

Laboratory study of the frictional rheology of sheared till

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Received 16 April 2007; revised 2 January 2008; accepted 21 February 2008; published 7 June 2008.

[1] Deformation of till produces power law creep for low strain at stresses high enough to cause permanent deformation but below the shear strength. Experiments were conducted on till (a mixed size granular material) from Matanuska Glacier, Alaska, and the Scioto (Ohio) Lobe of the Laurentide Ice Sheet (Caesar till).

(creep). Normal stress ranged from 50 kPa to 5 MPa at shearing rates ranging from 1 to 300 $\mu\text{m/s}$ for 1 cm thick samples. Fabric development within till layers was investigated by varying shear strain prior to creep tests.

With increasing initial strain and inferred fabric development, the creep strain rate decreases while n increases. Experiments at a normal stress of 1 MPa and no initial strain show $n = 6.8$, increasing to $n = 17.5$ at higher shear strains; however, strain rate was still decreasing and thus these values represent maximum estimates. Our data show that in the absence of dilatant hardening till exhibits rate sensitivity at strain of order 1 or less.

Citation: Rathbun, A. P., C. Marone, R. B. Alley, and S. Anandakrishnan (2008), Laboratory study of the frictional rheology of sheared till, *J. Geophys. Res.*, 113, F02020, doi:10.1029/2007JF000815.

1. Introduction

[2] Glaciers can move by a variety of processes, with deformation of subglacial till playing a major role for some valley glaciers, ice streams, and ice sheets [e.g., Paterson, 1994; Alley, 2000; Clarke, 2005]. Seismic studies have shown that the material under fast-moving ice streams is highly porous water-saturated sediment (till) [Alley et al., 1986; Blankenship et al., 1986; Peters et al., 2006]. The actively deforming layer, which ranges in scale between centimeters and several meters [e.g., Kamb, 1991; Engelhardt and Kamb, 1998; Truffer et al., 2000], is produced by reworking of sediment or by wear processes such as abrasion or plucking. At high effective stress (overburden stress minus pore water pressure), the shear strength of till is larger than the shearing resistance of ice, so deformation only occurs within the ice; at low effective stress, till is more likely to deform than ice [e.g., Engelhardt et al., 1990; Paterson, 1994]. Effective stresses may range from a few kPa or less to a few hundreds of kPa with channelized water drainage favoring higher values [Alley et al., 1989a; Engelhardt

et al., 1990; Paterson, 1994; Engelhardt and Kamb, 1997]. Low effective stress can also limit bed deformation by allowing sliding at the ice-bed interface [e.g., Kamb, 2001; Iverson et al., 2003].

[3] Early studies of bed deformation assumed a viscous or near-viscous relation for till deformation [e.g., Boulton and Hindmarsh, 1987]. Laboratory tests of till deformation are sparse [e.g., Tulaczyk, 2006], but have produced important results leading to the idea that till is more nearly a Coulomb plastic (treiboplastic) material [e.g., Kamb, 1991, 2001; Iverson et al., 1997, 1998; Tulaczyk et al., 2000]. However, debate still exists as to whether till is best characterized by a “viscous” rheology (with a nearly linear relation between shear stress and strain rate) or “plastic” rheology (for which shear stress is related to strain rate raised to a power of 10 or more) [e.g., Kamb, 1991; Jenson et al., 1995; Hindmarsh, 1997; Iverson et al., 1998; Alley, 2000; Tulaczyk et al., 2000; Fowler, 2003; Kavanaugh and Clarke, 2006].

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2. Terminology

[4] Technical terms differ between the various communities that study friction and the rheology of glacial and