulm} = -7.72$; $z_{shear} = -6.09$; all P << 0.001). Cormorants, which do not breed on Shemya Island but are commonly encountered in the protected nearshore waters, were moderately abundant in the middens and represented about 7% of the total. Their

presence in the midden deposits were within expected abundances, as were most of other non-alcid species.

The relative abundance patterns of alcids (e.g, murrelets, auklets, murres, puffins [*Fratercula* spp.]) complexly varied through time. For example, Ancient Murrelets (*Synthliboramphus antiquus*) were far more common in the most recent sites (ATU-021, ATU-022) than was expected, rarer than expected in ATU-003, and occurred as expected in the earliest levels (RA: $z_{021} = +2.35$, P <<0.001; $z_{022} = +3.51$, P <<0.001; $z_{003} = -3.86$, P <<0.001, $z_{061} = +0.14$, ns). By contrast, puffins showed nearly the opposite pattern, occurring far less than expected in more recent levels, and more commonly in the earliest (RA: $z_{021} = -4.54$; $z_{022} = -6.23$; $z_{003} = -5.16$; $z_{061} = +5.52$; all P << 0.001).

When the three most recent levels were treated as one, the patterns are a little clearer: the procellariids (albatrosses, fulmars, storm-petrels, and shearwaters) were much more common than the other species roughly 2,000 ybp, while the alcids were more common in ATU-061, a thousand years earlier. Cormorants, geese, eiders, and larids (gulls [*Larus* spp.]and kittiwakes) occurred in expected numbers.

Similar to the pattern observed from Buldir Island, when the birds were grouped into diet and foraging classes (Table 3), the temporal profiles were different (GFA: $D_{OF/OP} = 0.128$, $\chi^2 = 37.7$; $D_{OF/N} = 0.101$, $\chi^2 = 15.1$; $D_{OP/N} = 0.108$, $\chi^2 = 13.6$; all $P \ll 0.001$). The greatest sources of divergence were that offshore planktivorous birds were very underrepresented in the earliest site (ATU-061), offshore piscivorous birds were very abundant in site ATU-003, and nearshore birds were more common than expected in site ATU-021 (RA: $z_{OP} = -5.85$; $z_{OF} = +5.46$; $z_N = +4.49$; all P << 0.001). As before, the temporal profiles of birds classed within each group were alike.

The temporal profiles of midden bones from ledge-nesting and surface-nesting birds did not differ significantly from that of non-breeding and other inaccessible birds.

TEMPORAL PATTERNS OF CLIMATIC CHANGE

The history of relative sea level in the Aleutian Islands was summarized by Black (Black, 1974, 1977, 1980, 1982). Based on radiocarbon dating of tephra sequences, he suggested that relative Holocene sea level rose to its present level about 5000 ¹⁴C y BP on Attu Island, 6500 ¹⁴C y BP on Amchitka Island, and 11,000 ¹⁴C y BP in Cold Bay (Alaska Peninsula). Data from Shemya Island (Savinetsky *et al.*, 2004) and Bering Island (55° 12'N 165° 59'E)—450 km to the west—indicate that sea levels stabilized within \pm 2 m by 6000 y ¹⁴C bp (Razjigaeva *et al.*, 2004). Sea levels oscillated within this period (Fig. 3a), however, with peaks at about 3800 ¹⁴C y BP (+2.0 m), 2700 ¹⁴C y BP (+1.5 m), and 1000 ¹⁴C y BP (+0.9 m); and regressions at 3100 ¹⁴C y BP (-0.9 m), 1800 ¹⁴C y BP (-0.5 m), and 400 ¹⁴C y BP (-0.45 m).

These oscillations in sea level had profound local effects on island vegetation, coastlines, and reefs. On Bering Island, for example, the sea level rise of about 3800 y 14 C bp caused the development of small estuarine lagoons characterized by brackish and warmer water diatoms. A thousand years later, the time of the next sea level peak, these areas were characterized by interbedded beach sand and peat (Razjigaeva *et al.*, 2004). On Shemya Island, sea level fluctuations of ± 2 m have had great effect on the spatial extent of the intertidal and subtidal areas that provide important foraging grounds for humans and birds. For example, a drop in sea level of 0.5 m at 2880 ypb and 1800 ybp will have expanded outward by 200 m the present shoreline on the northwestern coast where the most extensive reefs on the island are located. This in turn expanded the intertidal and subtidal areas

outward another 500 m from shore, creating an additional 2500 ha of foraging habitat for birds and humans.

Reconstruction of temperature and precipitation changes (Fig. 3) was based on changes in diatom distribution and abundance from Bering Island (Razjigaeva *et al.*, 2004) and Gulf of Alaska (Heusser *et al.*, 1985), and the rate of decomposition and allogenic ash content of peat samples taken from Shemya Island (Savinetsky *et al.*, 2004). During the Early and Middle Holocene, conditions in the western Aleutians were typified by high amounts of precipitation until about 5000 ¹⁴C y BP. It then decreased monotonically until about 2000 ¹⁴C y BP when annual precipitation achieved more or less present conditions. By contrast, in the Gulf of Alaska, precipitation increased approximately 1.5 times over this period (i.e., 1500 to 2250 mmH₂0), with noticeable decreases at about 3700 ¹⁴C y BP and 2200 ¹⁴C y BP (Heusser *et al.*, 1985).

Temperatures varied over time, with peaks observed at about 5500 - 5250 14 C y BP and 3000 - 2750 14 C y BP (Fig. 3a,b). The most recent warm periods on Bering and Shemya Islands correspond well with the Medieval Warm Period of Northern Europe (ca. AD 900 – 1350); in the eastern Aleutians and Gulf of Alaska this was a period of moderate cooling or stable temperatures (Fig. 3c). Temperatures dropped considerably from 3000 – 2000 14 C y BP in the western Aleutians and Commander Islands (Fig. 3a, b), and then stabilized except for a 250 y period experienced throughout the Aleutians and Gulf of Alaska corresponding to the Little Ice Age of Europe (ca. AD 1350 - 1800).

These data support other paleoclimate reconstructions for the region (Bartlein *et al.*, 1998; Mann *et al.*, 1998; Calkin *et al.*, 2001; Gorbarenko *et al.*, 2002; Gorbarenko *et al.*, 2004), which indicate that by the mid-Holocene (i.e., \sim 5000 ybp) the climate of the

Aleutians was changing from relatively warm and dry conditions to the current cool and wet pattern. This change was associated with an intensification of the Aleutian low-pressure center and a corresponding increase in storms and cyclonic activity (Stabeno and Reed, 1992; Niebauer *et al.*, 1999). The role that the ocean environment plays in the interaction between climate and oceanography is less clearly understood (Schumacher and Alexander, 1999), but it is clear from several studies that ocean temperatures, salinities, and currents were similarly variable in the Gulf of Alaska, Aleutian Islands, and North Pacific during the late Holocene (Francis and Hare, 1994; Francis *et al.*, 1998; Finney *et al.*, 2000; Calkin *et al.*, 2001; Tunicliffe *et al.*, 2001; Finney *et al.*, 2002).

Climate change and the response of the oceanic ecosystem to climatic forcing may affect marine foodwebs in both directions: for example, significant shifts in phytoplankton production have been linked with the NE Pacific climatic regime shift of 1976/1977 (Brodeur and Ware, 1992; Polovina *et al.*, 1995; Roemmich and McGowan, 1995), but direct effects on top predators are much more difficult to detect. It is much more likely that indirect responses related to prey availability and abundance of top-trophic level species like birds and mammals will be most evident (Francis *et al.*, 1998). The significant declines in piscivorous marine birds and marine mammals in the Gulf of Alaska since the 1980s have been attributed to fundamental shifts in the distribution and abundance of their food resources (Springer, 1992; Piatt and Anderson, 1995), a pattern well documented throughout Beringia and the North Pacific (Montevecchi, 1993; Decker *et al.*, 1995; Hunt and Byrd, 1999; Jones *et al.*, 2002). For example, shearwaters, murres, and auklets responded quickly to changes in the California Current during 1985-1994 in ways predictable by their morphology and natural history (Oedekoven *et al.*, 2001). This was particularly evident by

lower population sizes (Sooty Shearwaters *Puffinus griseus*, Cassin's Auklets), shifting distributions to avoid areas of low productivity (shearwaters), and feeding on alternative prey (Common Murre). If similar patterns hold for the past, then it is likely that the paleodynamics of the avifauna collected by early human hunters should reflect the paleodynamics of the Aleutian Island ecosystem.

TEMPORAL AND SPATIAL PATTERNS OF LATE HOLOCENE AVIFAUNA

What is immediately evident from the data and analyses presented here is that the populations of very few species in the central and western Aleutians were constant during the past 3,000 years. For example, on Shemya and Buldir Islands the fish-eating alcids, murres and puffins, varied considerably and inconsistently through time: during some periods, murres predominated and during others, puffins. It is difficult to ascertain whether this variability is due to changes in food availability, changes in offshore or onshore environments, or changes associated with indirect effects on breeding habitats. A few patterns stand out, however: procellariids (Short-tailed Albatrosses and Short-tailed Shearwaters) were present and abundant in every level at Amchitka and Shemya Islands. During periods of increased temperatures (e.g., Medieval Warming Period, 650 – 1100 ybp) and precipitation, nearshore foragers rose in abundance (Fig. 2, Spearman rank correlations: $r_s(T)_N = +0.518$, $r_s(P)_N = +0.571$; p < 0.05). In addition, during periods of cooling (e.g., Little Ice Ages, 200 – 650 ybp and 1800 – 2100 ybp), offshore piscivorous birds—including the aforementioned procellariids—similarly increased (Fig 2, same test: $r_s(T)_{OF} = -0.696$, p < 0.04; $r_s(T)_{Proc} = -0.554$; p < 0.05).

These sustained and large-scale patterns over millennia track with analogous studies focused on historical changes in fisheries productivity elsewhere in the North Pacific.

Finney et al. (2002) identified significant multi-century shifts in fish abundance at about 2100 ybp and 800-1200 ybp, which corresponded to periods of major change in the atmospheric and oceanic circulation. The earliest period (~100 BC) was transitional: some areas of the Northeast Pacific (i.e., Gulf of Alaska, Santa Barbara basin) by this time were characterized by warming waters (Heusser *et al.*, 1985; Brodeur and Ware, 1992), while in the Northwest Pacific (Gorbarenko *et al.*, 2004; Razjigaeva *et al.*, 2004; Savinetsky *et al.*, 2004) the warming period had not yet begun. During the centuries immediately preceding this time (i.e., 2200 ybp and earlier), the climate and oceanography was cooler and more turbulent, and characterized by change at the largest scales.

Marine birds were associated with environmental change at the highest scales. Previously, in a separate analysis of twelve studies based on material recovered from all excavated Aleutian middens, Savinetsky *et al.* (2004) found that most late Holocene marine bird populations in the southern Bering Sea region demonstrated a negative correlation with temperature and a positive correlation with precipitation, but small sample sizes made statistical estimates unreliable in their study. We undertook a similar analysis and found that for the western Aleutians, a similar pattern obtained (Spearman rank correlations: $r_s(T)_{ALL} =$ -0.506, $r_s(P)_{ALL} = +0.601$; p < 0.075).

Other broad scale patterns are obscured by variability in species-level occurrences found in individual levels and sites and local anomalies abound. For example, the excavations from Buldir Island revealed pulses of abundance in the planktivorous species throughout time, which may reflect a corresponding increase in population sizes or chance successes by early Aleut hunters. Also on Buldir, Aleutian Canada Geese were rarely encountered in midden deposits by contrast to Shemya and Amchitka Islands, yet at present the island represents one of the most critical breeding habitats for the species. Additional studies using higher-resolution reconstruction of local and regional climates, larger sample sizes, and a broader taxonomic base including fish and invertebrates, are needed to clarify and further refine these findings.

HUMAN IMPACTS VS. INTRINSIC CHANGE IN NEAR SHORE ECOSYSTEMS

During precontact times, the Aleutian Islands supported one of the world's densest concentrations of humans with a hunting/fishing/gathering economic base (Liapunova, 1979; Corbett *et al.*, 2001; Corbett *et al.*, In Press). Exact population sizes are unknown due to declines following contact with Russian explorers in the mid-1700s; estimates range from 7,000 to over 30,000 in the Aleutian chain (Laughlin, 1980; Lantis, 1984). The distribution of early Aleut groups (Fig. 1) corresponds almost exactly with the location of major islands passes, for example the territories of the Sasignan and Qaxun Aleut groups were bounded by the Near Islands, Buldir, and Amchitka passes (Bergsland and Dirks, 1990). Island passes define the physical and biological structure of the Aleutian Islands (Byrd *et al.*, In Press; Ladd et al., in press; Sinclair et al., in press), and helped influence the development of human cultural preferences. For example, the Sasignan Aleut evolved a unique and distinctive hunting technology that was finely tuned to the biological conditions of their environment, and was unlike in many aspects from technologies and material cultures found in their neighbors, the Quxun, Naahmigus, and Niigugis Aleuts (McCartney, 1984; Corbett et al., 1997; Veltre, 1998; McCartney and Veltre, 1999; Corbett et al., In Press). Our evidence suggests that the early cultural structure of the central and western Aleutians reflects the environmental structure, that is, the biodiversity, geography, and physical oceanography.

The size of a human population obviously has a direct correlation with the impact of that population on its environment and local depletion of vertebrates used for food have been identified for many early maritime cultures (Gonzalez, 1999; Hildebrandt and Jones, 2002; Broughton and Bayham, 2003). We detected hunter-related depletions of populations breeding in accessible colonies at small scales of space and time, but we did not observe widespread or long-term effects. For example, cormorants, gulls, and murres were harvested in significantly greater numbers than the control group at levels E and H on Amchitka Island and level E on Buldir Island, and were subsequently rarer afterwards. These species were used for food and clothing, and nest on the flat shoulders of cliffs or on wide ledges, often on the flat tops of islets and rock stacks (Whittam and Siegel-Causey, 1981; Lefèvre et al., 1997). Disturbance by harvesting tends to drive the nesting birds to more isolated spots and diminish their availability during breeding season, but solitary birds will continue to forage in the intertidal and nearshore waters far past the breeding season and into winter (Meehan et al., 1996). Numerical abundance of affected species in these cases will decrease but local populations will rarely vanish from a region.

Another example, but based in more recent human activities, is illustrated by the abundance of Short-tailed Albatrosses in the Aleutian Islands. At some localities in the eastern Aleutians (e.g., Kodiak Island, Umnak Island), this species was the most abundant bird in terms of harvested biomass for over 3,000 years (Yesner, 1976; Causey, In Press). These patterns were also seen in the central and western Aleutians, and their abundances as reflected by presence in the midden levels were high yet varied insignificantly through time (Kolmogorov-Smirnov One-Sample tests, P > 0.10). The abrupt drop in abundance in the

past century is due entirely to hunting pressure and habitat destruction by European hunters in the early years of the 20th century (Carboneras, 1992; Causey, In Press).

Overall, however, historical hunting practice of the early Aleut hunters did not seem to have a demonstrable impact on the seabird community in the Aleutian Islands. Results shown here indicate that while some changes have taken place, they are not consistent with long-term effects associated with local extirpation by overhunting, selective harvesting, or habitat perturbation. Instead, it appears that trends in abundance for many species have been locally stable (e.g, Short-tailed Albatross, and Pelagic and Red-faced cormorants) or changing in similar patterns over large geographic scales. This is best shown where we have nearly contemporaneous material from approximately 2000 – 2500 ybp on Amchitka and Shemya Islands. These contemporaneous patterns separated by over 800 km suggest ecosystem changes rather than coordinated shifts in hunting patterns among the early Aleut occupants.

The diversity of the marine bird fauna of the central and western Aleutians 3,000 years ago was similar to what it is today, but populations differ primarily in their relative abundance. Then as now, these patterns of species change were very likely driven by environmental changes, particularly by those parameters directly affecting the abundance and distribution of food. If present patterns are indicative of past processes, it appears that local oceanography and regional changes in prey bases had a significant impact on the distribution and abundance of Aleutian birds and mammals.

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Site ¹	Location ²	Region and	Material ⁴	Species	Sampling	Inclusive	Levels	Metadata ⁹	Reference ¹⁰
		Pass ³		ID ⁵	Design ⁶	Dates (ybp) ⁷	Dated ⁸		
Chugachik I,	59° 45′ N	G	B, M, I	\sim	ð	1700 - 2300	No	No	Yesner 1977
Kachemak Bay	151° 03′ W	Gulfof							
		Alaska							
Kodiak I	Various	Ð	B, M	IJ	U	Unknown	No	No	Friedmann, 1935 –
		Gulfof							1938,
		Alaska							Hrdlicka 1945
Larsen Bay,	57° 56' N	U	B, M, I	∞	0	620 - 1290	No	No	Yesner 1989
Kodiak I	154° 02' W	Gulfof							
		Alaska							
Peterson Lagoon,	54° 45' N	Е	Ч	S	0	100 - 1000	Yes	Γ	Hoffman et al. 2000
Unimak I	165° 58' W	Unimak Pass							
Amoknak I	53° 53' N	Щ	M, I	\mathbf{N}	Ŋ	Unknown	No	No	Eyerdam 1934

Table 1. Aleutian Zooarcheological Studies.

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166° 32' W Unalaska

		Pass							
Chaluka, Umnak	59° 45′ N	Щ	М	IJ	ð	200 - 3700	Μ	No	Lippold 1966, 1972
Island	151° 03′ W	Samalga							
		Pass							
Oglodax,	59° 45′ N	Е	B, M	\mathbf{S}	Ŏ	200 - 2400	No	No	Yesner and Aigner
Anangula, Umnak	151° 03′ W	Samalga							1976
Island		Pass							
Atka I	52° 12′ N	C	B, M	Μ	ð	500 - 1900	Yes	No	Veltre 1979
	174° 12' E	Atka Pass							
Amchitka Island	51° 23′ N	C	B, M	\mathbf{S}	Ø	890 - 2650	Yes	Γ	Harrington 1987,
	179° 16' E	Amchitka							Siegel-Causey et al.
		Pass							1991
Buldir Island	52° 22′ N	M	B, M, F, I	S	ð	460 - 1160	Yes	L	Lefevre et al. 1977
	176° 55' E	Buldir Pass							
Shemya Island	52° 43′ N	M	B, M, F, I	S	ð	1790 - 3540	Yes	L	Siegel-Causey et al.
	174° 07' E	Attu Pass							1994, Corbett et al.

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In Press

- 1 Sites identified by geographic name and ordered from East to West.
- Geographic location is mid-point for densely aggregated archeological sites (e.g., SW Umnak Island) 2
- C: Central Aleutians; G: Gulf of Alaska; E: Eastern Aleutians; W: Western Aleutians ξ
- 4 B: birds, M: mammals, F: fish, I: invertebrates.
- M: Missing, identified only by group (i.e., "fish"); G: General, identified or accurate to genus (i.e., "Gull" or "Larus"); S: Species, S

identified or accurate to species.

- 6 U: Uncontrolled; Q: Quadrat test pits.
- 7 Dates represent ranges of uncorrected radiocarbon dates associated with faunal deposits.
- 8 Individual excavation levels with published dates.
- 9 No: metadata missing or unreported; L: Faunal material listed by levels or dated
- 10 Unpublished work is not listed.

Island	Site	Level	Date	Lab Reference
Amchitka	RAT-031	В	890 ± 90	I-4736 (1)
		Н	1980 ± 95	I-4737 (1)
		J	2550 ± 95	I-4735 (1)
Buldir	KIS-008, pit 1	1	460 ± 50	Beta-54253 (2)
		6	280 ± 50	Beta-54254 (2)
	KIS-008, pit 2	С	530 ± 60	Beta-54255 (2)
		G	1160 ± 50	Beta-54256 (2)
Shemya	ATU-003	7	2556 ± 135	IEMAE-1177 (3)
	ATU-021	2	1980 ± 60	Beta-131313 (2)
	ATU-022	1	2020 ± 110	Beta-110114 (2)
	ATU-061	Bottom	3325 ± 60	Beta-39104 (2)

Table 2. Uncorrected radiocarbon dates (ybp ¹⁴C) for levels and sites used in this study.

1 Unknown laboratory (Desautels et al. 1971).

- 2 Beta Analytic (Corbett et al. 1997).
- 3 Institute of Ecology and Evolution, Russian Academy of Sciences, Moscow

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Table 3. Nesting, Foraging, and Diet strategies of common seabirds found in the Aleutian Islands. Data are based on Byrd et al. (in

press).

Species	Common	Nesting ¹	Foraging ²	Diet ³
Aethia cristatella	Crested Auklet	C	0	Ь
Aethia pusilla	Least Auklet	C	0	Ρ
Aethia pygmaea	Whiskered Auklet	C	Z	Ρ
Brachyramphus marmoratus	Marbled Murrelets	S	Z	Ц
Branta canadensis	Canada Goose	S	Z	Λ
Cyclorhamphus psittacula	Parakeet Auklet	C	Z	Р
Fratercula cirrhata	Tufted Puffin	В	0	Ч
Fratercula corniculata	Horned Puffin	C	Z	Ч
Fulmarus glacialis	Northern Fulmar	S	0	Ρ
H histrionicus	Harlequin Duck	S	Z	Λ
Larus spp.	Gulls	S	Z	Щ

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Ь	Ц	Щ	Ь	Ц	Ц	Ц	Ц	Ь	Ц
C	7	С	С	C	С	7	7	С	0
В	Γ	X	В	X	Γ	\mathbf{N}	\mathbf{N}	В	Γ
Storm-petrel	Cormorants	Short-tailed Albatross	Cassin's Murrelet	Short-tailed Shearwater	Kittiwakes	Common Eider	Aleutian Tern	Ancient Murrelet	Murres
Oceanodroma spp	Phalacrocorax spp	Phoebastria albatrus	Ptychorhamphus aleuticus	Puffinus tenuirostris	Rissa spp	Somateria mollissima	Sterna aleuticus	Synthliboramphus antiquus	<i>Uria</i> spp

2 N: nearshore, O: offshore.

B: burrow, C: crevice, L: ledge, S: surface, X: non-breeder in Aleutians.

3 F: fish, P: plankton, V: vegetarian

						Site 31				
Levels ¹	A	В	С	E	U	Η	Ι	ſ	K	Total ²
		890			1,890			2,650		
Date ¹⁴ C ybp ³		ybp			ybp			ybp		
Phoebastria albatrus	n	6	6	4	ю	5	9	7	0	293(41)
Fulmarus glacialis	0	7	3	ŝ	1	7	10	1	1	75(23)
Puffinus tenuirostris	7	9	7	7	7	7	12	С	0	178(31)
Phalacrocorax auritus ⁴	0	9(3)	2(2)	1(1)	2(2)	0	2(1)	0	0	16(9)
Phalacrocorax pelagicus ⁴	12(3)	323(29)	217(25)	363(48)	55(6)	94(15)	199(25)	9(3)	22(3)	1294(157)
Phalacrocorax urile ⁴	1(1)	76(14)	33(7)	76(12)	14(3)	18(3)	35(5)	(0)0	4(1)	257(50)
Branta canadensis	1	٢	12	29	5	С	7	0	9	23(65)
Somateria mollissima	1	11	6	L	5	4	10	1	7	279(50)
Larus glaucescens	1	9	\mathfrak{S}	5	ς	7	4	1	З	143(28)

Table 4. Distribution of identified avian bony elements excavated from Amchitka Island, site 31, by stratigraphic level. Data are

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										3/15/05
<i>Uria</i> spp.	1	8	10	7	7	4	9	1	1	201(40)
<i>Aethia</i> spp.	4	56	44	39	11	16	43	9	8	557(146)
Misc spp.	8	62	51	44	12	20	53	7	12	1407(269)
NISP	26	201	161	175	63	71	178	24	39	4723(909)
1 A – M, discrete level	s. Level	5 D, F, L, 2	ind M are n	iot reported	here beca	use of smal	l samples (< 5 elemen	ts).	

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2 Totals abstracted from Tables 5 and 7 in Harrington (1987).

3 Radiocarbon dating for levels B, G, and J reported by Desautels et al. (1971).

4 Data abstracted from Siegel-Causey et al. (1991).

Table 5. Distribution of identified avian bony elements excavated from Buldir Island by stratigraphic level. Data are from Lefèvre et

al. 1997.

	Pit 1				Pit 2				
Level	Total	V	B	C	D	E	Ч	G	
	280			530				1160	Total
Species	ybp ¹			ybp				ybp	
Phoebastria albatrus	2 (2) ²	0	0	1 (1)	0	0	0	3 (1)	6 (4)
Fulmarus glacialis	3 (1)	0	0	0	0	0	0	0	3 (1)
Oceanodroma spp.	59 (13)	0	0	0	0	85 (19)	1 (1)	0	145 (33)
Phalacrocorax spp.	14 (5)	0	0	1 (1)	0	0	0	0	15 (6)
Branta canadensis	18 (7)	0	0	4 (2)	0	51 (21)	0	0	73 (30)
Larus spp.	61 (14)	0	3 (1)	0	0	11 (5)	0	0	75 (20)
<i>Rissa</i> spp.	119 (19)	0	0	0	0	0	0	0	119 (19)

								Cause	sy et al. 3/15/05
Uria spp.	14 (6)	2 (1)	0	2 (1)	0	0	0	0	18 (8)
Synthliboramphus antiquus	250 (35)	6 (1)	12 (5)	0	3 (1)	0	17 (3)	20 (4)	308 (49)
Ptychoramphus aleuticus	25 (7)	2 (1)	17 (2)	6 (2)	4 (1)	410 (54)	14 (3)	37 (4)	515 (74)
Cyclorrhynchus psittacula	57 (15)	0	0	2 (1)	0	0	1 (1)	0	60 (17)
<i>Aethia</i> spp.	525 (90)	11 (2)	16 (5)	311 (32)	14 (4)	405 (64)	20 (4)	25 (5)	1327 (206)
Cerorhinca monocerata	11 (3)	1 (1)	7 (2)	14 (6)	0	147 (21)	3 (1)	8 (3)	191 (37)
Fratercula spp.	9 (4)	0	10 (3)	44 (7)	24 (3)	87 (18)	2 (1)	1 (1)	177 (37)
Misc spp. ³	361	21	46	530	58	1906	43	06	3037
Totals	1528 (221)	43 (6)	111 (18)	915 (53)	103 (9)	3102 (202)	101 (14)	184 (18)	6087 (541)

¹ Radiocarbon dating reported by Corbett et al. 1997.

2 NISP (MNI).

3 MNI were not calculated for miscellaneous and unidentifiable bone fragments.

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Table 6. Distribution of identified avian bony elements excavated from Shemya Island by excavation pits. Data are from Corbett et

al. In Press.

Excavation	ATU-021	ATU-022	ATU-003	ATU-061	
Species	1980 ybp	2020 ybp	2556 ybp	3325 ybp	Totals
Phoebastria albatrus	$338(10)^{1}$	321 (17)	175 (10)	27 (4)	861 (41)
Fulmarus glacialis	54 (8)	254 (38)	117 (14)	5 (1)	430 (61)
Oceanodroma spp.	22 (7)	2 (1)	5 (2)	1 (1)	30 (11)
Puffinus spp.	204 (25)	371 (68)	313 (56)	56 (7)	944 (156)
Phalacrocorax spp.	101 (8)	61 (7)	30 (5)	27 (4)	219 (24)
<i>Branta</i> spp. ²	2 (1)	4 (1)	13 (2)	1 (1)	20 (5)
Somateria spp.	9 (2)	20 (4)	8 (2)	2 (1)	39 (9)
Larus spp.	11 (3)	23 (2)	7 (4)	8 (3)	49 (12)
Rissa spp.	1 (1)	15 (3)	4 (2)	2 (2)	22 (8)

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Uria spp.	18 (5)	34 (5)	14 (6)	47 (6)	113 (22)
Synthliboramphus antiquus	107 (9)	78 (18)	17 (6)	16 (3)	218 (36)
Ptychoramphus aleuticus	0	0	7 (2)	1 (1)	8 (3)
Cyclorrhynchus psittacula	0	0	0	8 (3)	8 (3)
<i>Aethia</i> spp.	43 (9)	32 (12)	14 (5)	50 (3)	139 (29)
Cerorhinca monocerata	0	6 (3)	4 (1)	6 (3)	16 (7)
Fratercula spp.	15 (3)	19 (4)	6 (3)	63 (8)	103 (18)
Misc spp. ³	750	482	366	665	2263
Totals	1675 (91)	1722 (183)	1100 (120)	985 (51)	5482 (445)

1. NISP (MNI).

- 2. Indeterminate; may represent Branta canadensis, B. nigricans, or Chen canagica.
- 3. MNI were not calculated for miscellaneous and unidentifiable bone fragments.

Figure Captions

Figure 1. Map of island groups and islands passes of the Aleutian Islands and location of precontact Aleut tribal groups (redrawn from Byrd *et al.*, In press, and Bergsland and Dirks, 1990). Passes are indicated by number: 1) Near Island Pass; 2) Buldir Pass; 3) Amchitka Pass; 4) Tanaga Pass; 5) Seguam Pass; 6) Amukta Pass; 7) Umnak Pass; 8) Samalga Pass; 9) Akutan Pass; 10) Unimak Pass.

Figure 2. Cumulative and relative (inset) frequencies of bird bones in midden deposits. Solid black: nearshore species; shaded: offshore, piscivorous species; open: offshore, planktivorous species. See Table 3 for species assignments.

Figure 3. Paleoenvironment of the Aleutian Islands reconstructed by several proxy studies. Shaded boxes indicate the approximate period of the Medieval Warm Period (MWP) and the Little Ice Age (LIA). A) Bering Island: bold dashed-dotted line is sea level change (m) from present, dashed line is temperature change (°C) from present, based on standard transfer functions of diatom presence and abundance. Data redrawn from Razjigaeva *et al.* (2004). B) Shemya Island: solid line is percent decomposition of dated peat samples (precipitation), dashed line is percent ash content of peat (temperature). Data redrawn from Savinetsky *et al.* (2004). C) Gulf of Alaska: solid line is mean annual precipitation (mm H₂O), dashed line is mean July temperature (°C), based on standard transfer functions of diatom presence and abundance. Data redrawn from Heusser *et al.* (1985).





