

# Snowpack-Mediated Aerosol-Climate Interactions

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(On the Web at [http://dust.ess.uci.edu/smn/smn\\_clm\\_snw\\_aer.pdf](http://dust.ess.uci.edu/smn/smn_clm_snw_aer.pdf))

## Abstract

1. Examine forcing of and response to snowpack heating by dust (natural) and soot (FF/BF/BB) in current and LGM climates
2. Quantify relative forcing efficacies of dust, soot in snowpack
3. Speculate about dust's role in glacial terminations

## 1. Absorbing Aerosols in Cryosphere

1. Warren and Wiscombe (1980); Clarke and Noone (1985): Soot perturbs *in situ* snowpack reflectance
2. Jacobson (2004): FF/BF snowpack forcing warms climate  $\sim 0.06$  K
3. Hansen and Nazarenko (2004); Hansen et al. (2005): Soot-snowpack forcing  $0.08 \text{ W m}^{-2}$  warms climate  $0.065$  K. Efficacy  $E \approx 1.7 \times \text{CO}_2$
4. Flanner et al. (2006): Soot-snowpack forcing Efficacy  $E \approx 1.8 \times \text{CO}_2$

This study examines dust *and* soot snowpack in current and LGM climates:

1. Current Climate: Dust+Soot-snowpack forcing  $0.050 \text{ W m}^{-2}$  warms climate  $0.16$  K ( $\sim 50\%$  is Anthropogenic). Efficacy  $E \approx 4.2 \times \text{CO}_2$
2. LGM Climate: Dust+Soot-snowpack forcing  $0.28 \text{ W m}^{-2}$  warms climate  $0.95$  K. Efficacy  $E \approx 4.8 \times \text{CO}_2$
3. Dust+Soot forcing efficacy very high due to
  - (a) Representing positive feedbacks in snowpack
  - (b) More absorption to trigger snow-ice-albedo feedback non-linearity
  - (c) Deposition timing and location

## 2. Methods

1. Community Atmosphere Model with Slab Ocean Model (CAM/SOM)
2. Industrial Carbonaceous aerosol emissions ([Bond et al., 2004](#))
3. Global Fire Emissions Database (GFED) ([van der Werf et al., 2003](#); [Randerson et al., 2005](#))
4. BC/OC Emissions factors ([Andreae and Merlet, 2001](#), updated)
5. BC Optical properties ([Chang and Charalampopoulos, 1990](#); [Bond and Bergstrom, 2005](#); [Bond et al., 2006](#))
6. Dust Entrainment and Deposition model (DEAD) ([Zender et al., 2003](#); [Mahowald et al., 2006b](#))
7. Glaciogenic dust sources for LGM ([Mahowald et al., 2006a](#))
8. Dust Optical Properties ([Sokolik and Toon, 1999](#); [Dubovik et al., 2002](#); [Sinyuk et al., 2003](#))
9. SNow, ICe, and Aerosol Radiative model (SNICAR) ([Flanner and Zender, 2005, 2006](#); [Flanner et al., 2006](#))
10. Experiment(Control): Dust+Soot are (not) radiatively active in snow-pack

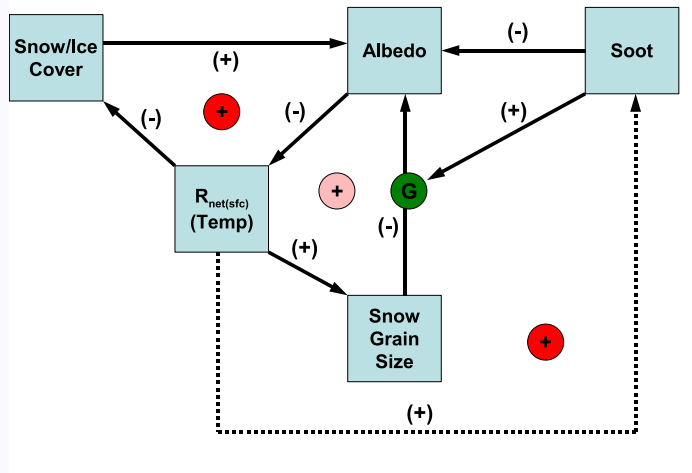


Figure 1: Soot and Dust amplify snow-albedo feedback via multiple paths.

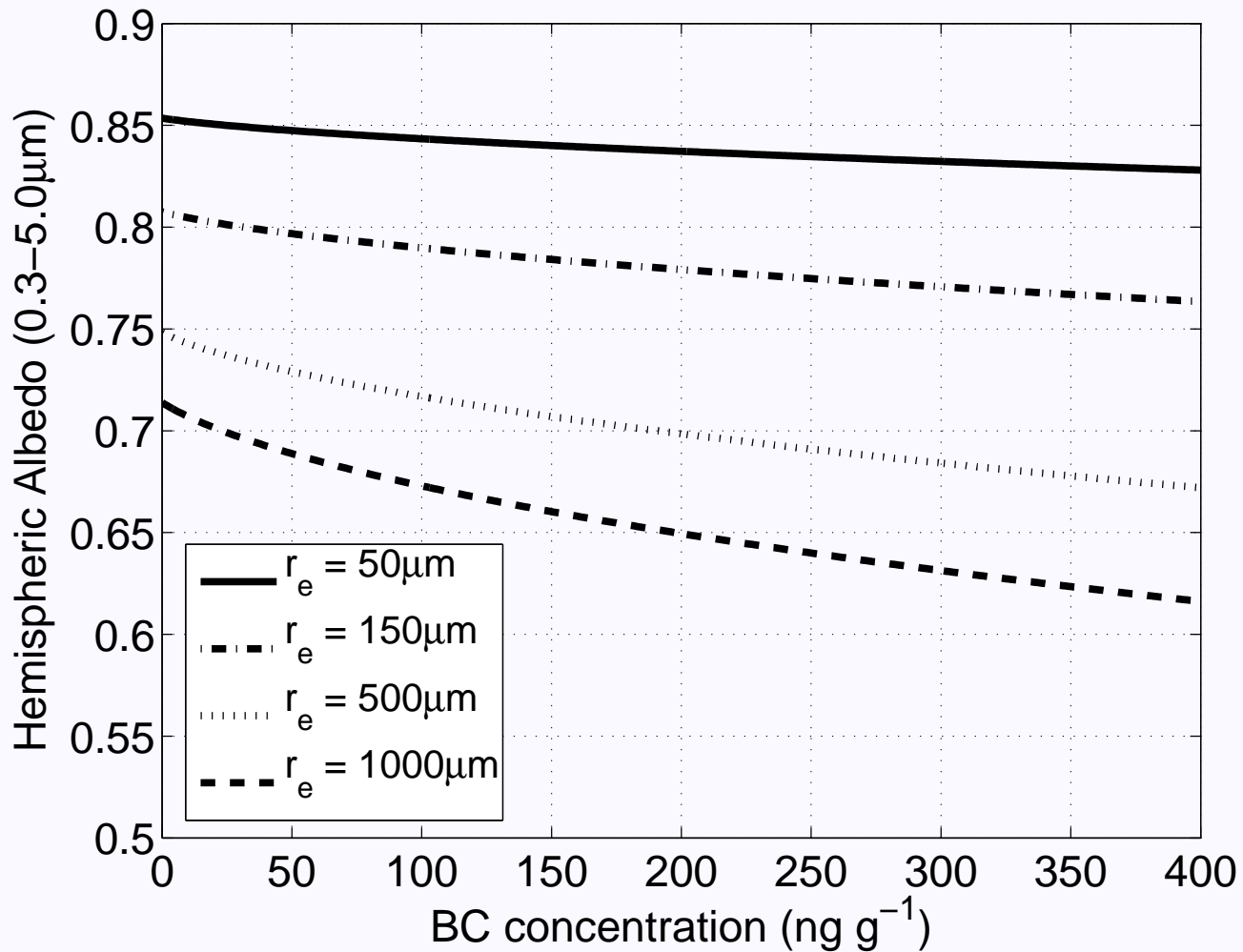


Figure 2: Hemispheric broadband albedo decreases with soot concentration and snow effective radius. (Flanner and Zender, 2006, *JGR*)

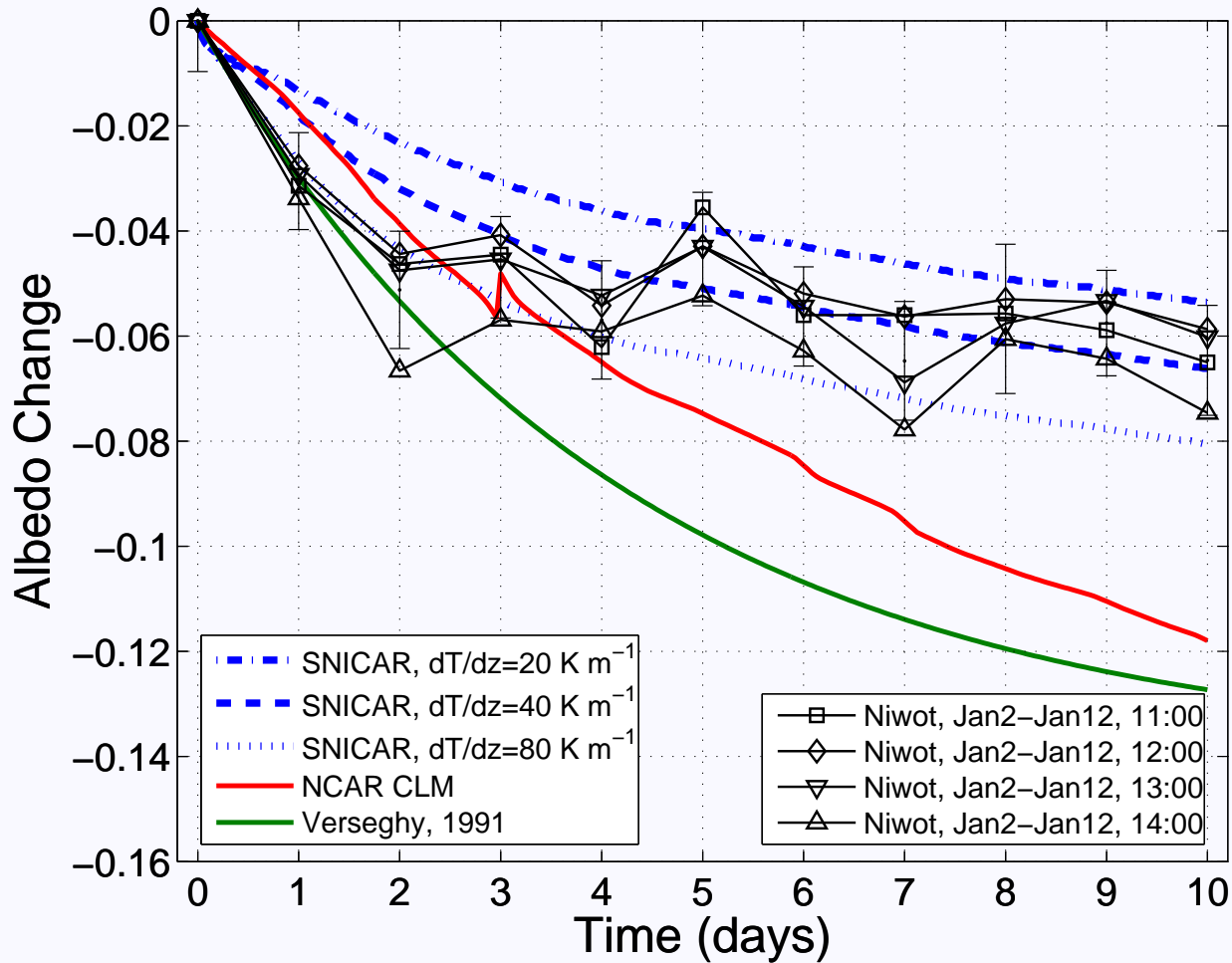


Figure 3: Observed and modeled albedo decay at Niwot Ridge January 2, 2001. (Flanner and Zender, 2006, *JGR*)

## Dust Refractive Index: AERONET, TOMS, Model

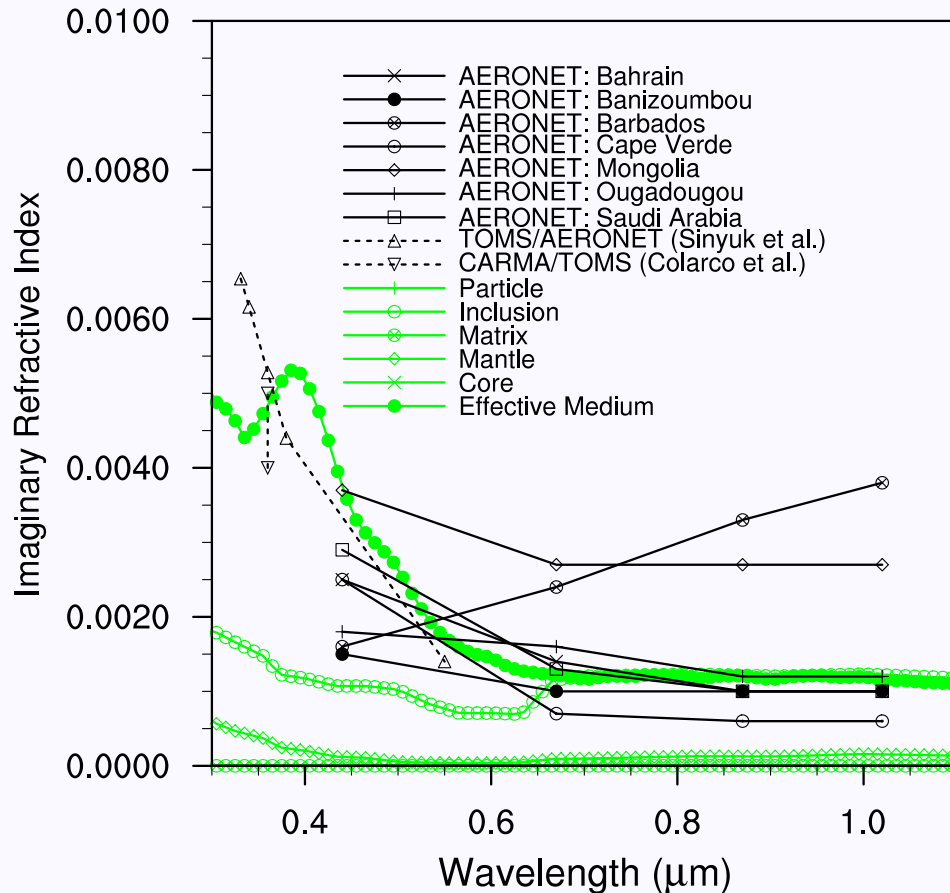


Figure 4: Refractive index and single scattering albedo for dust based on 48% Quartz, 25% Illite, 25% Montmorillonite, 1% Limestone, 1% Hematite. Data from [Dubovik et al. \(2002\)](#), [Colarco et al. \(2002\)](#), and [Sinyuk et al. \(2003\)](#).





Figure 5: February 15 dust-fall in snowpack near Niwot Ridge, Colorado on May 17, 2006. (Courtesy Tom Painter, NSIDC)



Figure 6: Dust and soot on snow near Niwot Ridge, Colorado on May 17, 2006. Soot deposited by hand same day. Dust-fall February 15. (Courtesy Tom Painter, NSIDC)

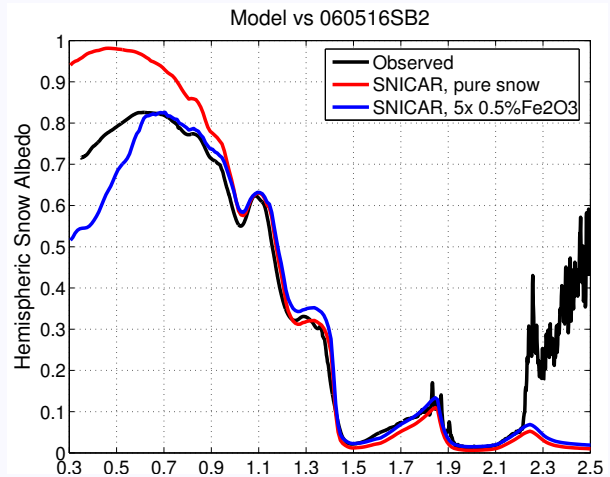
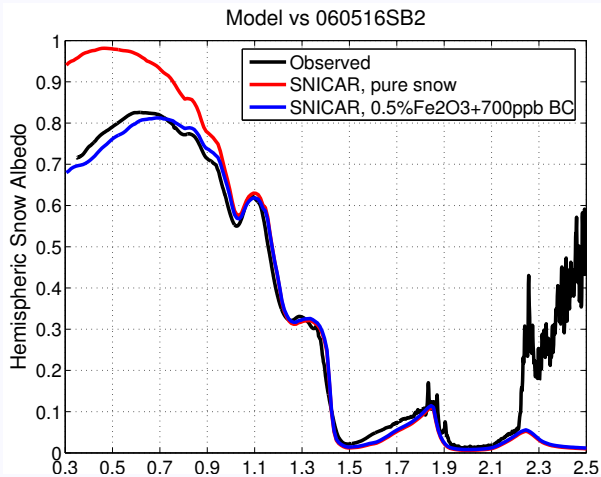
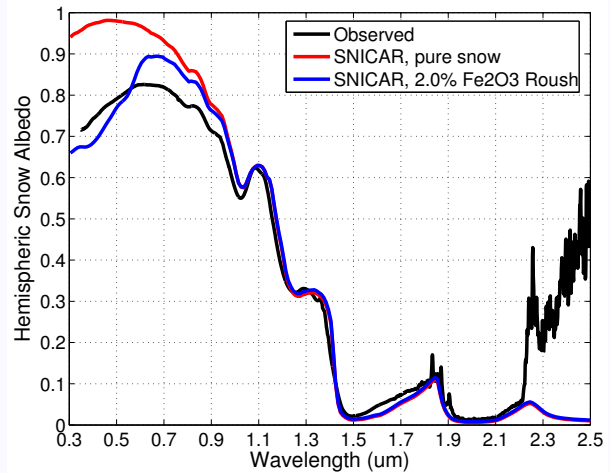
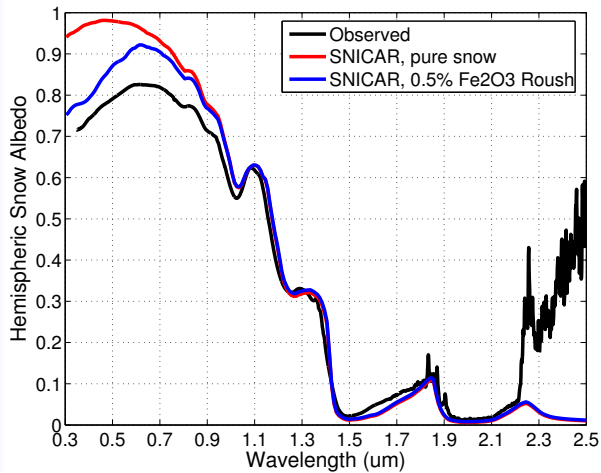


Figure 7: Measured (black) and modeled snowpack spectral hemispheric albedo  $\mathcal{R}(\lambda)$ . Model with (blue) and without (red) measured  $100 \mu\text{g g}^{-1}$  dust contribution. (Courtesy Tom Painter, NSIDC)

## Present BC/Snow Surface Concentration

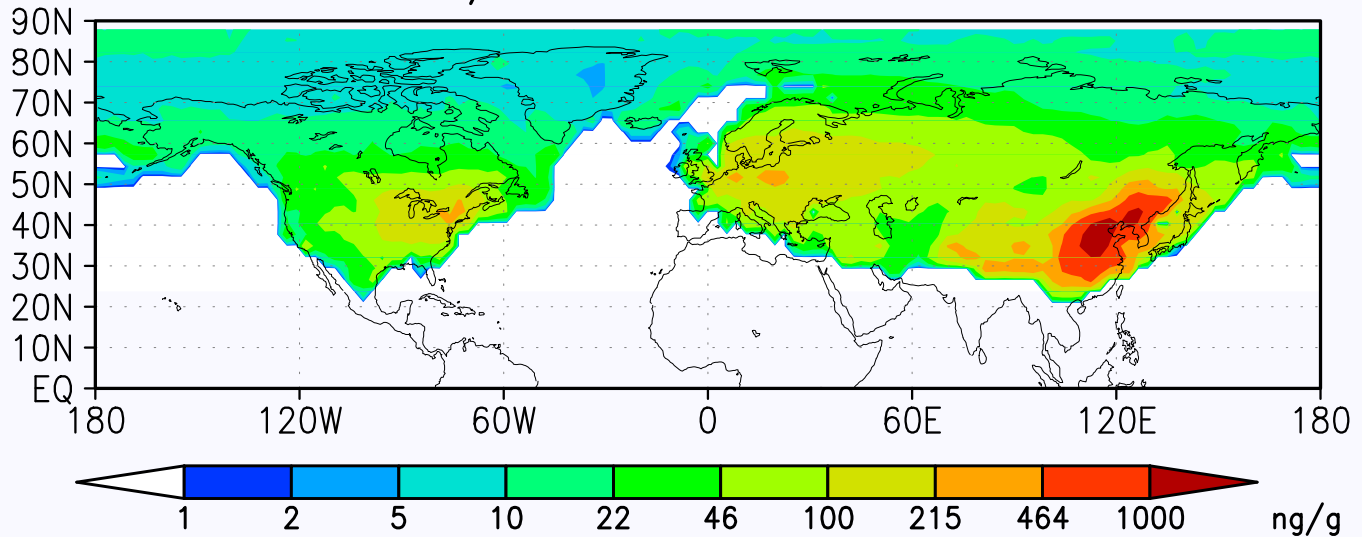


Figure 8: Present day soot concentration in snowpack, 1998 emissions. Global mean  $\sim 24 \text{ ng g}^{-1}$ .

# Present Dust/Snow Surface Concentration

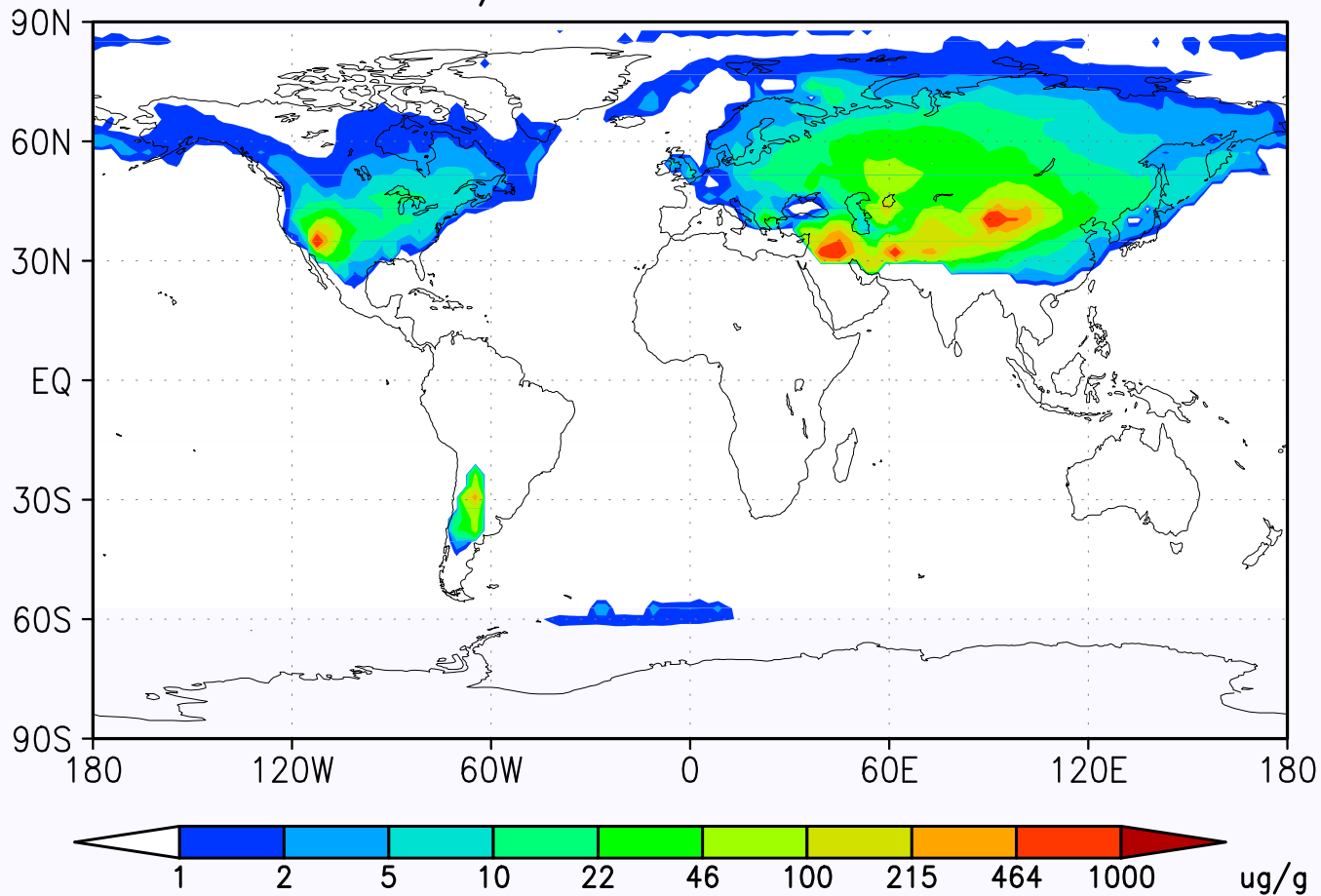
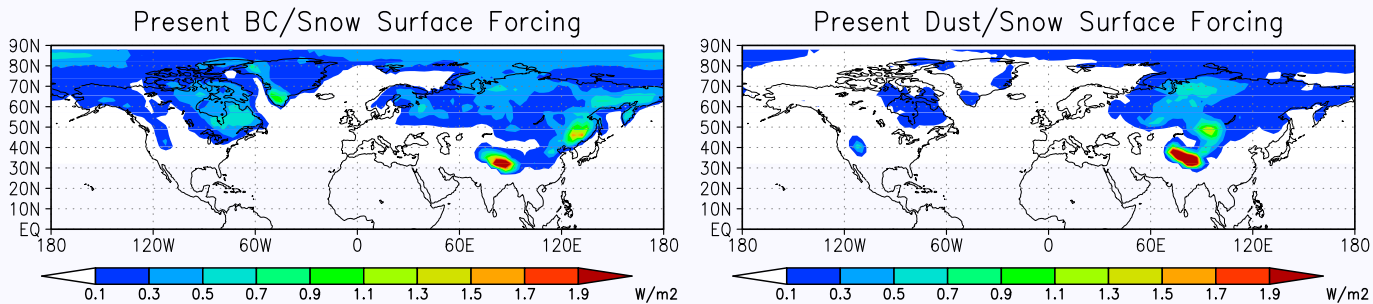


Figure 9: Present day dust concentration in snowpack. Global mean  $\sim 5.7 \mu\text{g g}^{-1}$ .



## Present BC+Dust/Snow Surface Forcing

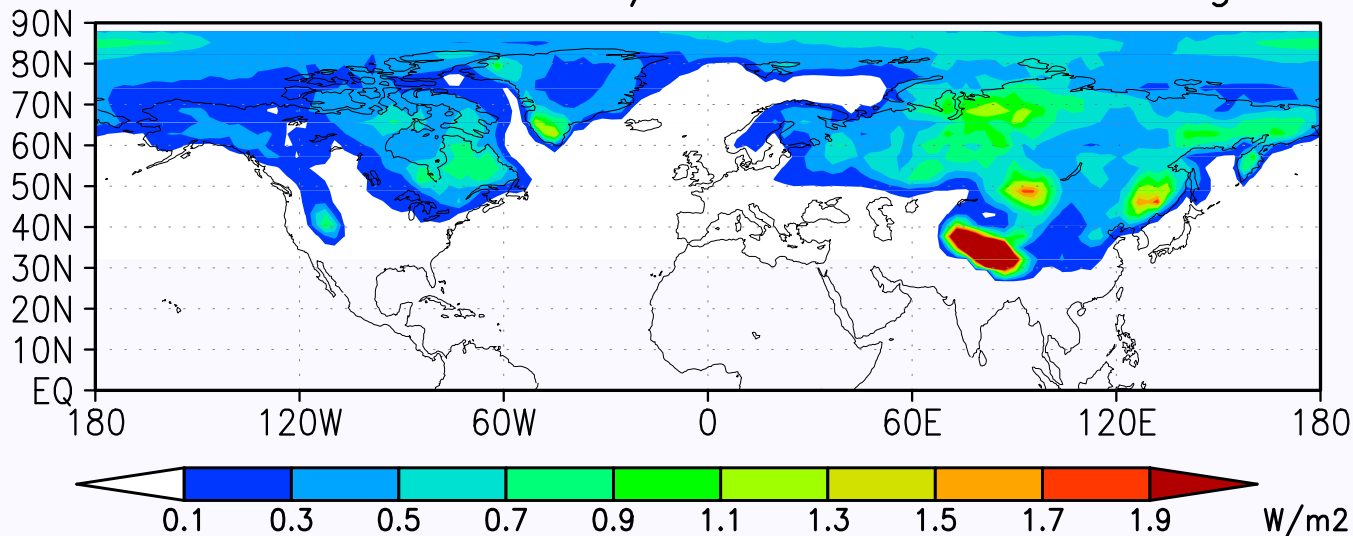
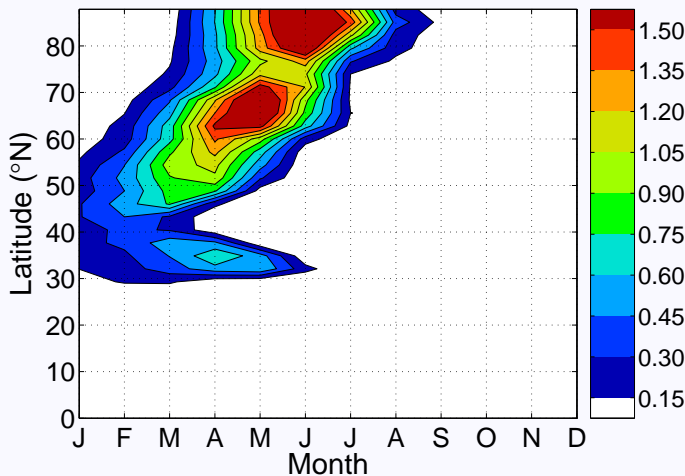
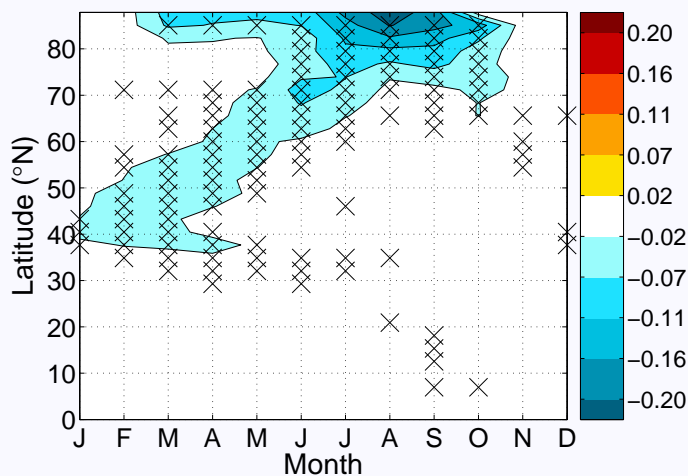


Figure 10: Present day global surface forcing by soot and dust in snowpack. Global mean  $\sim 0.056 \text{ W m}^{-2}$ .

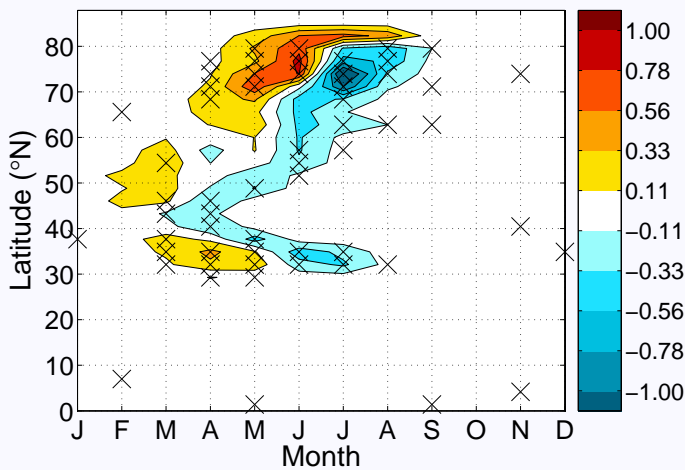
Present BC+Dust Snow Forcing ( $W m^{-2}$ )



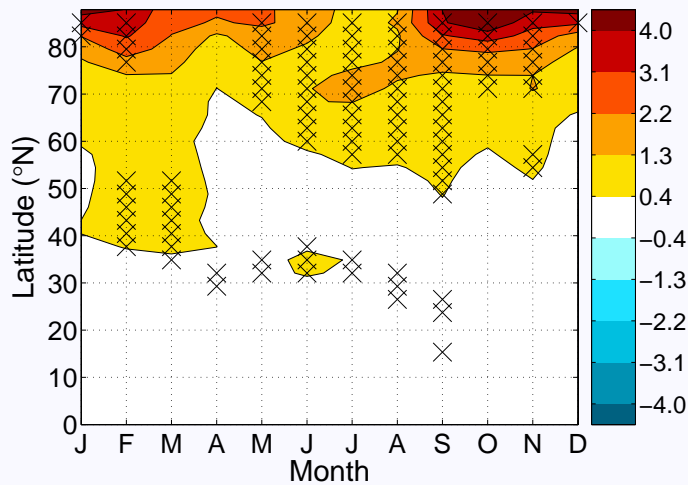
Pres. BC+Dust/Snow ALBS Change



Pres. BC+Dust/Snow QMELT Change ( $mm day^{-1}$ )



Pres. BC+Dust/Snow Temp. Change ( $^{\circ}C$ )



# LGM BC/Snow Surface Concentration

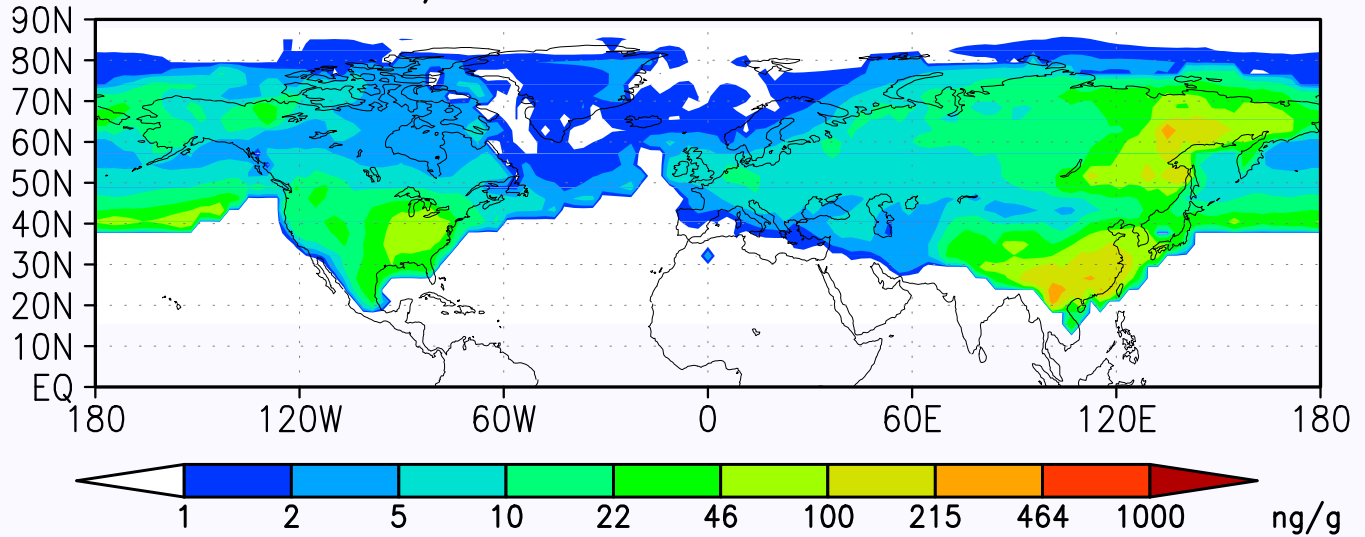


Figure 12: LGM soot concentration [ $\text{ng g}^{-1}$ ] in snowpack, 1998 fire emissions from non-ice regions. Global mean  $\sim 1.5 \text{ ng g}^{-1}$ .



# LGM Dust/Snow Surface Concentration

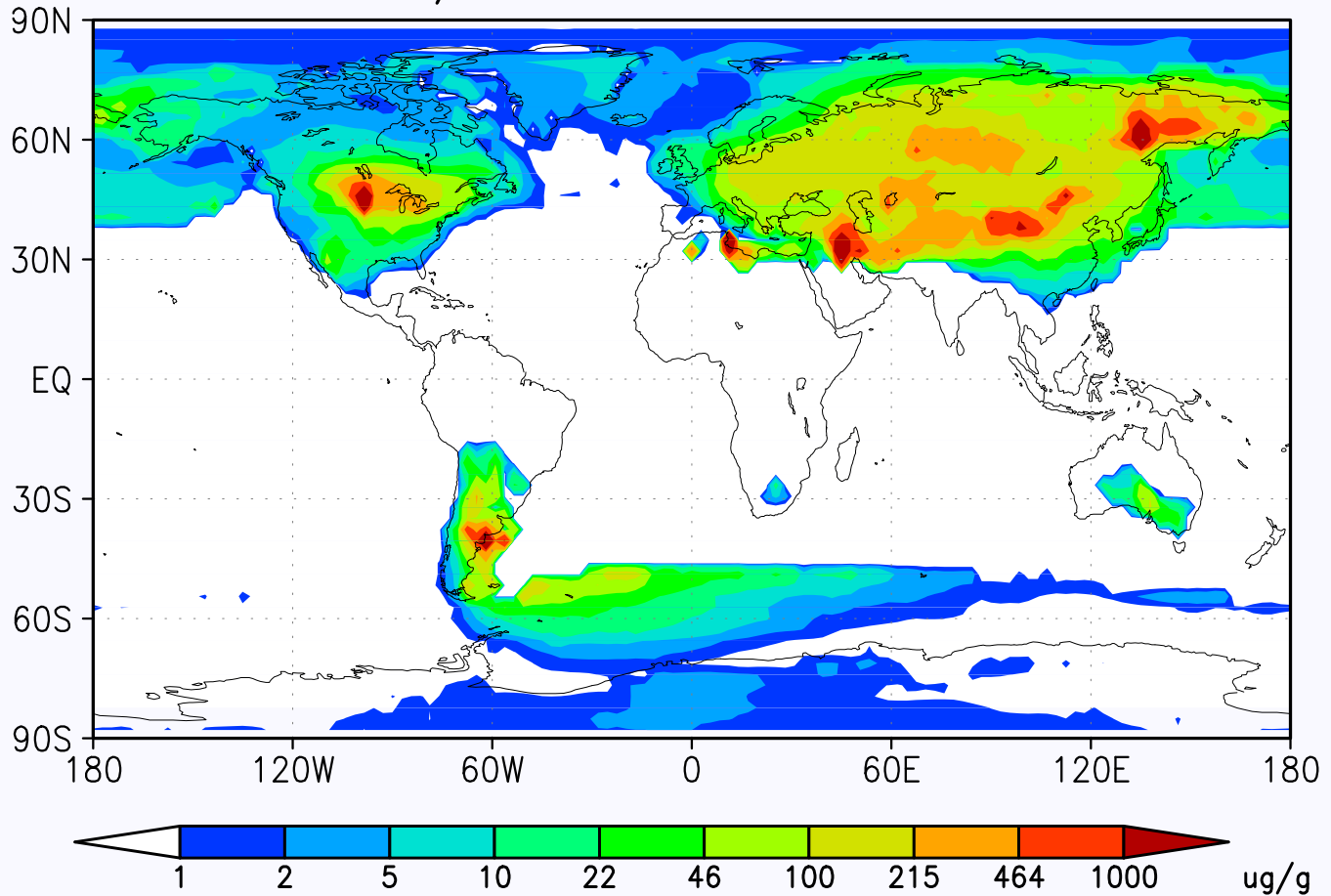


Figure 13: LGM dust concentration [ $\mu\text{g g}^{-1}$ ] in snowpack using Glaciogenic sources inferred by Mahowald et al., 2006, *JGR*. Global mean  $\sim 6.9 \mu\text{g g}^{-1}$ .

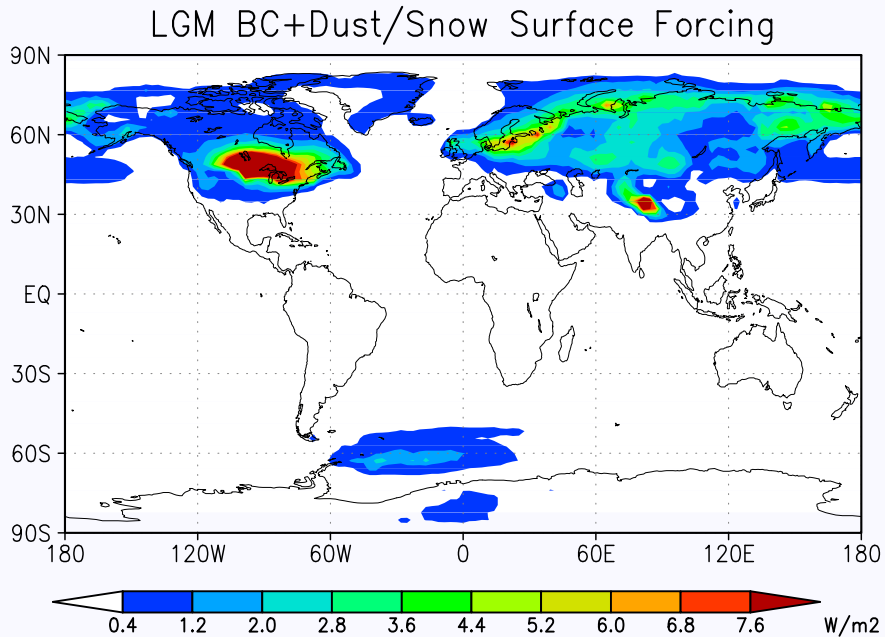
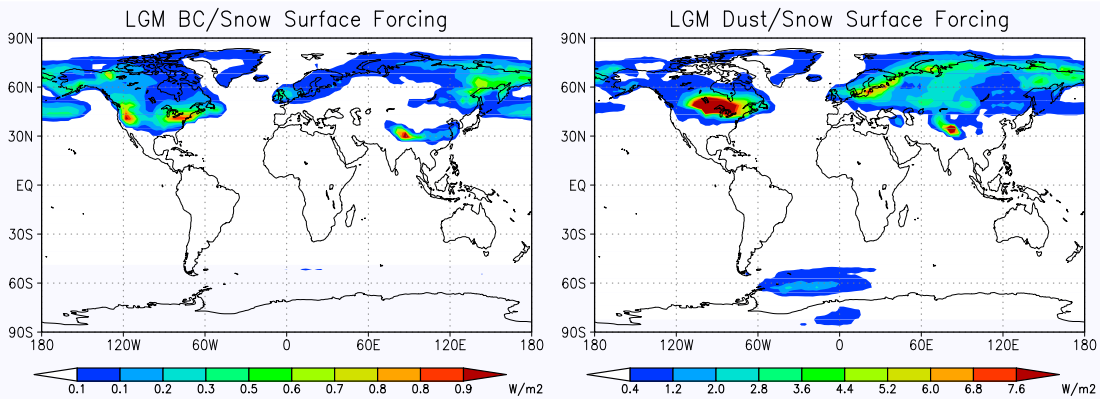
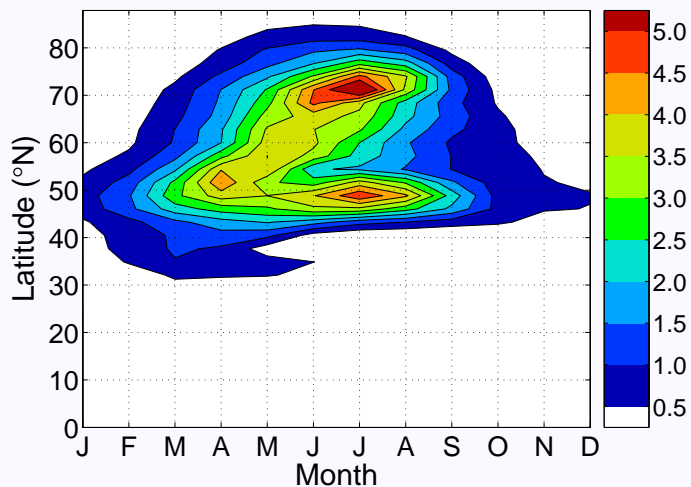
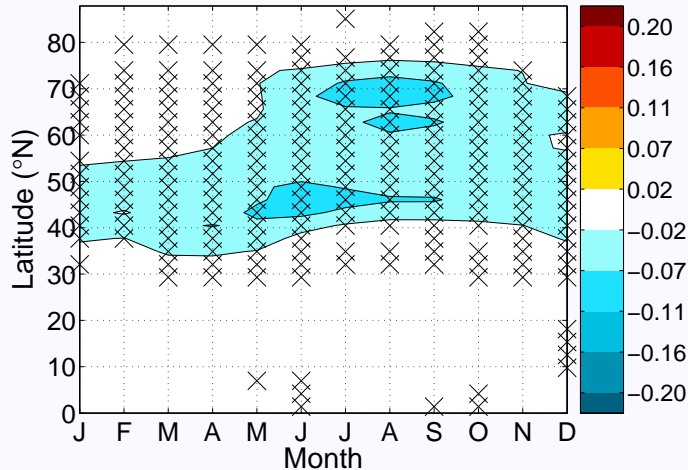


Figure 14: LGM forcing by soot and dust in snowpack.

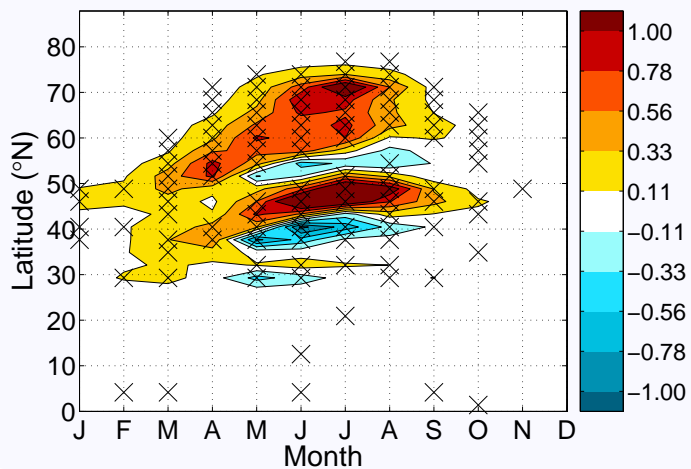
LGM BC+Dust Snow Forcing ( $W m^{-2}$ )



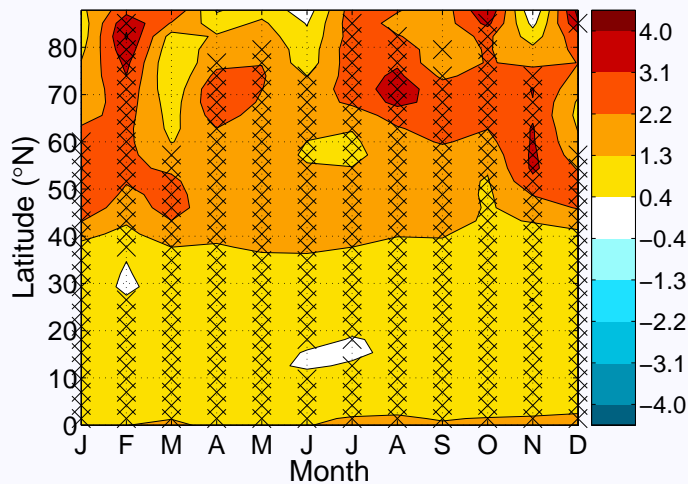
LGM BC+Dust/Snow ALBS Change



LGM BC+Dust/Snow QMELT Change ( $mm day^{-1}$ )



LGM BC+Dust/Snow Temp. Change (°C)



### 3. Conclusions

Current climate:

- Dust-snowpack forcing  $<$  soot forcing, dust efficacy  $>$  soot efficacy
- Dust+Soot-snowpack forcing  $0.050 \text{ W m}^{-2}$  warms climate  $0.16 \text{ K}$  ( $\sim 50\%$  by anthropogenic soot) with efficacy  $E \sim 4.2 \times \text{CO}_2$
- Significant climate affects on melt seasonality, Arctic  $T$

LGM climate:

- Dust-snow forcing increases ten-fold, causes  $90\%$  of  $0.95 \text{ K}$  warming
- Sustained ice-sheet darkening can (did?) change global climate

Overall:

- Introducing new forcing agents (dust) and feedback processes in snowpack model (subsurface melt, snow aging) dramatically increases forcing efficacy by triggering snow-ice-albedo feedbacks
- Polar amplification mechanisms are stronger than previously thought