# Gravity and the Geoid in the Nepal Himalaya

NAGW-2704



Semi Annual Report to NASA Washington

1 January 1992

Roger Bilham University of Colorado Boulder CO 80309-0216 303 492 6189

(NASA-CR-189523) GRAVITY AND THE GEDID IN N92-14560 THE NEPAL HIMALAYA Semiannual Report (Colorado Univ.) 41 p CSCL 08E Unclas

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#### Uplift and erosion in the Himalaya

Materials within the Himalaya are rising due to convergence between India and Asia. If the rate of erosion is comparable to the rate of uplift the mean surface elevation will remain constant. Any slight imbalance in these two processes will lead to growth or attrition of the Himalaya.

The process of uplift of materials within the Himalaya coupled with surface erosion is similar to the advance of a glacier into a region of melting. If the melting rate exceeds the rate of downhill motion of the glacier then the terminus of the glacier will receed up-valley despite the downhill motion of the bulk of the glacier. Thus although buried rocks, minerals and surface control points in the Himalaya are undoubtably rising, the growth or collapse of the Himalaya depends on the erosion rate which is invisible to geodetic measurements.

Erosion rates are currently estimated from suspended sediment loads in rivers in the Himalaya. These typically underestimate the real erosion rate since bed-load is not measured during times of heavy flood, and it is difficult to integrate widely varying suspended load measurements over many years. An alternative way to measure erosion rate is to measure the rate of change of gravity in a region of uplift. If a control point moves vertically it should be accompanied by a reduction in gravity as the point moves away from the Earth's center of mass. There is a difference in the change of gravity between uplift with and without erosion corresponding to the difference between the free-air gradient and the gradient in the acceleration due to gravity caused by a corresponding thickness of rock. Essentially gravity should change precisely in accord with a change in elevation of the point in a free-air gradient if erosion equals uplift rate.

We were funded by NASA to undertake a measurement of absolute gravity simultaneously with measurements of GPS height within the Himalaya. Absolute gravity is estimated from the change in velocity per unit distance of a falling corner-cube in a vacuum. Time is measured with an atomic clock and the unit distance corresponds to the wavelength an iodine stabilised laser. Since both these are known in an absolute sense to 1 part in  $10^{10}$  it is possible to estimate gravity with a precision of 0.1 µgal. Known systematic errors reduce the measurement to an absolute uncertainty of 6 µgal. The free air gradient at the point of measurement is typically about 3 µgals/cm. At Simikot where our experiment was conducted we determined a vertical gravity gradient of 4.4 µgals/cm.

The accompanying report records the experiment that we undertook in the Himalaya in 1991. The site description is provided together with a description of the instrument. The measured value of gravity at Nagarkot is 978494834.7 $\pm$ 6.7 µgals. It is our intention to remeasure this point in 1993 or 1994.

Publications and reports:

- Winester, D., J. Fried, B. Bernard, L. Shrestha, B. N. Shrestha, G. Adiga, R. Bilham and J. Faller (1990)) Absolute Gravity at Nagarkot Geodetic Observatory. pp.30. Archives of His Majesties Government of Nepal, Survey Department.
- Jackson, M., S. Barrientos, J. Behr, B. Bernard, R. Bilham, P. Bodin, G. Chitrakar, R. DeConto, L. Denham, J. Faller, J. Fried, D. Kauffman, D. Kayastha, P. Molnar, J. Normandeau, G. Peter, B. Phuyal, T. Pradhananga, B. Sharma, B. Shrestha, K.Shrestha, F. Sigmundsson, B. Stephens, B. Washburn, Wang Wenying, D. Winister, Zhao Guogang, Trans-Himalayan Geodesy, (1991). Eos Trans. Amer. Geophys. Un. 72, 44, 112
- Adhikari, K, R Bilham, M Jackson, N Karki, Kayastha, B Phuyal, T Pradhananga, B Sharma, B Shrestha, K Shrestha (1991). Interseismic Himalayan Subsidense: Uplift of Everest, *Eos Trans. Amer. Geophys. Un.* 72, 44, 497.

# ABSOLUTE GRAVITY

Nagarkot Geodetic Observatory, Nepal

March/April 1991

Observations, corrections and results. Gravity ties to Kathmandu and Simira airports.

Dan Winester, Jack Fried and Brent Bernard National Geodetic Survey, Rockville Md Laxman Shrestha, Buddhi N. Shrestha and Gajanan Adiga HMG Survey Department, Dilli Bazar, Nepal Roger Bilham and Jim Faller University of Colorado, Boulder, CO, 80309

# ABSOLUTE GRAVITY, Nagarkot, Nepal 1991

NGS Rockville Md: Survey of Nepal: Coordinated by: Dan Winester, Jack Fried and Brent Bernard Laxman Shrestha and Gajanan Adiga Roger Bilham, Jim Faller and Buddhi N. Shrestha

#### Summary of measurements

The purpose of measuring absolute gravity in the Himalaya was to establish a reference datum for the local gravity network in Nepal and to establish points that may be remeasured to reveal changes of elevation in future years. The original plan was to measure absolute gravity at three locations: in the Greater Himalaya, in the Lesser Himalaya and in the Terrai bordering the northern plains of India. Each absolute gravity point was scheduled to be co-located with a GPS control point so that an independent estimate of vertical deformation might be possible.

The plan we adopted differed in three ways from the above:

1) One absolute-g site only was measured at Nagarkot (FAGS-1). The corrected value of the FAGS-1 indoor point at ground level for the period 3/30/91-4/2/91 is  $978494834.7\pm6.7$  µgal. The gravity gradient at floor level (zero to 0.43m) was 4.4194 µgal/cm.

2) Relative ties were made to three GPS points: Nagarkot, Kathmandu airport and Simira Airport. The relative differences from FAGS-1 to these points are listed on the next page.

The ties were undertaken using a pair of Model D LaCoste Romberg meters. For Nagarkot the GPS point is less than 10 m from the brick building where GPS measurements were made. The Kathmandu Airport tie was undertaken using road transport (multiple ties over the 33-km-long 1.5 hour road linking Nagarkot to the capital). The Simira tie was made by flying several times between Simira and Kathmandu. The Model D gravimeter has just sufficient range to accommodate the gravity variation associated with the vertical change in height between Nagarkot and Kathmandu, and also the latitude change and vertical range combination between Kathmandu and Simira.

3) The limited number of sites suitable for gravity measurements has resulted in no gravity measurements at points suspected to be rising in the Greater and Lesser Himalaya. Simira is south of the Lesser Himalaya and Kathmandu and Nagarkot lie between the Lesser and the Greater Himalaya. Future Model D or Model G gravimeter ties be made from Kathmandu airport to GPS points elsewhere in Nepal are needed to correct this limitation in the 1991 measurements.

A removal truck was used to meet the several hundred pounds of equipment from the plane and to store the packaging at Nagarkot. The power at Nagarkot was found to be unreliable for the gravity measurements as was the portable generator used to provide backup power. Measurements for this reason were spread over a longer period than is usual. Air conditioning was requested for the gravimeter but was found to be unnecessary in Nagarkot. A decision to occupy only one point "absolutely" and the other points using Model D gravimeters was made because:

a) the absolute gravimeter was damaged in transit to Kathmandu or on the road to Nagarkot and might have further been damaged by additional road transport.

b) suitable temperature control from air conditioners was unavailable at the other selected sites, and an air conditioner would have had to have been trucked in from India together with a 15 kw generator.

c) Power outages at Nagarkot reduced the time available for measurements at additional sites.

The new gravity base stations provide a framework for the local Nepal gravity network. It is anticipated that future gravity measurements will extend this network throughout the country. The absolute accuracy of the 1991 measurements is  $\pm 6 \mu$ gals or approximately  $\pm 1.5$  cm in elevation.

Funding support for the measurements was provided by NASA grant NAGW-2704. A description of the JILA absolute gravimeter follows the observational data.



UNITED STATES DEPARTMENT OF COMMERCE National Oceanic and Atmospheric Administration NATIONAL OCEAN SERVICE

CALLE OF CHARTING AND CEODETIC CERVICES ROCKVILLE, MARYLAND 20852

Coast and Geodetic Survey 11 June 1991

Dr. Roger Bilham CIRES. Univ. of Colorado Boulder. 30 80309

Dear Roger:

Enclosed are gravity base station descriptions for occupied sites in Nepal. A copy of these will be sent to Buddhi Shrestha. The NAGARKOT FAGS-1 absolute gravity value will be available from Dr. Feter. The gradients at NAGARKOT FAGS-1 from floor to 33 cm is 0.44194 mgal/m and from floor to 120 cm is 0.43923 mgal/m. Relative to the floor value at NAGARKOT FAGS-1 at the following gravity transfers:

> NAGARKOT GPS KATHMANDU J SIMARA J SIMARA GPS

- 0.691 ± 0.002 mgals +166.469 ± 0.005 +368.599 ± 0.017 +368.706 ± 0.013

Sincerely.

Čaniel Winester. Geodesist National Geodetic Survey, N/CG 161N

















## JILA #4 ABSOLUTE GRAVITY DETERMINATION

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Add water table correction Add polar motion correction Add laser drift correction Add laser-head temp. correct	tion .	<u>_</u>
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PAGE 3 OF

COMMENTS

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PAGE 4 OF

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_ Strould be uniform for first 6 sets, we will
continue to update.
_ Cravity readings from Last night Are excellent
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tolerances (10:59 Local). Skies starting to clear.
_ Sets proceeding without incident.
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frequency / Volt meter to Topazis to watch
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And equipment if rucessary. UPS voltage @ 1055 117VAC
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whole system crashed. I turned all even thing.
ups to pump still Running. Got to go 0
Next Page 0

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COMMENTS NEORO KOT CONTRA CONTIN
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Re-initiallized system. Smiled DDT: Data for set #10
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that it was all only I have and its temp is
atready stable - the decreased current is the
expected Response to the indrease in temp.
Will begin next set on original observation
schedule. SET # 11 Power Failure AGRIN While
Rodace & I were attempting to wire up and
stort the germator is the everit of a power
failure. To LATE NATOR! System down Again.
Generator now Rupping, system powered up
waiting to relock spring and Losen: Generator
will Require Republing even 2 to 3 hours, Current
PIAN is to operate until we have 18 sets (19)
is the data Looks good - very good and
the problems pre many.
Concration Palueling schedule: 5:30P
<u> </u>
lite carosen/te 10:30P
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- 1 0 03:30 <b>P</b>
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PAGE 6 OF

COMMENTS NAGRAKAT Observatory

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ABSOLUTE GRAVITY STATION ORIENTATION DIAGRAM



DIRECTOR GENERAL HMG SURVEY DEPT



Final Processed set

MEANS.

1

3 STD errors For error bars

Scatter due to Field operators have trouble setting system correctly with damaged interferometer.



# time (2.63 days) AVG. CORR. = 4.3



time ( 2.65 days )



time ( 2.65 days )



time ( 2.65 days )

#### NGS ABSOLUTE GRAVITY OBSERVATIONS From aakath91.089 This drop set has been previously processed for: three sigma acceptance limit gravitational tide correction

#### DROP SET MEANS SUMMARY

d s	rop et	num of drops	laser mode	mean date/time	mean grav (ugal)	sd mean (ugal)	sd obs (ugal)
#	1	250	RED	910330152056	978 494 428.8	.9	14.6
#	2	247	BLUE	910330172055	978 494 409.3	.5	8.6
#	3	249	RED	910330192056	978 494 427.7	.6	9.0
#	4	248	BLUE	910330212101	978 494 404.1	.5	8.1
#	5	249	RED	910330232055	978 494 423.6	.5	8.0
#	6	246	BLUE	910331012055	978 494 402.7	.6	9.7
#	7	248	RED	910331032056	978 494 424.7	.9	13.7
#	8	234	RED	910331072101	978 494 425.9	.9	13.6
#	9	234	RED	910401142106	978 494 416.2	1.1	16.1
#	10	237	BLUE	910401162101	978 494 397.7	.7	10.2
#	11	250	RED	910401182100	978 494 415.6	.5	8.2
#	12	248	BLUE	910401202055	978 494 395.1	.5	8.3
#	13	247	RED	910401222055	978 494 413.2	. 4	6.8
#	14	236	BLUE	910402002100	978 494 394.3	.6	8.9
#	15	242	RED	910402022055	978 494 418.4	.6	9.7
#	16	246	BLUE	910402042110	978 494 399.9	1.0	16.2
#	17	215	RED	910402062136	978 494 419.8	2.0	28.6

10 dropsets weighted mean of red mode observations = 4 420.15.77 dropsets weighted mean of blue mode observations = 4 400.75.3average of weighted red and blue means = 4 410.4

average standard deviation of observation = 11.7

#### NGS ABSOLUTE GRAVITY OBSERVATIONS From aakath91.089 This drop set has been previously processed for: three sigma acceptance limit gravitational tide correction

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#### DROP SET MEANS SUMMARY OFFSET CORRECTED

d s	rop et	num of drops	la <b>ser</b> mode	mean date/time	mean grav (ugal)	re (	sidual ugal)
#	1	250	RED	91033015205	6 978 494	419.1	8.7
#	2	247	BLUE	91033017205	5 978 494	418.9	8.6
#	3	249	RED	91033019205	6 978 494	418.1	7.7
#	4	248	BLUE	91033021210	1 978 494	413.8	3.4
#	5	249	RED	91033023205	5 978 494	413.9	3.5
#	6	246	BLUE	91033101205	5 978 494	412.4	2.0
#	7	248	RED	91033103205	6 978 494	415.1	4.7
#	8	234	RED	91033107210	1 978 494	416.3	5.9
#	9	234	RED	91040114210	6 978 494	406.5	-3.9
#	10	237	BLUE	91040116210	1 978 494	407.4	-3.0
#	11	250	RED	91040118210	0 978 494	405.9	-4.5
#	12	248	BLUE	91040120205	5 978 494	404.8	-5.6
#	13	247	RED	91040122205	5 978 494	403.5	-6.9
#	14	236	BLUE	910402002100	0 978 494	404.0	-6.4
#	15	242	RED	91040202205	5 978 494	408.7	-1.7
#	16	246	BLUE	910402042110	978 494	409.6	8
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average of weighted red and blue means = 4 410.4 s.d. mean = 5.5average standard deviation of observation = 11.7 NGS ABSOLUTE GRAVITY OBSERVATIONS From aakath91.089 This drop set has been previously processed for: gravitational tide correction

#### DROP SET MEANS SUMMARY OFFSET CORRECTED

d s	rop et	num of drops	laser mode	mean date/time	mean grav (ugal)	res (l	sidual 1gal)
#	1	250	RED	91033015205	5 978 494	419.7	8.8
#	2	250	BLUE	91033017205	5 978 494	418.7	7.8
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#	4	250	BLUE	91033021210	1 978 494	413.4	2.6
#	5	250	RED	91033023205	5 978 494	414.4	3.6
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#	9	250	RED	91040114210	5 978 494	403.9	-6.9
#	10	250	BLUE	91040116210	1 978 494	406.9	-4.0
#	11	250	RED	910401182100	978 494	406.5	-4.3
#	12	250	BLUE	910401202055	5 978 494	404.5	-6.4
#	13	250	RED	910401222055	5 978 494	403.8	-7.0
#	14	250	BLUE	910402002100	978 494	403.2	-7.6
#	15	250	RED	910402022055	5 978 494	408.8	-2.0
#	16	250	BLUE	910402042110	978 494	410.1	8
#	17	250	RED	910402062136	5 978 494	411.8	.9

average of weighted red and blue means = 4 410.9 s.d. mean = 5.8average standard deviation of observation = 20.1

#### NGS ABSOLUTE GRAVITY OBSERVATIONS From aakath91.089 This drop set has been previously processed for: three sigma acceptance limit gravitational tide correction local atmospheric pressure correction

#### DROP SET MEANS SUMMARY

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d s	rop et	num of drops	laser mode	mean date/time	mean grav (ugal)	sd mean (ugal)	sd obs (ugal)
#	1	250	RED	910330152056	978 494 430.6	.9	14.6
#	2	247	BLUE	910330172055	978 494 411.0	.5	8.6
#	3	249	RED	910330192056	978 494 429.3	.6	9.0
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#	5	249	RED	910330232055	978 494 425.0	.5	8.0
#	6	246	BLUE	910331012055	978 494 404.4	.6	9.7
#	7	248	RED	910331032056	978 494 426.8	.9	13.7
#	8	234	RED	910331072101	978 494 427.6	.9	13.6
#	9	234	RED	910401142106	978 494 417.2	1.1	16.1
#	10	237	BLUE	910401162101	978 494 399.0	.7	10.2
#	11	250	RED	910401182100	978 494 416.9	.5	8.2
#	12	248	BLUE	910401202055	978 494 396.3	.5	8.3
#	13	247	RED	910401222055	978 494 414.2	.4	6.8
#	14	236	BLUE	910402002100	978 494 395.6	.6	8.9
#	15 ່	242	RED	910402022055	978 494 420.1	.6	9.7
#	16	246	BLUE	910402042110	978 494 401.8	1.0	16.2
#	17	215	RED	910402062136	978 494 421.4	2.0	28.6

10 dropsets weighted mean of red mode observations = 4 421.55.97 dropsets weighted mean of blue mode observations = 4 402.15.5average of weighted red and blue means = 4 411.85.7

average standard deviation of observation = 11.7

NGS ABSOLUTE GRAVITY OBSERVATIONS From aakath91.089 This drop set has been previously processed for: three sigma acceptance limit gravitational tide correction local atmospheric pressure correction

#### DROP SET MEANS SUMMARY OFFSET CORRECTED

d: se	rop et	num of drops	laser mode	mean date/time	mean grav (ugal)	res (u	sidual 1gal)
#	1	250	RED	91033015205	6 978 494	420.9	9.1
#	2	247	BLUE	91033017205	5 978 494	420.7	8.9
#	3	249	RED	91033019205	6 978 494	419.6	7.8
#	4	248	BLUE	91033021210	1 978 494	415.1	3.3
#	5	249	RED	91033023205	5 978 494	415.3	3.5
#	6	246	BLUE	91033101205	5 978 494	414.0	2.2
#	7	248	RED	910331032050	5 978 494	417.2	5.4
#	8	234	RED	91033107210	L 978 494	417.9	6.1
#	9	234	RED	910401142100	5 978 494	407.5	-4.3
#	10	237	BLUE	91040116210	L 978 494	408.6	-3.2
#	11	250	RED	910401182100	978 494	407.2	-4.6
#	12	248	BLUE	91040120205	5 978 494	406.0	-5.8
#	13	247	RED	910401222055	5 978 494	404.5	-7.3
#	14	236	BLUE	910402002100	978 494	405.3	-6.6
#	15	242	RED	910402022055	5 978 494	410.4	-1.4
#	16	246	BLUE	910402042110	978 494	411.5	3
#	17	215	RED	910402062136	5 978 494	411.8	.0

average of weighted red and blue means = 4 411.8 s.d. mean = 5.7 average standard deviation of observation = 11.7 NGS ABSOLUTE GRAVITY OBSERVATIONS From aakath91.089 This drop set has been previously processed for: three sigma acceptance limit gravitational tide correction local atmospheric pressure correction ocean loading correction

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### DROP SET MEANS SUMMARY OFFSET CORRECTED

S	lrop set	num of drops	laser mode	mean date/time	mean grav (ugal)	res (u	sidual Igal)
#	1	250	RED	910330152056	978 494	419.6	7.7
#	2	247	BLUE	910330172055	978 494	419.5	7.6
#	3	249	RED	910330192056	978 494	419.4	7.4
#	4	248	BLUE	910330212101	978 494	416.2	4.3
#	5	249	RED	910330232055	978 494	416.6	4.6
#	6	246	BLUE	910331012055	978 494	414.6	2.6
#	7	248	RED	910331032056	978 494	416.5	4.5
#	8	234	RED	910331072101	978 494	417.5	5.5
#	9	234	RED	910401142106	978 494	407.2	-4.7
#	10	237	BLUE	910401162101	978 494	407.4	-4.5
#	11	250	RED	910401182100	978 494	405.9	-6.0
#	12	248	BLUE	910401202055	978 494	406.0	-6.0
#	13	247	RED	910401222055	978 494	405.7	-6.3
#	14	236	BLUE	910402002100	978 494	406.8	-5.1
#	15	242	RED	910402022055	978 494	411.0	-1.0
#	16	246	BLUE	910402042110	978 494	410.9	-1.1
#	17	215	RED	910402062136	978 494	410.6	-1.4

average of weighted red and blue means = 4 412.0 s.d. mean = 5.5average standard deviation of observation = 11.7

#### ABSOLUTE GRAVITY: A RECONNAISSANCE TOOL FOR STUDYING VERTICAL CRUSTAL MOTIONS

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Abstract. A major effort is under way to develop highly portable absolute gravimeters having an ultimate accuracy of  $3-5~\mu$ Gal, an accuracy which translates into a height sensitivity of several centimeters. We are just finishing the construction of six such units. Measurements at the Joint Institute for Laboratory Astrophysics with one of these new instruments agree well with the earlier measurements made in 1981 and 1982 with a previous generation instrument. Recent measurements at the International Bureau of Weights and Measures in Sevres, France, as a part of an international intercomparison of absolute gravimeters, also show good agreement with the other instruments.

Measurement of the absolute value of the freefall acceleration "g" has long been a matter of scientific interest. Present-day methods of measuring the absolute value of g employ ballistic systems involving either direct free-fall or symmetrical rise-and-fall methods. The earliest such measurements employed the direct free-fall method and geometrical optics to determine the position of the dropped object as a function of time. More recently, laser interferometry has been used almost exclusively.

A major effort to develop a new generation of high-precision absolute gravimeters is in the final stages at the Joint Institute for Lahoratory Astrophysics (JILA) located at the University of Colorado in Boulder, Colorado. These gravimeters interferometrically measure the position of a free-falling object as a function of time and thereby permit the determination of the free-fall acceleration. This paper will discuss the use of absolute gravity for the study of vertical motions, the status of the JILA absolute gravity instruments, and the advantages and nearterm prospects of using them for this purpose.

Traditionally, vertical height information has been derived mainly from leveling data. However, even using automated leveling systems, the cost per kilometer is high, from \$350/km to rerun an existing line to between \$500 and \$600/km to run a new line (G. J. Mitchell, private communication, 1986). A number of extraterrestrial techniques and systems also exist for measuring vertical movements of the earth's surface such as laser satellite ranging, very long baseline interferometry, and using ground receivers together

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Paper number 585687. 0148-0227/86/0058-5687\$05.00 with the NAVSTAR global positioning system satellites. These methods are now capable of achieving the interesting accuracies of between 1 and 3 cm and are therefore likely to play an increasingly important role in determining vertical motions. Their costs are still high; but these costs, particularly those associated with the global positioning satellite system approach, should soon be lowered.

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Gravity measurements, both relative and absolute, given sufficient measurement precision, provide a comparatively inexpensive way to look for vertical crustal movements. A 1-cm vertical crustal motion would result in a gravity change of approximately 3 µGal were no change in the local mass distribution to occur. The actual change in gravity observed in connection with a 1-cm vertical displacement will generally be 2-3 µGal but can be outside this range for some crustal movement mechanisms [Jachens, 1978a,b]. To differentiate, however, between subsurface density changes and vertical height changes, one must use one of the geometrical geodetic systems. Gravity does, however, provide an excellent and low cost reconnaissance tool with which to gather large amounts of preliminary data which then, for those areas in which gravity changes are occurring, can be checked and interpreted in combination with the other (geometrical) vertical data. If vertical motions are subsequently confirmed by other means, the observed gravity changes can help to determine the mechanism responsible for the motions.

In using gravity measurements as a reconnaissance tool to look for vertical movements, absolute gravity measurements have a number of advantages over relative gravity measurements: the most important of these being that absolute gravity is a "point technique." A single measurement produces a gravity value, in some sense a measure of the distance from the center of the earth, which depends only on the basic standards of length and time. Relative gravity measurements (as well as conventional leveling techniques) must necessarily be tied to a (presumed) stable external reference point which complicates the measurement process and inevitably raises questions about the stability of that reference point over the appropriate time frame. In a relative gravimeter (see, for example Clark [1984]), the spring, whose length is essentially the measured parameter, displays secular creep as well as episodic changes in its length. Vibrations encountered while transporting these devices and stresses due to clamping only serve to exacerbate these problems. In addition, nonlinearities in the adjusting screw and its associated lever reduction mechanism have to be carefully calibrated if their effects are to be removed. In practice, the measurement precision depends on the par-

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Fig. 1. Block diagram of free-fall method.

ticular instrument used, the station-to-station distance, and also on the gravity difference. Without special precautions, relative gravimeters typically reach precisions of between  $\pm 30$  and  $\pm 100 \ \mu$ Gal for a single measurement of a given difference in gravity. Extreme care is required to reduce this error to the  $\pm 5$  to  $\pm 10 \ \mu$ Gal range [Torge, 1985].

By contrast, the accuracy of absolute freefall instruments depends mainly on the reproducibility of the basic standards of length and time, and a stabilized laser provides the length standard and an atomic (rubidium) clock provides the time standard. The absolute wavelength of the laser and the frequency of the atomic clock can easily be measured directly in the laboratory. The drifts in these "standards" are low enough so that they can be used for months with negligible error contributions at the parts in 109 level of accuracy. Further, these "standards" are less subject to the ordinary vibration in transit, environmental temperature, etc., problems which have proven difficult with traditional relative gravimeters at the microGal level of sensitivity.

Modern-day absolute gravity instruments have been developed and improved over the past 30 years through the utilization of available

technology. In practice, they all measure the position of a freely falling mass as a function of time (with exquisite sensitivity) and from that motion determine the value of g (Figure 1). Two types of free-fall instruments have been developed: the first utilizes simple free fall, and the second uses an up-and-down trajectory [Faller and Sakuma, 1986). In each case, g is determined by fitting a quadratic expression to the measured trajectory. In practice, a Michelson-type laser interferometer is used to sense the position of the falling object during its fall. The dropped object contains a cube corner (a special type of optical mirror that reflects the laser directly back, independent of the cube's exact orientation). The occurrence times of the zero crossing of the fringes then provide the necessary information with which to calculate g.

The first laser interferometric g measurements were made in 1962 by J. E. Faller using an early commercially available He-Ne laser in what had been designed as a white-light-fringe g apparatus. The first portable laser interferometer absolute gravimeter was developed by J. A. Hammond and J. E. Faller at JILA and Wesleyan University with support from the Air Force Geophysics Laboratory (AFGL). With this apparatus, which had an accuracy of 50 µGal, data were taken at eight

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